

# Noise as a Source for Detecting Time-Lapse Changes

Roel Snieder

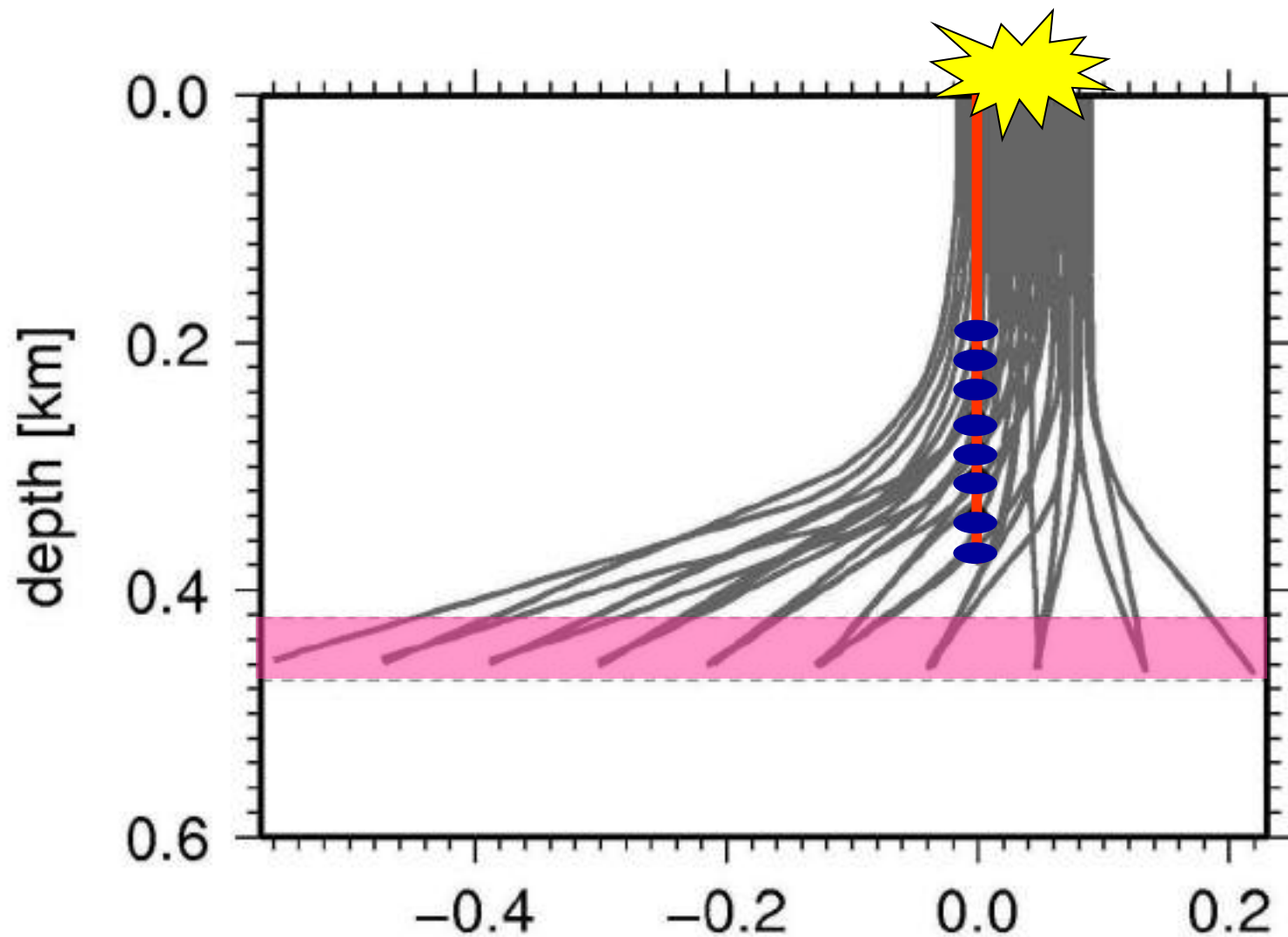
[rsnieder@mines.edu](mailto:rsnieder@mines.edu)

# Noise as a Source for Detecting Time-Lapse Changes

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Ichiro Kuroda, Xun Li, Masatoshi Miyazawa, Nori  
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Thomas Planès, Kaoru Sawazaki, John Scales,  
Christoph Sens-Schönfelder, Kasper van Wijk,  
Roel Snieder

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# Geometry of L08-pad



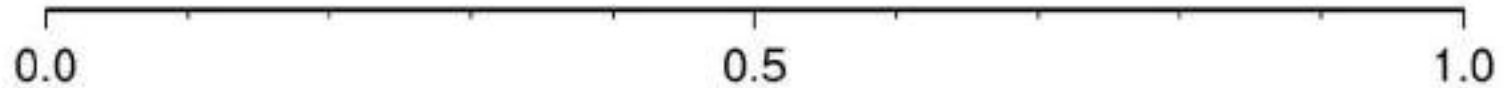
(Miyazawa, Snieder & Venkataraman, *Geophysics*, **73**, D35-D40, 2008)

# from Noise to Signal

Top station

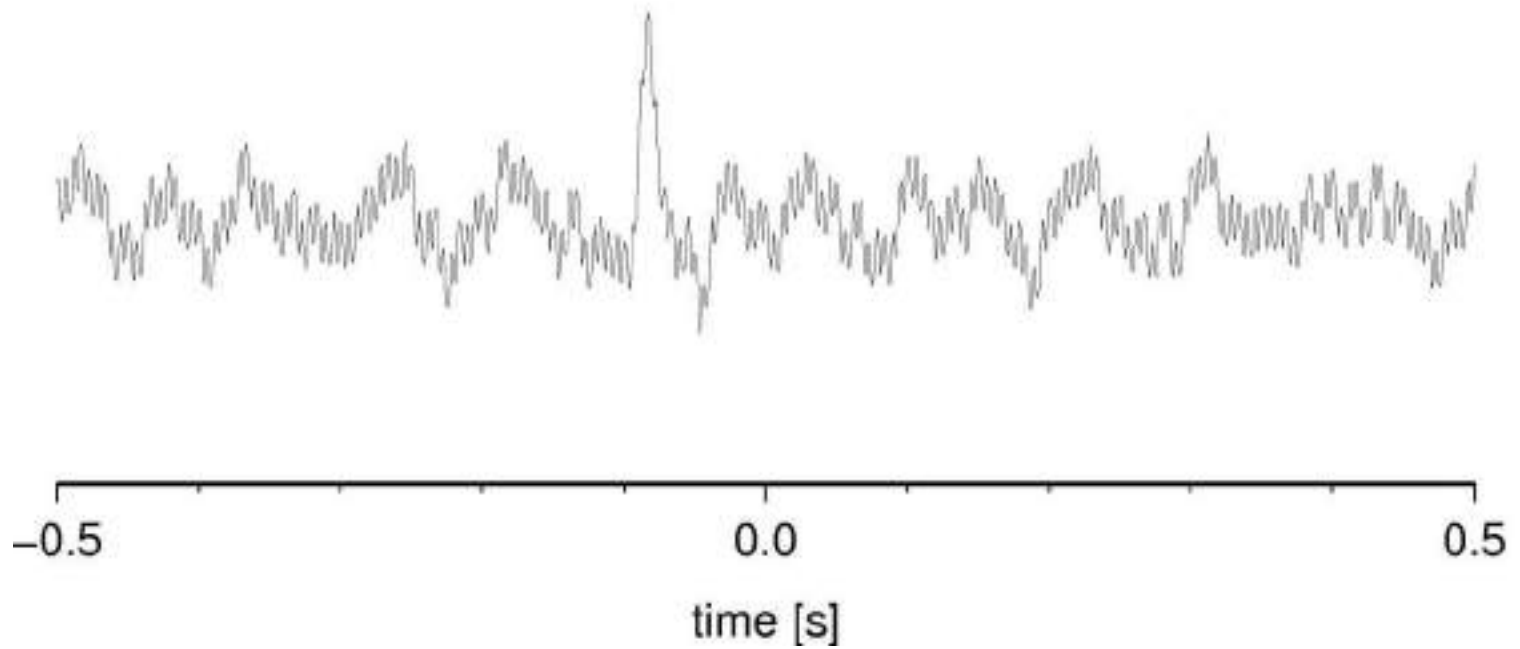


Bottom station

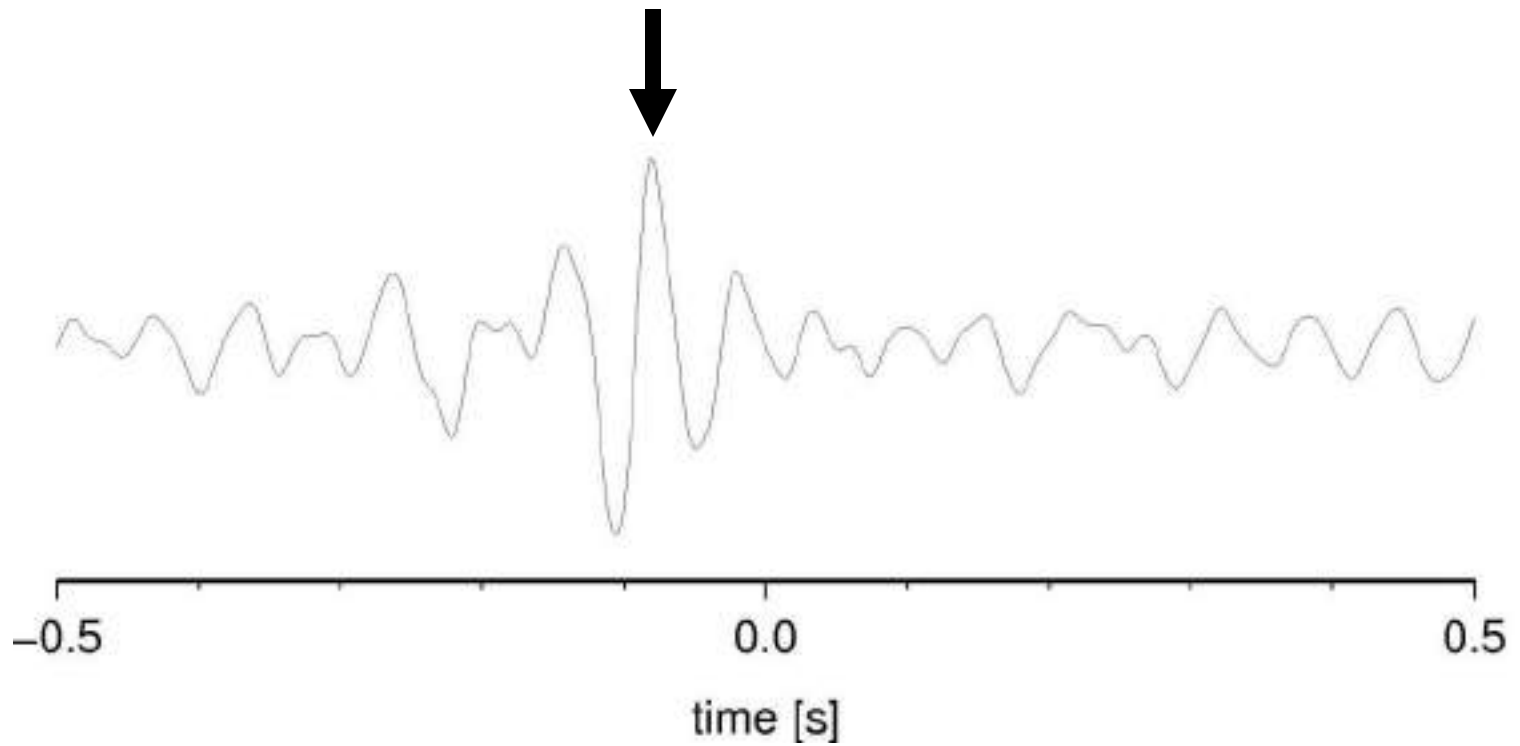


time [s]

# Cross-Correlation



# Cross-Correlation

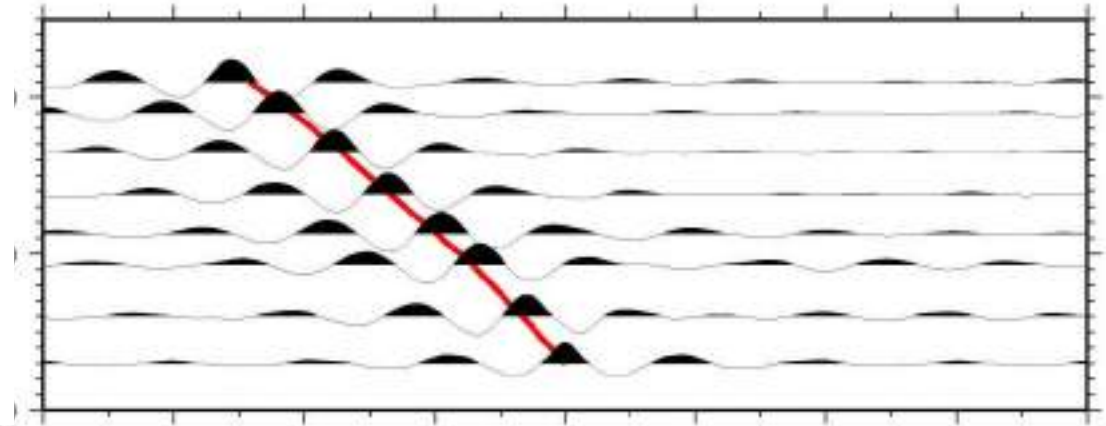
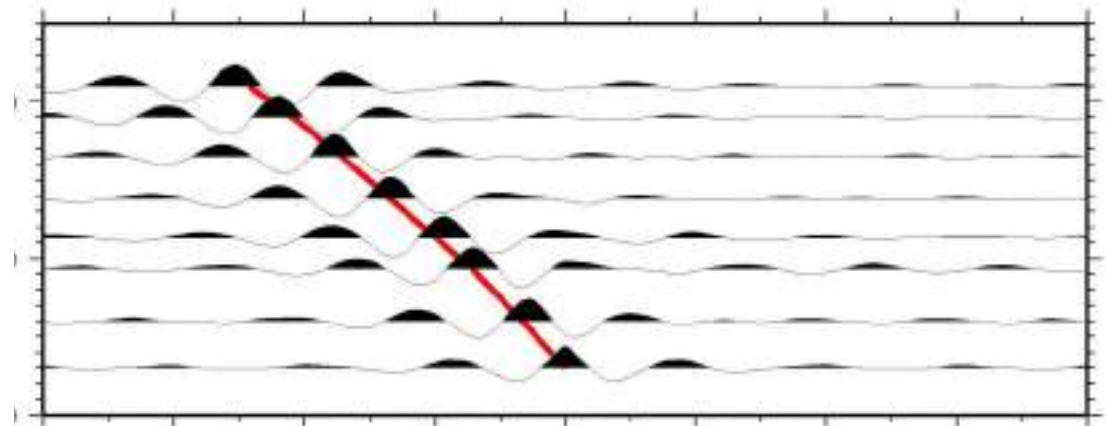
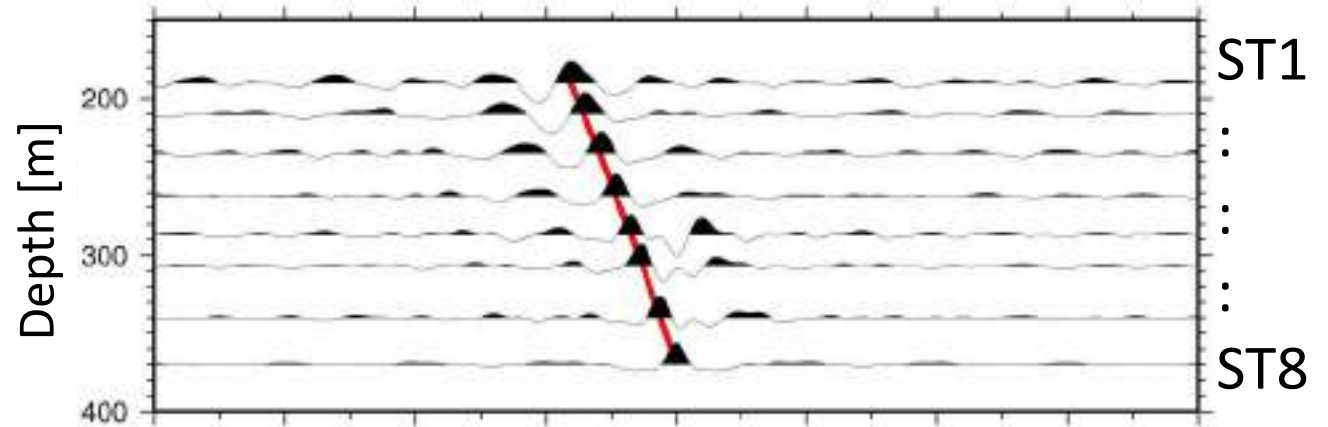


Cycle1

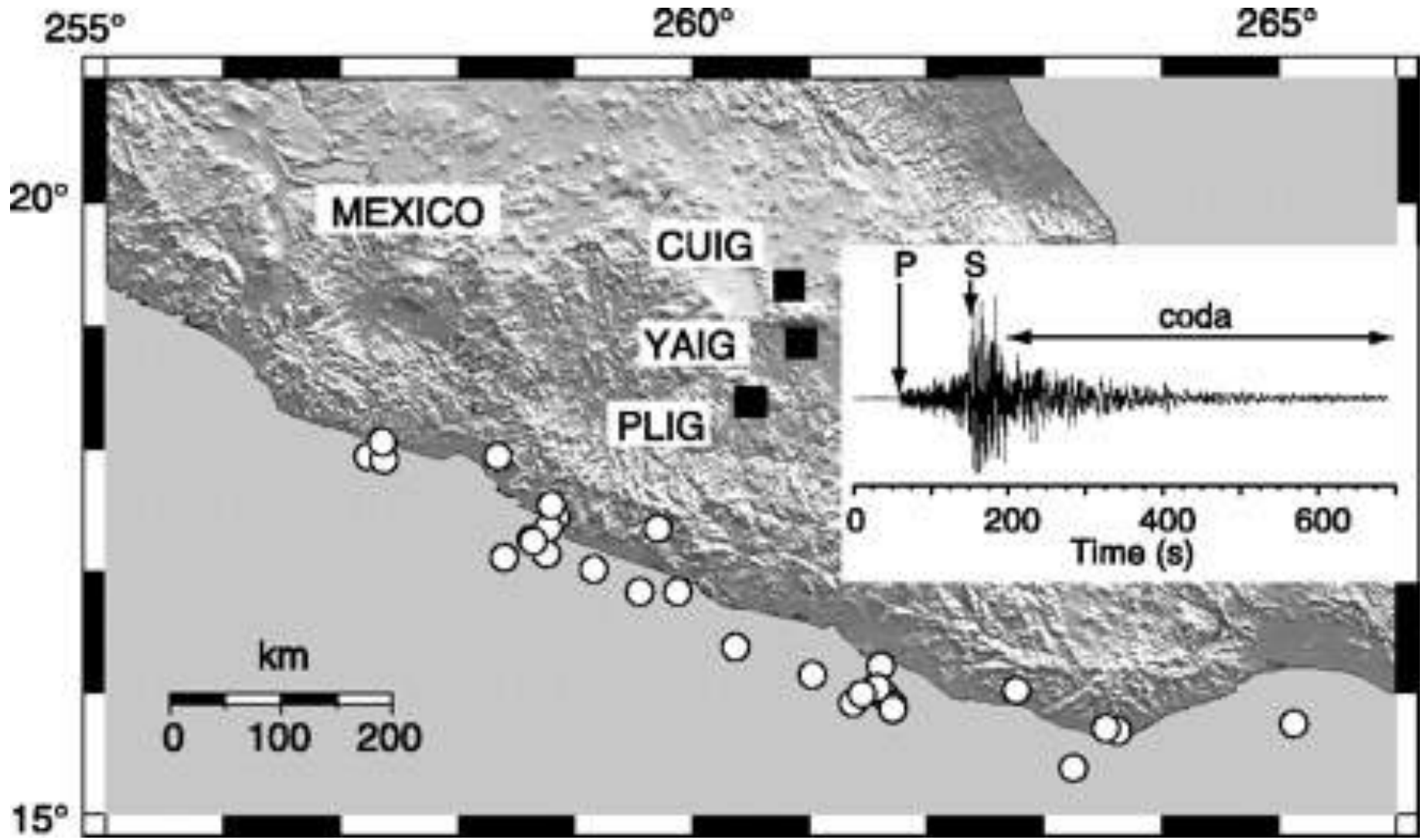
UD

EW

NS



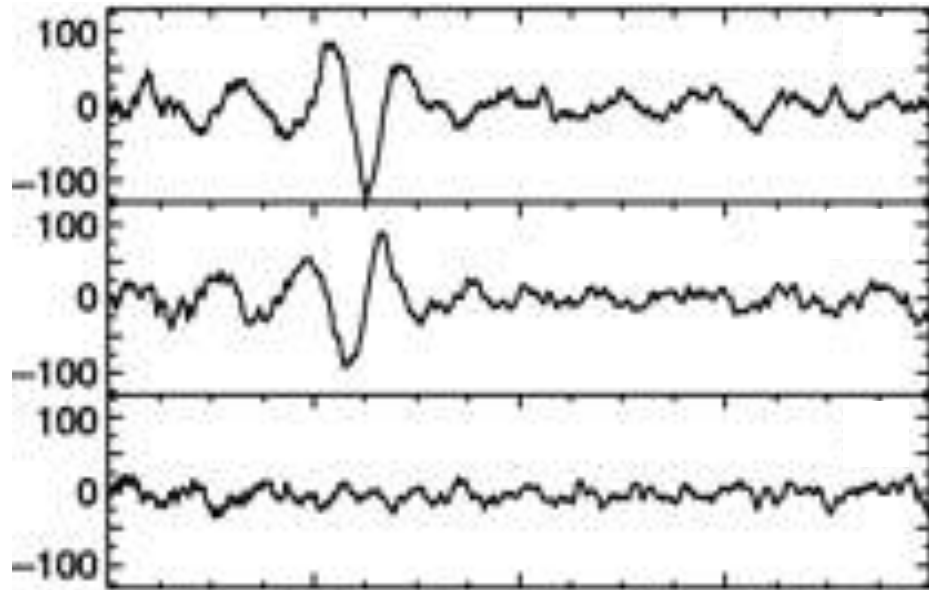
# Surface waves



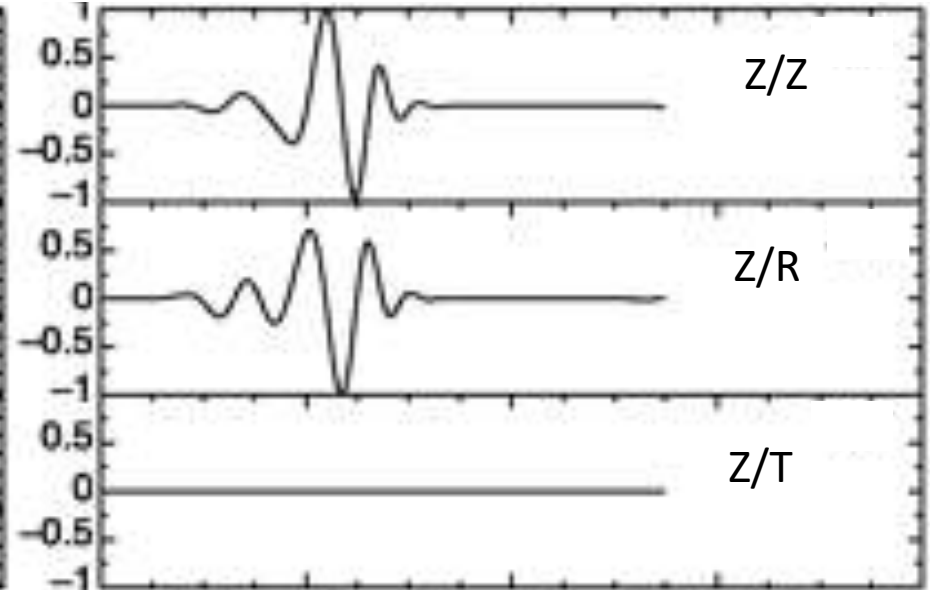
(Campillo and Paul, *Science*, **299**, 547-549, 2003)



# correlation

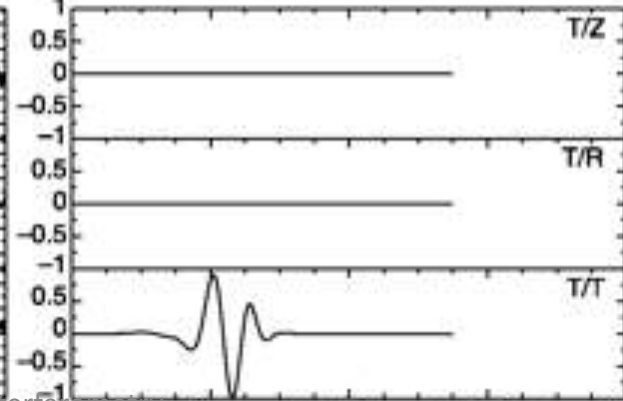
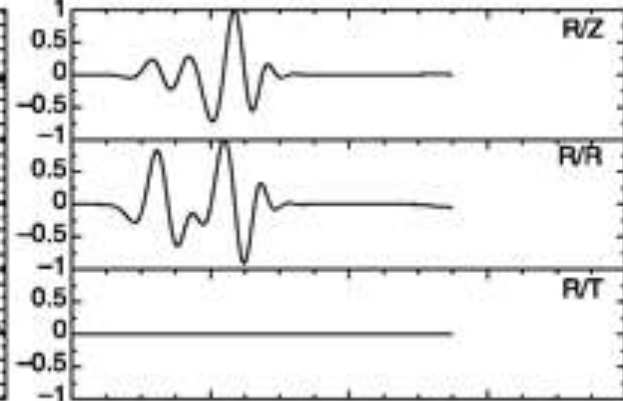
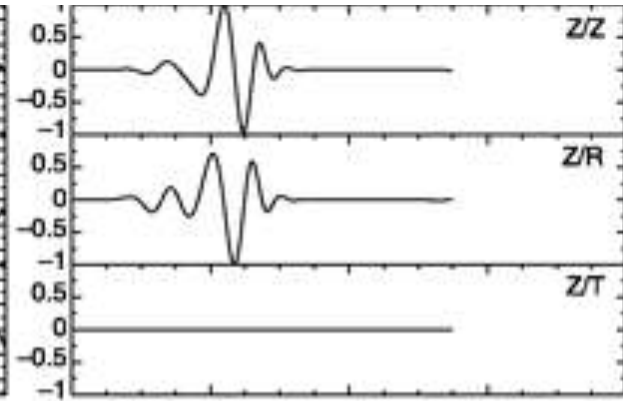
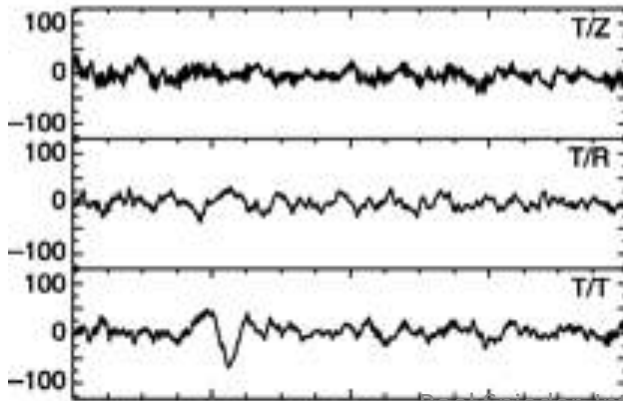
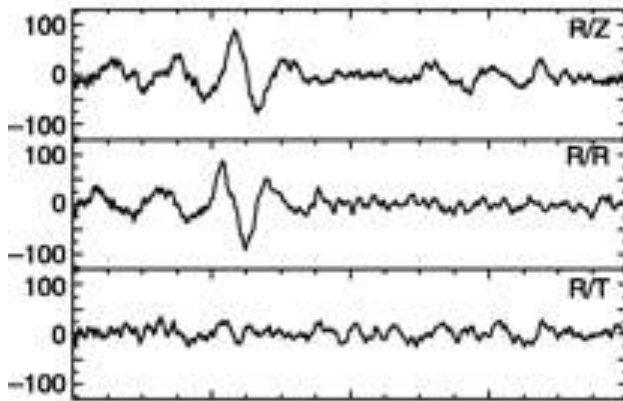
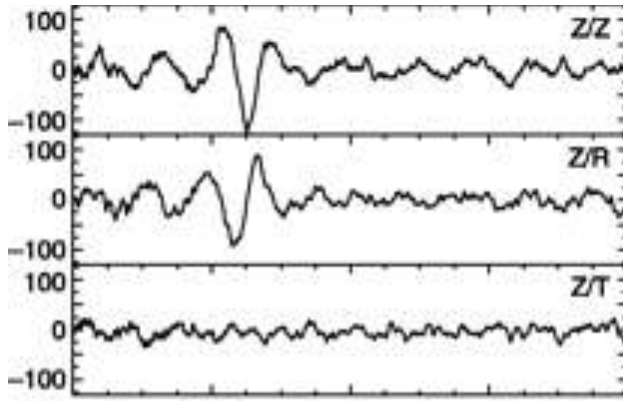


# Green's tensor



correlation

Green's tensor



Z/Z

Z/R

Z/T

R/Z

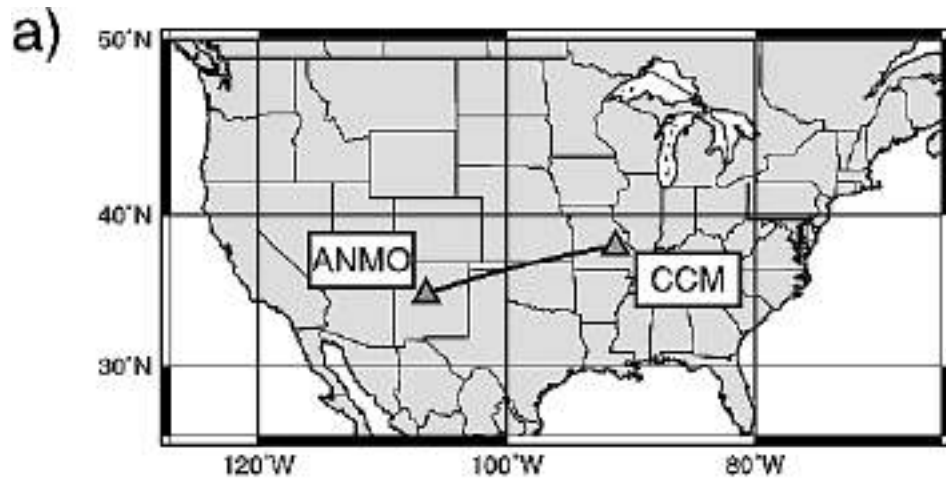
R/R

R/T

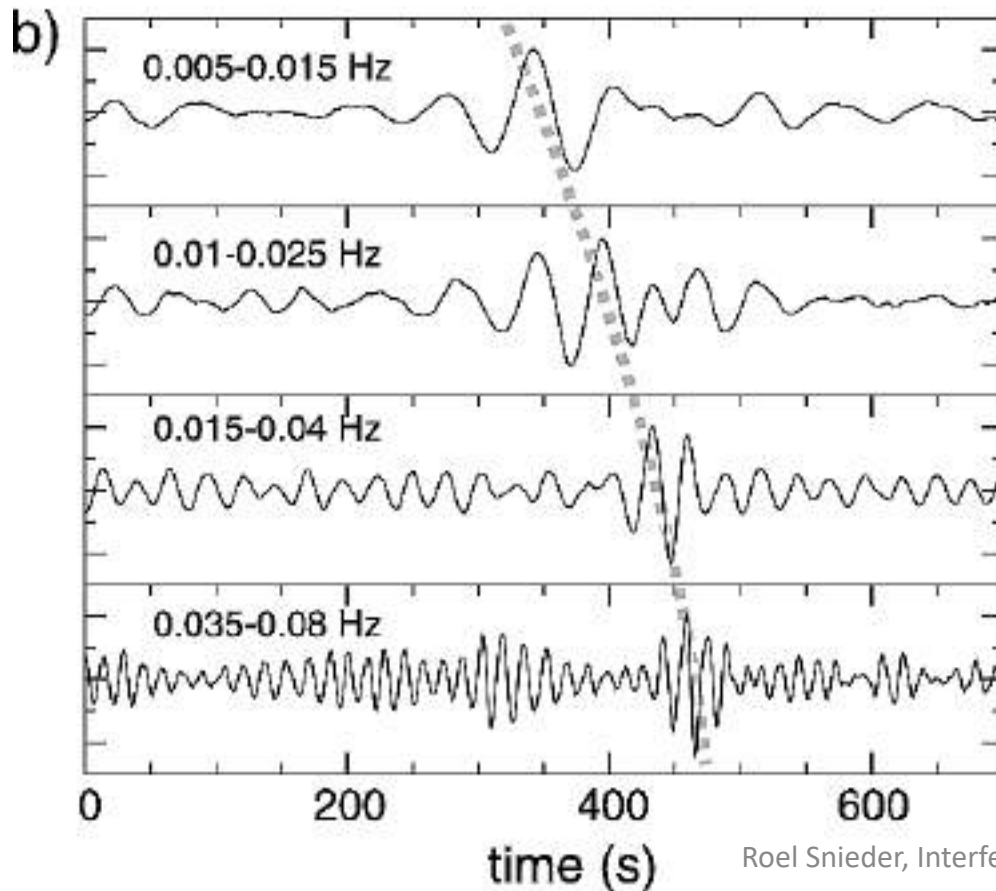
T/Z

T/R

T/T

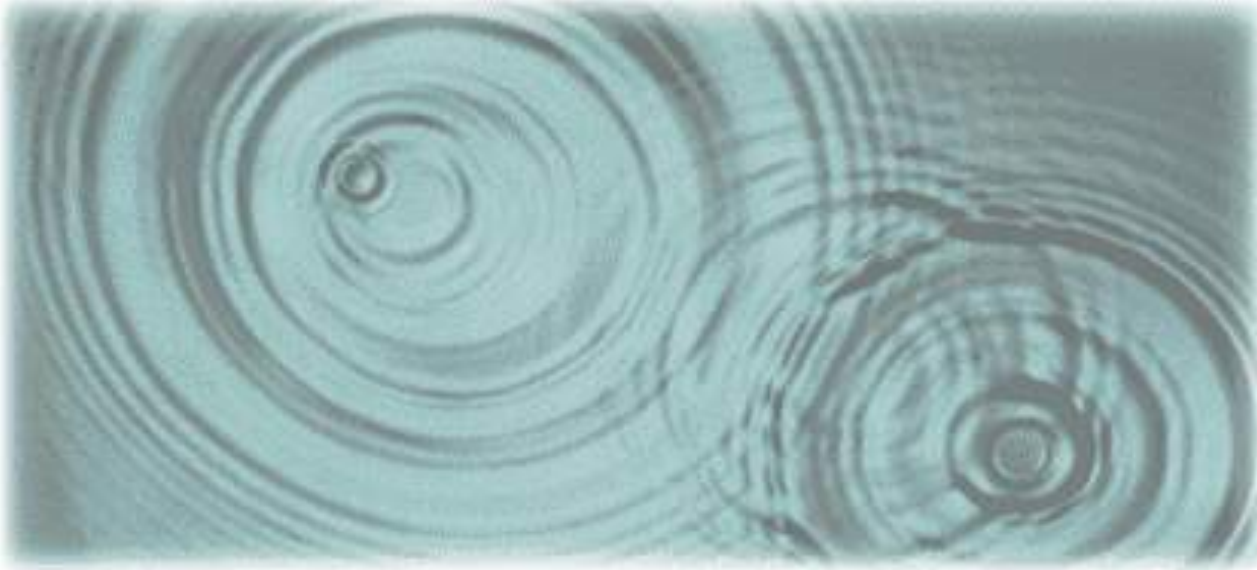


# Surface wave dispersion from noise

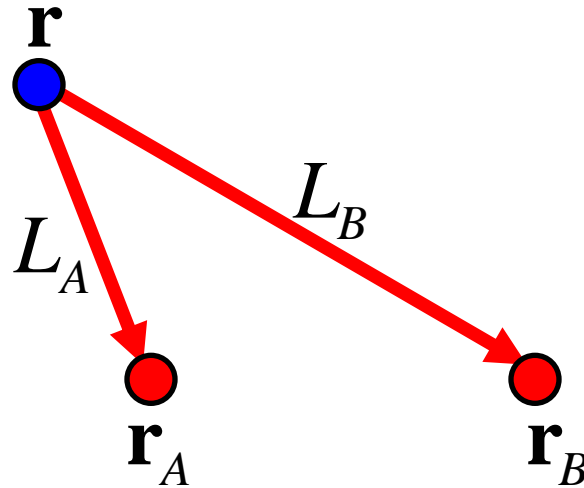


(Shapiro and Campillo,  
Geophys. Res. Lett.,  
**31**, L07614, 2004)

# Raindrop model

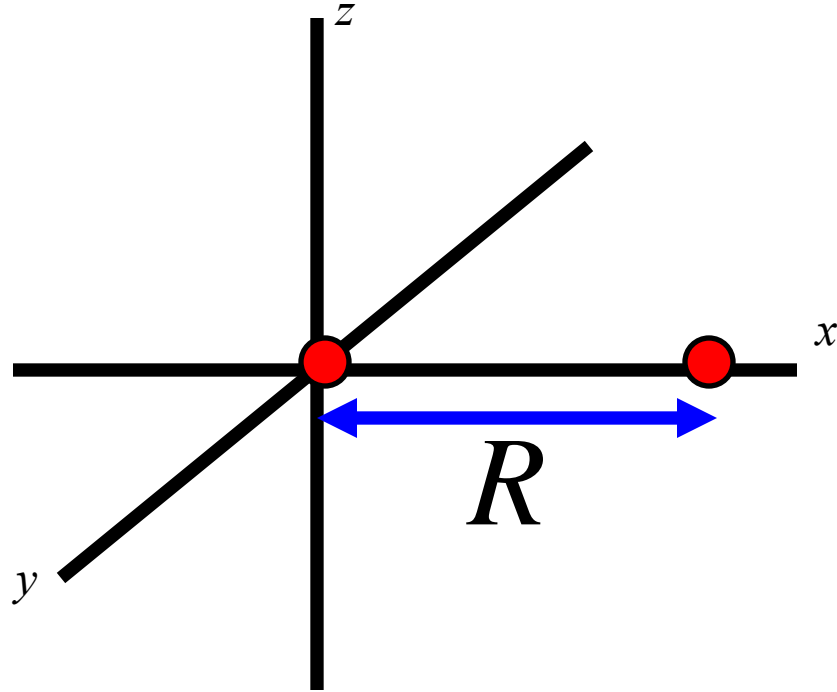


# Correlation as volume integral

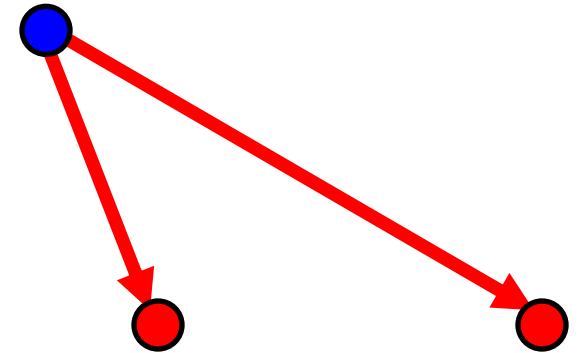


$$C_{AB}(\omega) = n|S(\omega)|^2 \int \frac{\exp(ik(L_A - L_B))}{L_A L_B} dV$$

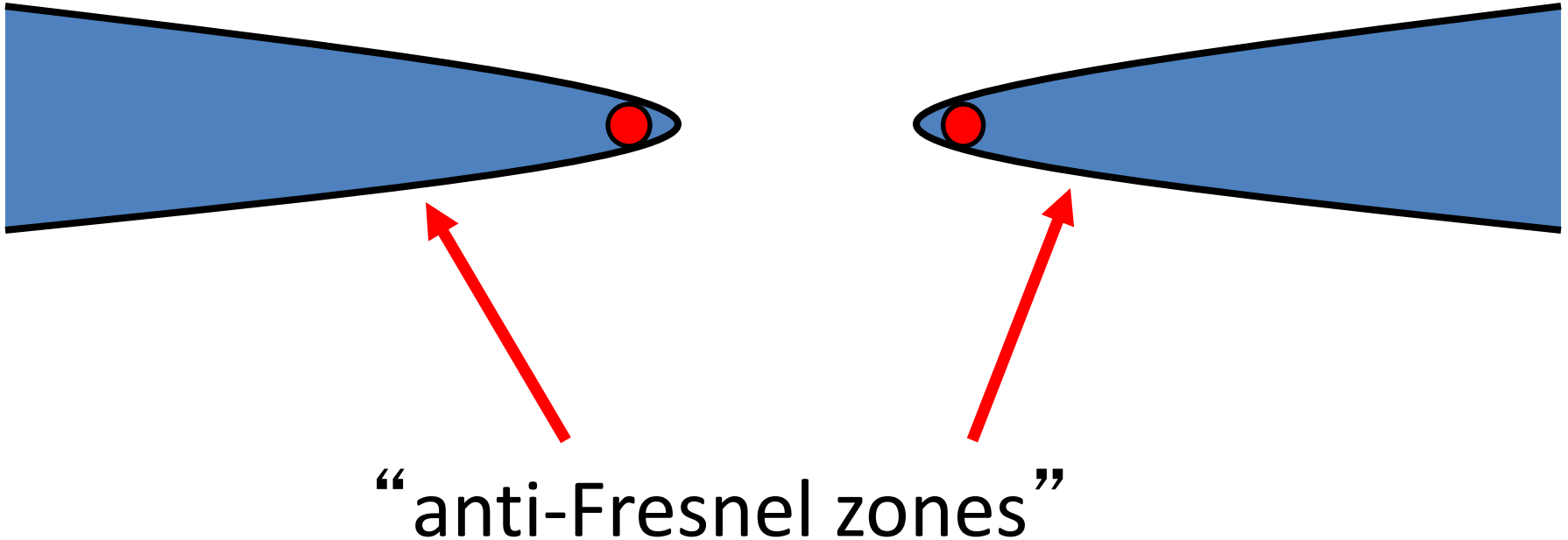
# Stationary phase contribution



$$y = z = 0$$



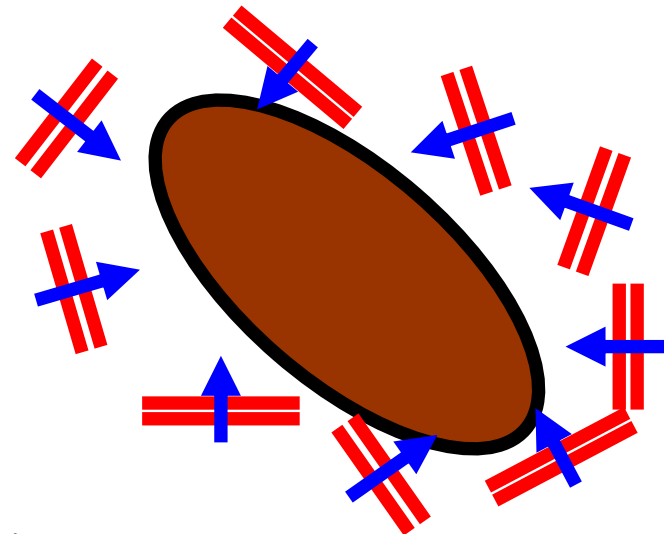
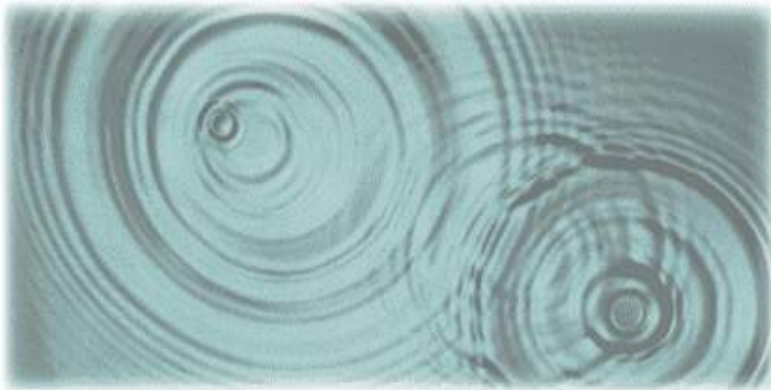
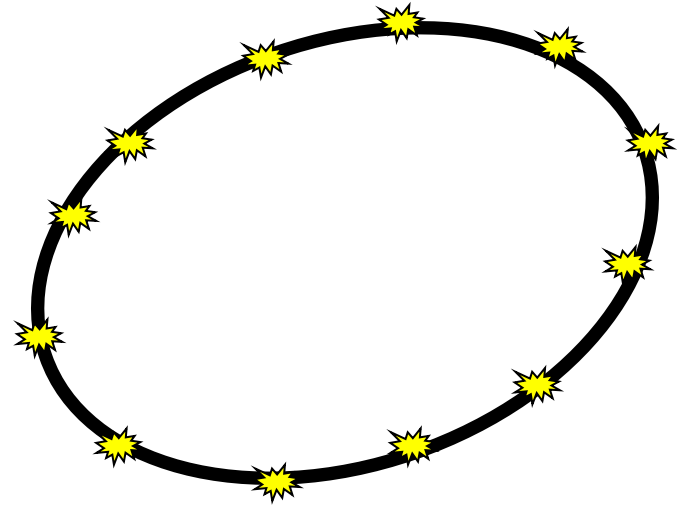
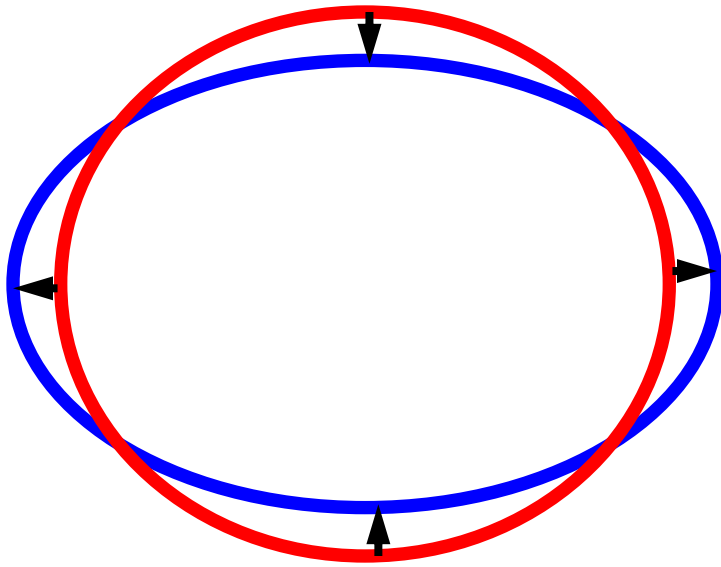
# Stationary phase regions



(Snieder, Phys. Rev. E, **69**, 046610, 2004,  
for reflected waves see:

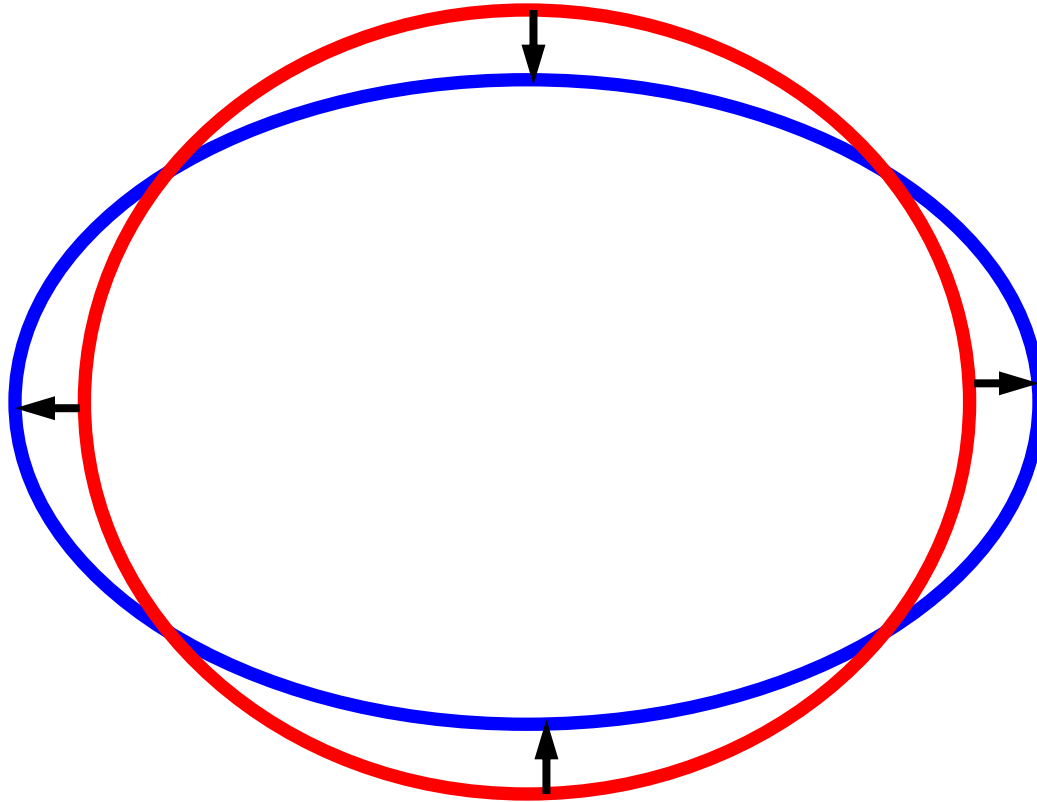
Snieder, Wapenaar, and Larner, Geophysics, 71, SI111-SI124, 2006)

# Four derivations





# Derivation based on normal-modes



(Lobkis and Weaver, JASA, **110**, 3011-3017, 2001)

# Displacement response

$$u(\mathbf{r}, t) = \sum_m \frac{u_m(\mathbf{r})}{\omega_m} (a_m \sin \omega_m t + b_m \cos \omega_m t)$$

$$G(\mathbf{r}, \mathbf{r}', t) = \sum_m \frac{u_m(\mathbf{r})u_m(\mathbf{r}')}{\omega_m} \sin \omega_m t \times H(t)$$



Heaviside function

# Velocity response

$$v(\mathbf{r}, t) = \sum_m u_m(\mathbf{r}) (a_m \cos \omega_m t - b_m \sin \omega_m t)$$

$$G^{(v)}(\mathbf{r}, \mathbf{r}', t) = \sum_m u_m(\mathbf{r}) u_m(\mathbf{r}') \cos \omega_m t \times H(t)$$

# Uncorrelated excitation

$$v(\mathbf{r}, t) = \sum_m u_m(\mathbf{r}) (a_m \cos \omega_m t - b_m \sin \omega_m t)$$

$$\langle a_n a_m \rangle = F^2 \delta_{nm}$$

$$\langle b_n b_m \rangle = F^2 \delta_{nm}$$

$$\langle a_n b_m \rangle = 0$$

# Correlation

$$C_{AB}^{(v)}(\tau) = \langle v(\mathbf{r}_A, t)v(\mathbf{r}_B, t + \tau) \rangle$$

$$v(\mathbf{r}_{A,B}, t) = \sum_m u_m(\mathbf{r}_{A,B}) (a_m \cos \omega_m t - b_m \sin \omega_m t)$$

# Correlation as a sum of modes

$$\begin{aligned} C_{AB}^{(\nu)}(\tau) &= \sum_{n,m} u_n(\mathbf{r}_A) u_m(\mathbf{r}_B) \\ &\quad \times \left\{ \langle a_n a_m \rangle \cos \omega_n t \cos \omega_m (t + \tau) \right. \\ &\quad - \langle a_n b_m \rangle \cos \omega_n t \sin \omega_m (t + \tau) \\ &\quad - \langle b_n a_m \rangle \sin \omega_n t \cos \omega_m (t + \tau) \\ &\quad \left. + \langle b_n b_m \rangle \sin \omega_n t \sin \omega_m (t + \tau) \right\} \end{aligned}$$

# For uncorrelated modes

$$\begin{aligned}
 C_{AB}^{(\nu)}(\tau) &= \sum_{n,m} u_n(\mathbf{r}_A) u_m(\mathbf{r}_B) \\
 &\quad \times \left\{ \langle a_n a_m \rangle \cos \omega_n t \cos \omega_m (t + \tau) \right. \\
 &\quad - \langle \cancel{a_n b_m} \rangle \cos \omega_n t \sin \omega_m (t + \tau) \\
 &\quad - \langle \cancel{b_n a_m} \rangle \sin \omega_n t \cos \omega_m (t + \tau) \\
 &\quad \left. + \langle b_n b_m \rangle \sin \omega_n t \sin \omega_m (t + \tau) \right\}
 \end{aligned}$$

$F^2 \delta_{nm}$

# For uncorrelated modes

$$C_{AB}^{(\nu)}(\tau) = \sum_{n,m} u_n(\mathbf{r}_A) u_m(\mathbf{r}_B) \times \left\{ \begin{aligned} & \langle a_n a_m \rangle \cos \omega_n t \cos \omega_m (t + \tau) \\ & - \langle a_n b_m \rangle \cos \omega_n t \sin \omega_m (t + \tau) \\ & - \langle b_n a_m \rangle \sin \omega_n t \cos \omega_m (t + \tau) \\ & + \langle b_n b_m \rangle \sin \omega_n t \sin \omega_m (t + \tau) \end{aligned} \right\}$$



$$= \cos \omega_m \tau$$



# Correlation

$$C_{AB}^{(v)}(\tau) = F^2 \sum_m u_m(\mathbf{r}_A) u_m(\mathbf{r}_B) \cos \omega_m \tau$$

## Green's function

$$G^{(v)}(\mathbf{r}_A, \mathbf{r}_B, \tau) = \sum_m u_m(\mathbf{r}_A) u_m(\mathbf{r}_B) \cos \omega_m \tau \times H(\tau)$$

# Correlation

$$C_{AB}^{(v)}(\tau) = F^2 \sum_m u_m(\mathbf{r}_A) u_m(\mathbf{r}_B) \cos \omega_m \tau$$

## Green's function

$$G^{(v)}(\mathbf{r}_A, \mathbf{r}_B, \tau) = \sum_m u_m(\mathbf{r}_A) u_m(\mathbf{r}_B) \cos \omega_m \tau \times H(\tau)$$

$$\tau > 0 \longrightarrow C_{AB}^{(v)}(\tau) = F^2 G^{(v)}(\mathbf{r}_A, \mathbf{r}_B, \tau)$$

# Correlation

$$C_{AB}^{(v)}(\tau) = F^2 \sum_m u_m(\mathbf{r}_A) u_m(\mathbf{r}_B) \cos \omega_m \tau$$

## Green's function

$$G^{(v)}(\mathbf{r}_A, \mathbf{r}_B, \tau) = \sum_m u_m(\mathbf{r}_A) u_m(\mathbf{r}_B) \cos \omega_m \tau \times H(\tau)$$

$$\tau < 0 \longrightarrow C_{AB}^{(v)}(\tau) = F^2 G^{(v)}(\mathbf{r}_A, \mathbf{r}_B, -\tau)$$

# Correlation and Green's function

$$C_{AB}^{(v)}(t) = F^2 \left\{ G^{(v)}(\mathbf{r}_A, \mathbf{r}_B, t) + G^{(v)}(\mathbf{r}_A, \mathbf{r}_B, -t) \right\}$$

- sum of causal and acausal Green's function
- holds for arbitrary heterogeneity

# Displacement instead of velocity

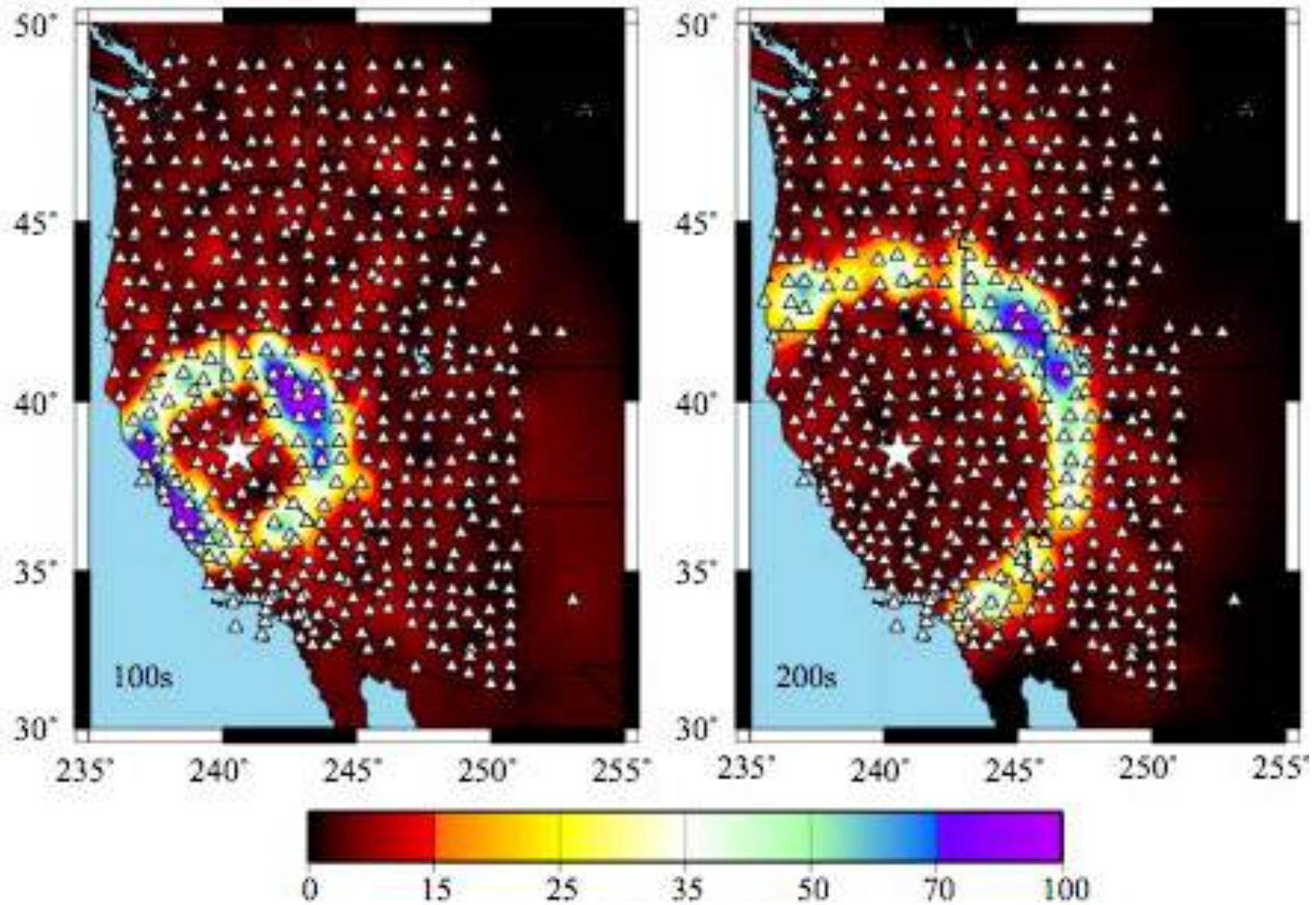
$$C_{AB}^{(v)} = \frac{d^2 C_{AB}^{(disp)}}{dt^2}$$

$$G^{(v)}(\mathbf{r}_A, \mathbf{r}_B, t) = \frac{\partial G^{(disp)}(\mathbf{r}_A, \mathbf{r}_B, t)}{\partial t}$$

$$\frac{dC_{AB}^{(disp)}}{dt}(t) = F^2 \left\{ G^{(disp)}(\mathbf{r}_A, \mathbf{r}_B, t) - G^{(disp)}(\mathbf{r}_A, \mathbf{r}_B, -t) \right\}$$

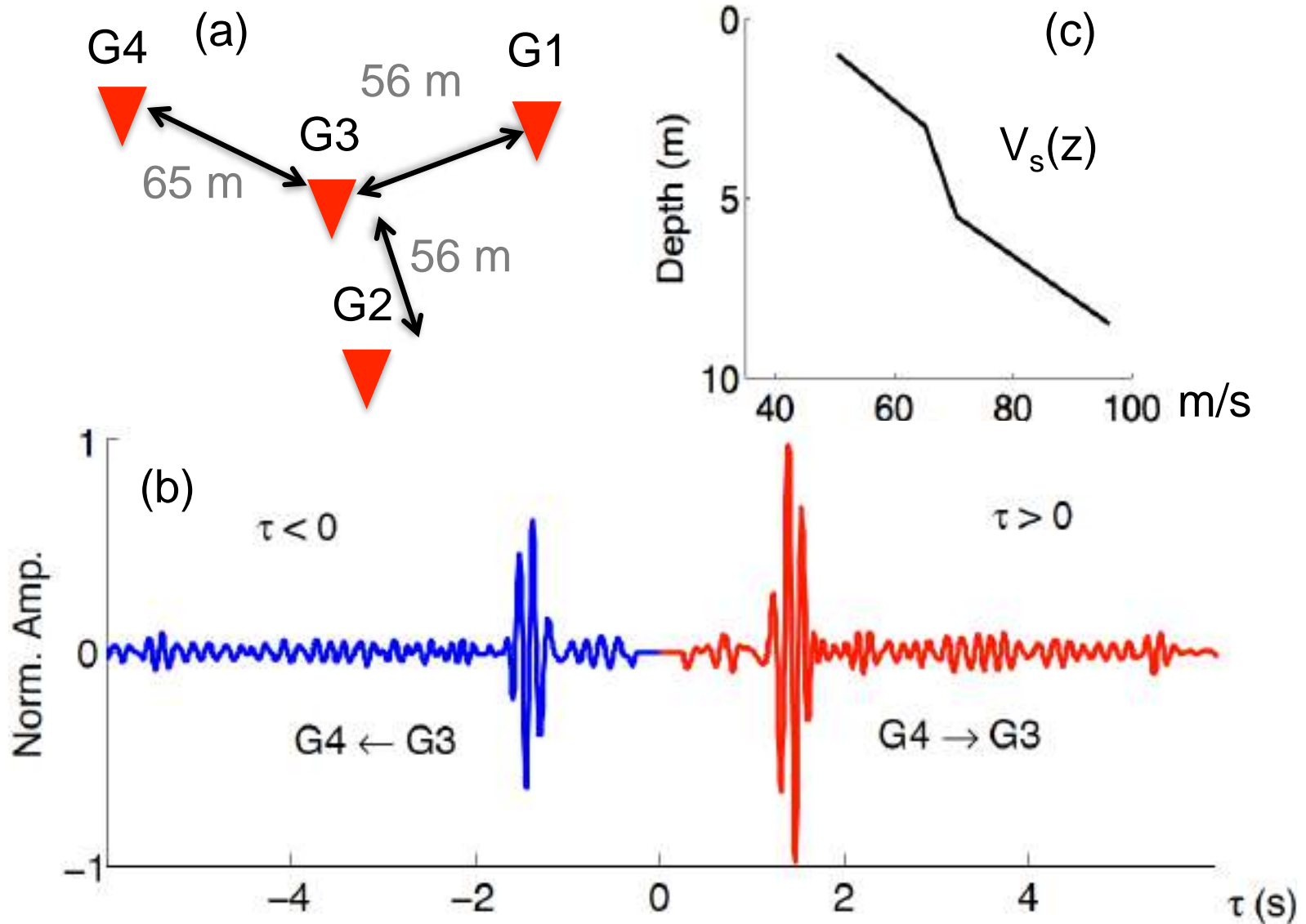
Conclusion: time derivative may appear

# Wave field reconstruction



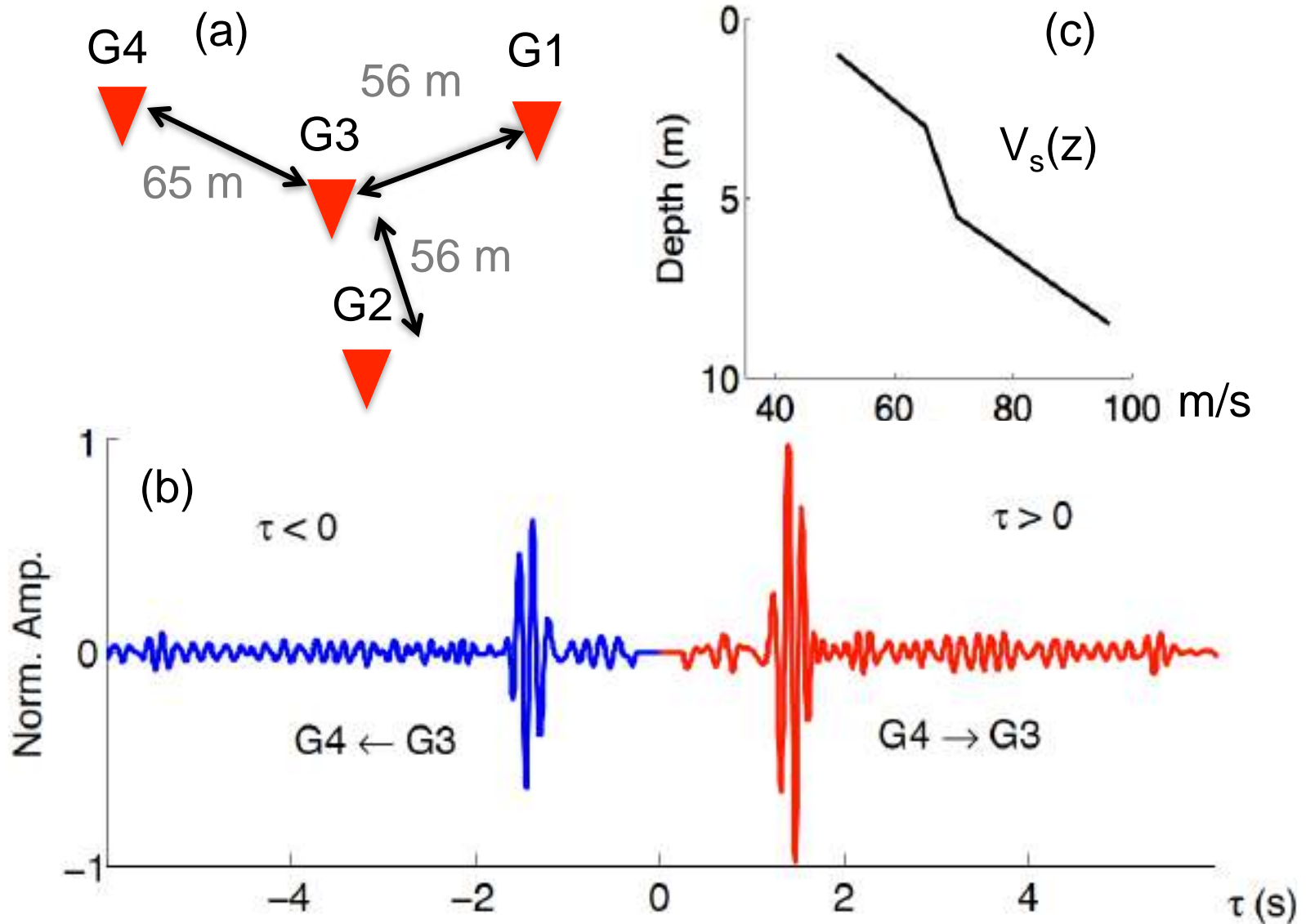
(Lin, Ritzwoller and Snieder, *Geophys. J. Int.*, 177, 1091–1110, 2009)

# Seismic interferometry



(Larose et al., Geophys. Res. Lett., 32, L16201, 2005)

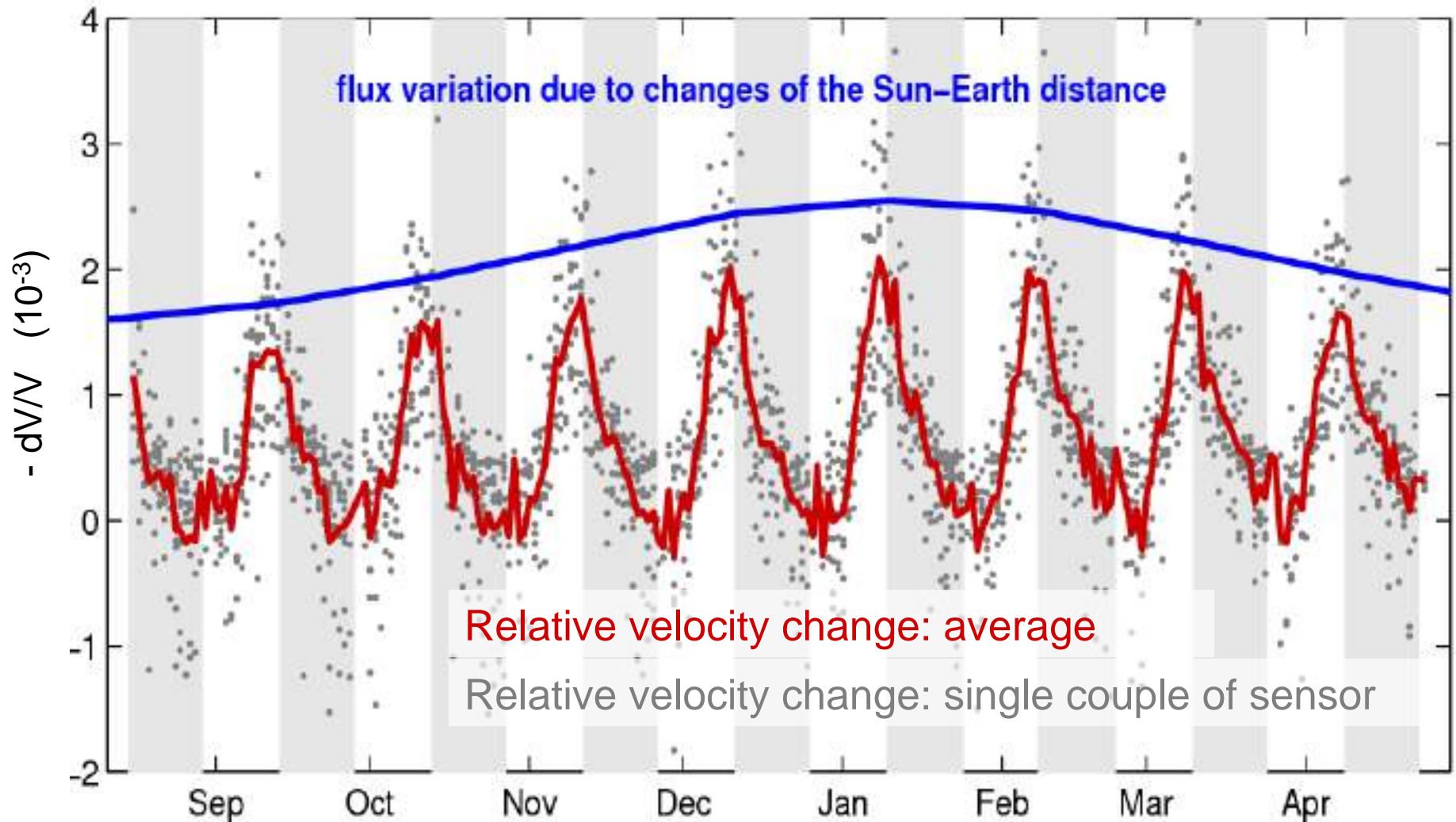
# Seismic interferometry **on the moon!**



(Larose et al., Geophys. Res. Lett., 32, L16201, 2005)



# Seismic interferometry on the moon



(Sens-Schönfelder & Larose, Phys. Rev. E, 78, 045601, 2008 )

# INSIGHTS

PERSPECTIVES



Downloaded from <https://science.sciencemag.org/> on May 17, 2019

## ENVIRONMENT

### *Why coal ash and tailings dam disasters occur*

Knowledge gaps and management shortcomings contribute to catastrophic dam failures

By **J. Carlos Santamarina<sup>1</sup>**,  
**Luis A. Torres-Cruz<sup>2</sup>**, **Robert C. Bachus<sup>3</sup>**

**O**n 25 January 2019, the structure damming a pond filled with iron ore mining wastes (tailings) burst at Brumadinho, Brazil (1), causing a massive mudslide that killed at least 232 people. This tailings dam failure was only the most recent in a long list of catastrophic tailings dam accidents (see the

first figure) (2, 3). Similar accidents also occur at electric power stations, where ponds are used to store coal combustion residuals such as fly and bottom ash. There are about 1000 operating ash ponds in the United States (4), and coal consumption patterns suggest that there may be more than 9000 worldwide. The catastrophic accident at the Kingston fossil power plant in Tennessee in 2008 (5) highlights the destructive potential of ash pond failures. Detailed analysis

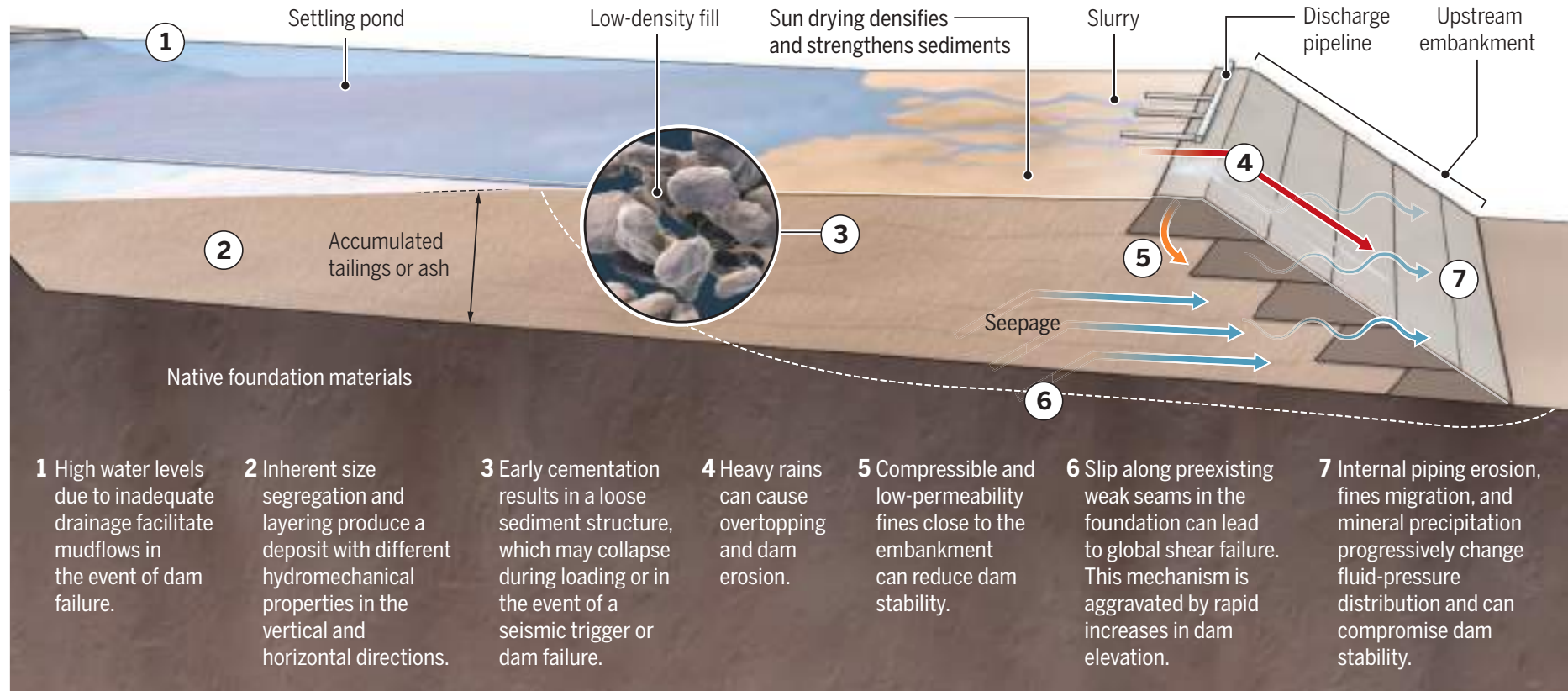
of tailings dam and ash pond failures shows that little-understood processes such as time-delayed triggering mechanisms are more likely to manifest when best engineering practices are disregarded.

Failure of the containment structure around mine tailings and coal ash is often followed by a fast-moving mudflow, which can run downstream for several miles, with catastrophic consequences. This liquefaction of the impounded materials may sug-

PHOTO: DOUGLAS MAGNO/AFR/GETTY IMAGES

# Mechanisms and processes of tailings dam and ash pond failures

Many different aspects of impounded ash and mine tailings and the associated dams can contribute to dam failures and the resulting fast-moving mudflows. The figure shows upstream construction as an example; most processes and mechanisms shown also apply to centerline and downstream construction methods.



**1** High water levels due to inadequate drainage facilitate mudflows in the event of dam failure.

**2** Inherent size segregation and layering produce a deposit with different hydromechanical properties in the vertical and horizontal directions.

**3** Early cementation results in a loose sediment structure, which may collapse during loading or in the event of a seismic trigger or dam failure.

**4** Heavy rains can cause overtopping and dam erosion.

**5** Compressible and low-permeability fines close to the embankment can reduce dam stability.

**6** Slip along preexisting weak seams in the foundation can lead to global shear failure. This mechanism is aggravated by rapid increases in dam elevation.

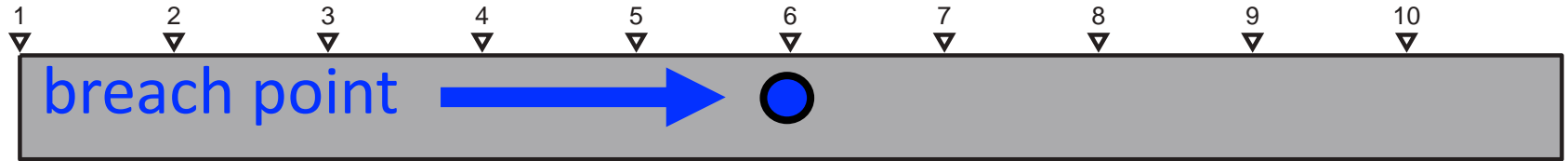
**7** Internal piping erosion, fines migration, and mineral precipitation progressively change fluid-pressure distribution and can compromise dam stability.

# Monitoring a dam

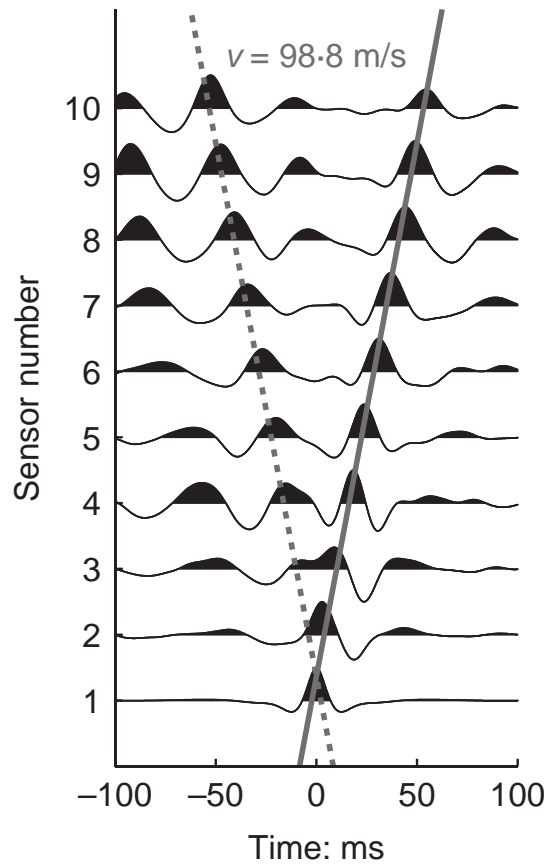
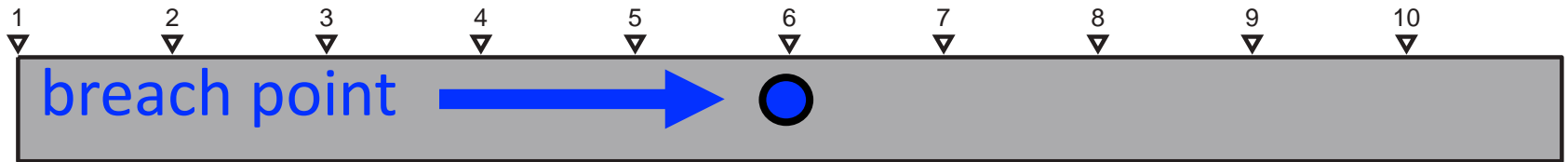
breach created by pulling a metal rod out of the dam



# Change in shear velocity in a dam

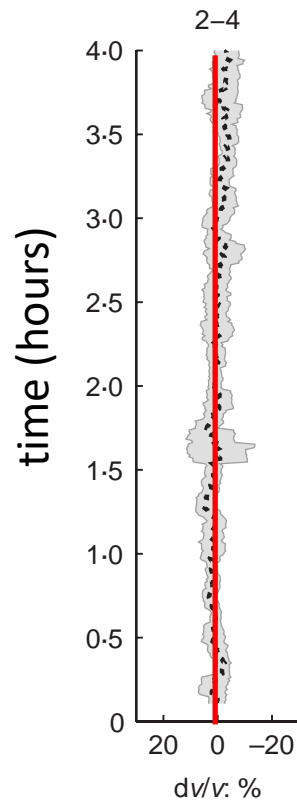
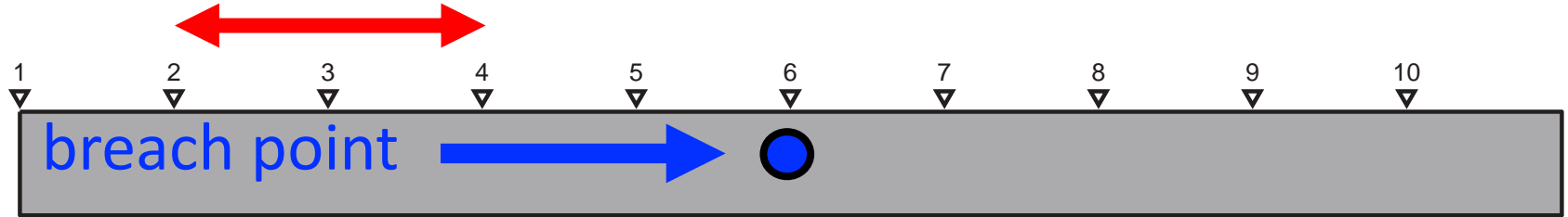


# Sensor 1 as virtual source

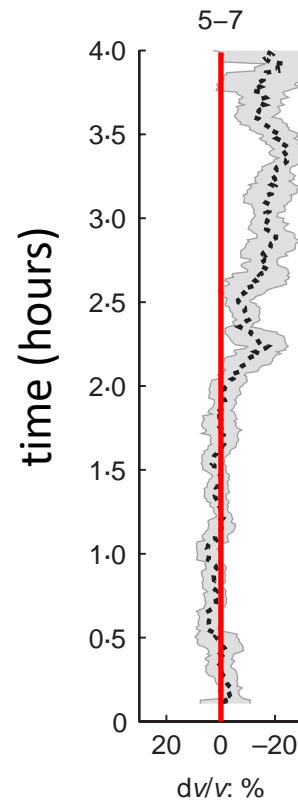
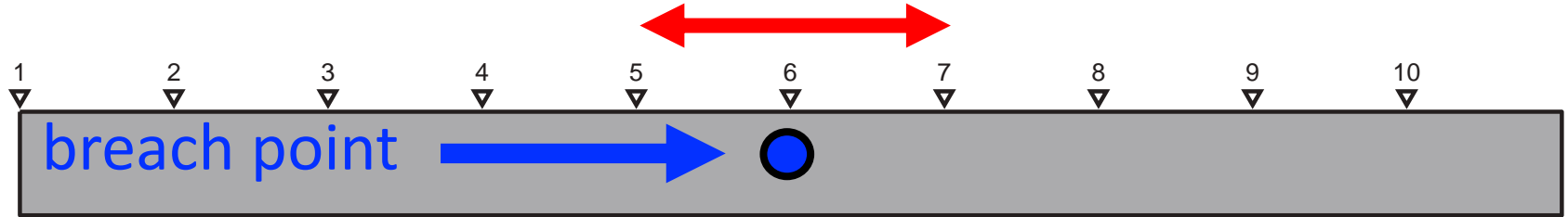


(Planes et al., *Geotechnique*, 66, 301-312, 2016)

# Change in shear velocity in a dam

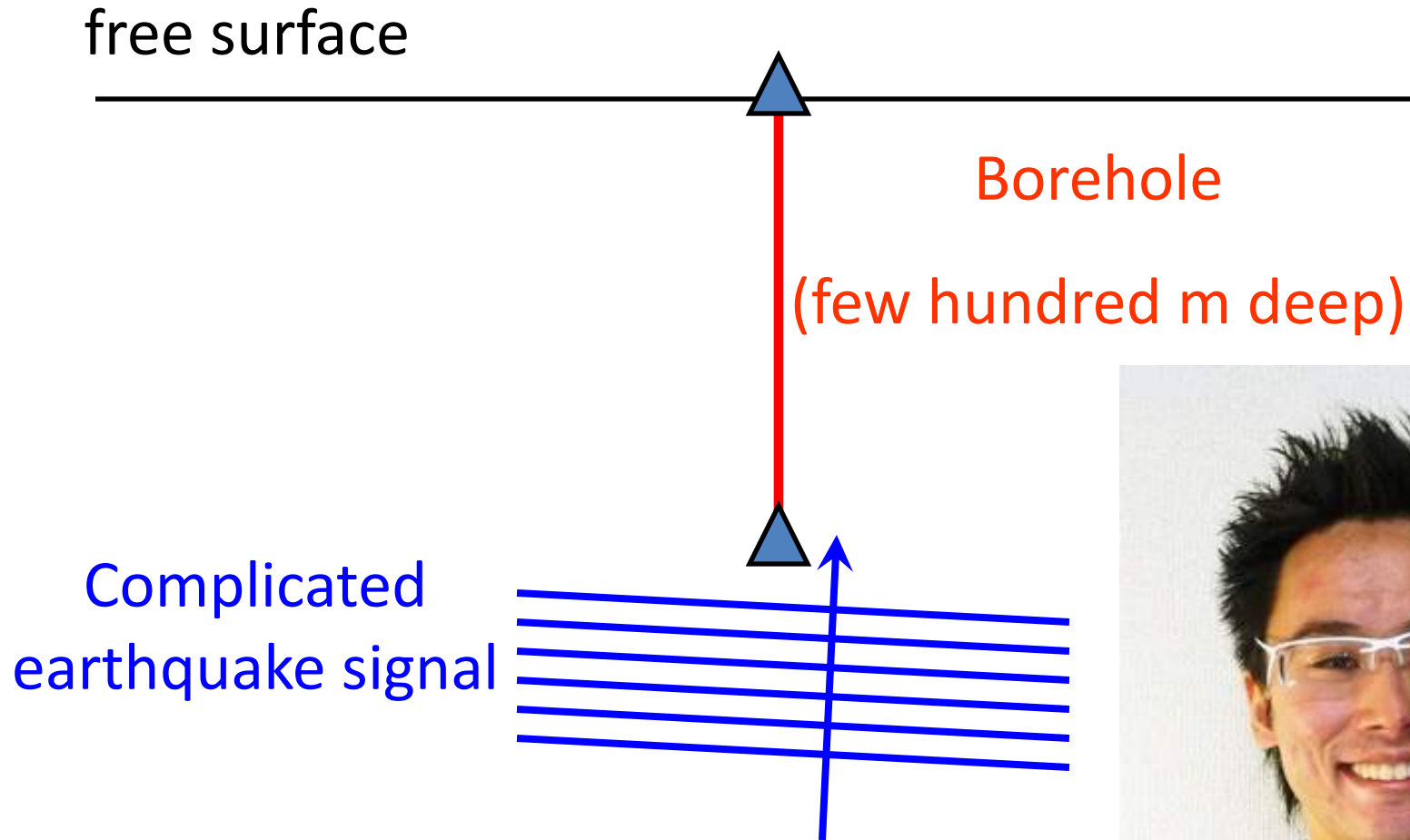


# Change in shear velocity in a dam

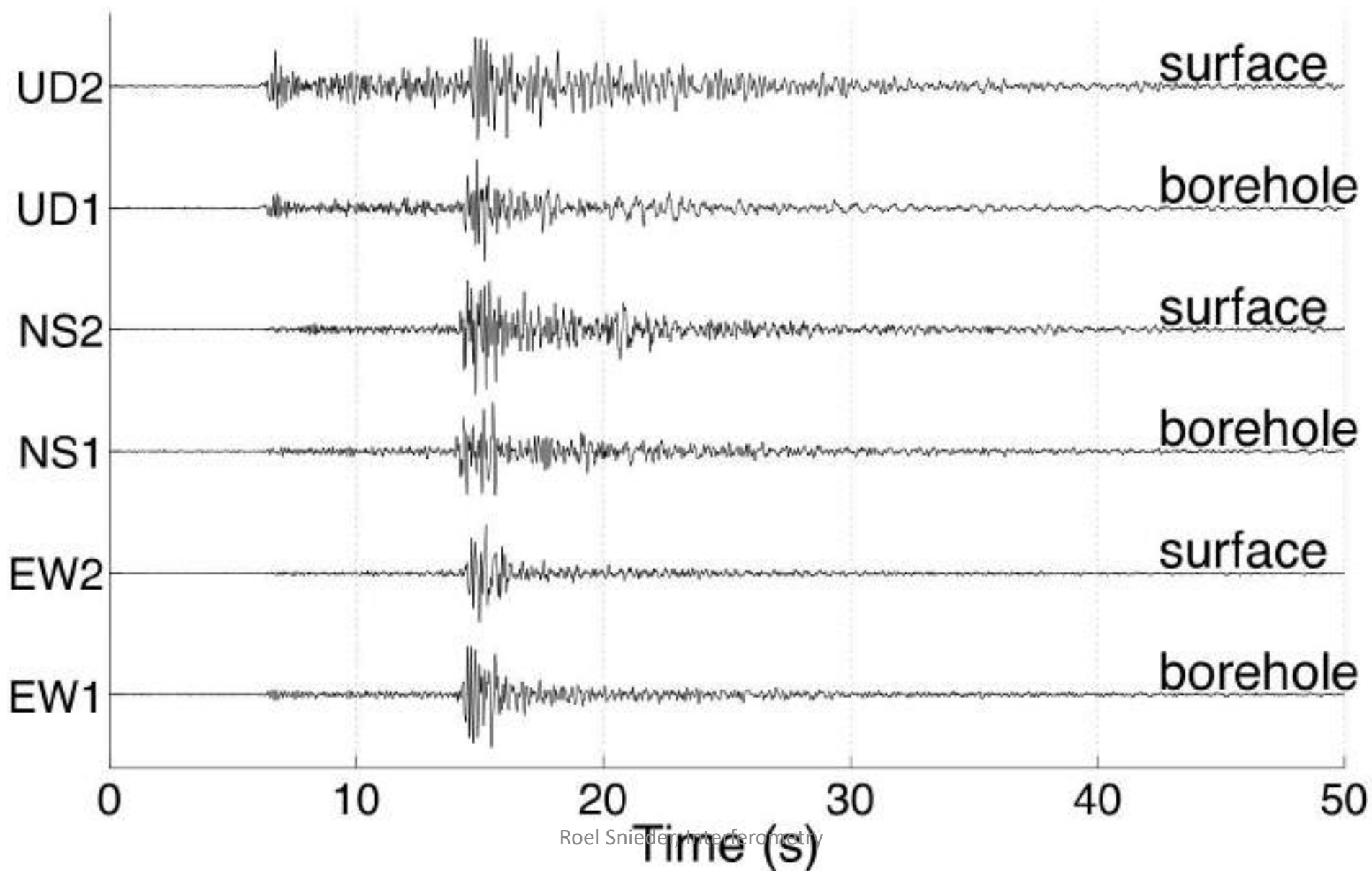




# Near-surface structure from Kik-Net

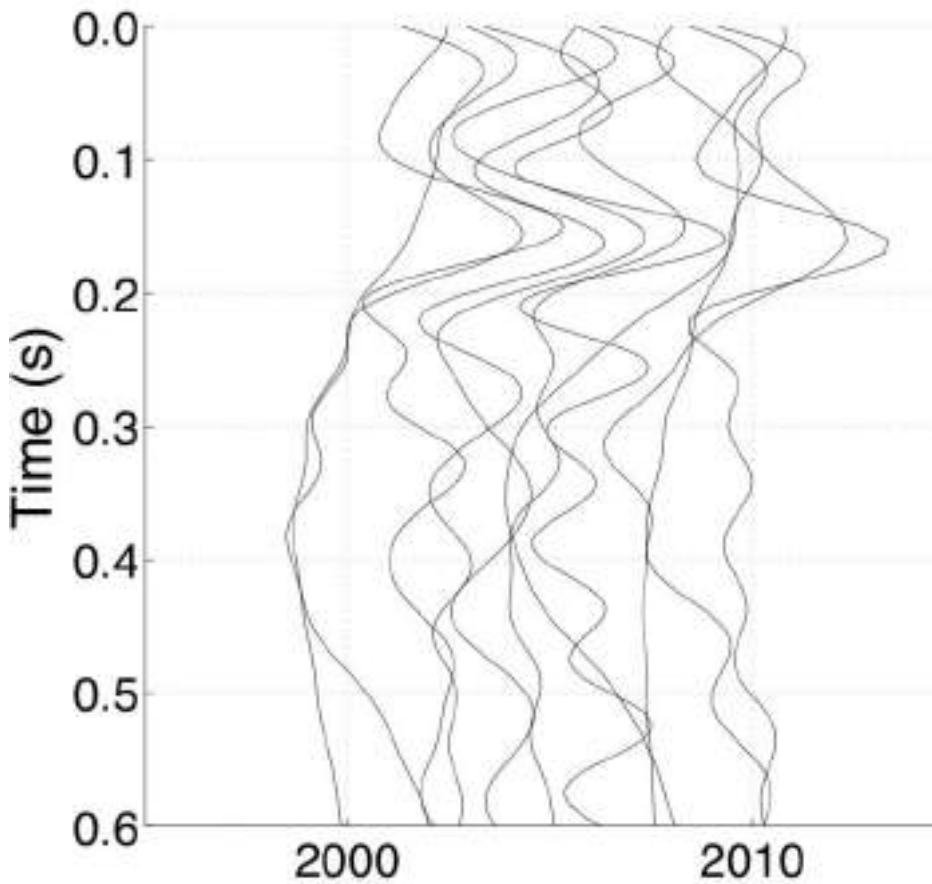


# Data at station NIGH13

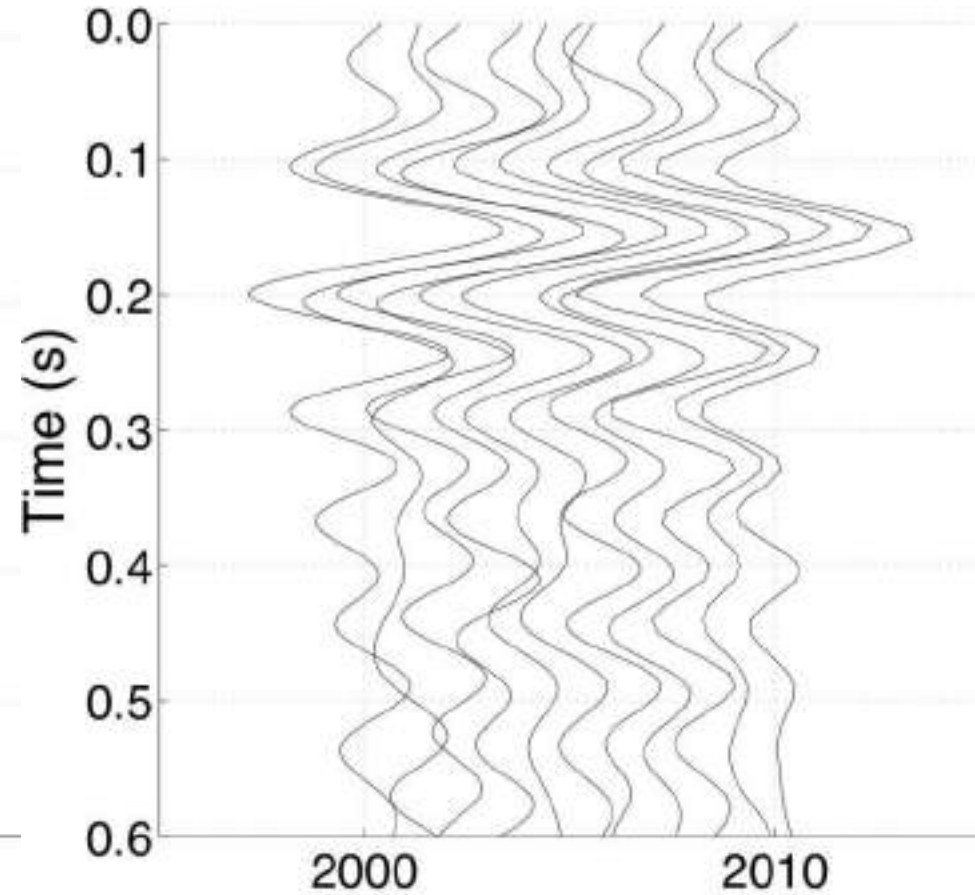


# Annual stacks at station NIGH13

correlation



deconvolution



# Correlation vs. deconvolution

$$u(z, \omega) = S(\omega)e^{-ikz}$$

$$\textit{correlation} = u(z = 0, \omega)u^*(z = D, \omega) = |S(\omega)|^2 e^{ikz}$$

$$\textit{deconvolution} = \frac{u(z = 0, \omega)}{u(z = D, \omega)} = e^{ikz}$$

# When there is one source

$$u_A = G_{AS}S$$

$$u_B = G_{BS}S$$

$$D = \frac{u_A}{u_B} = \frac{G_{AS}S}{G_{BS}S} = \frac{G_{SA}}{G_{SB}}$$

Independent of source signal

# When there are two sources

$$u_A = G_{AS_1} S_1 + G_{AS_2} S_2$$

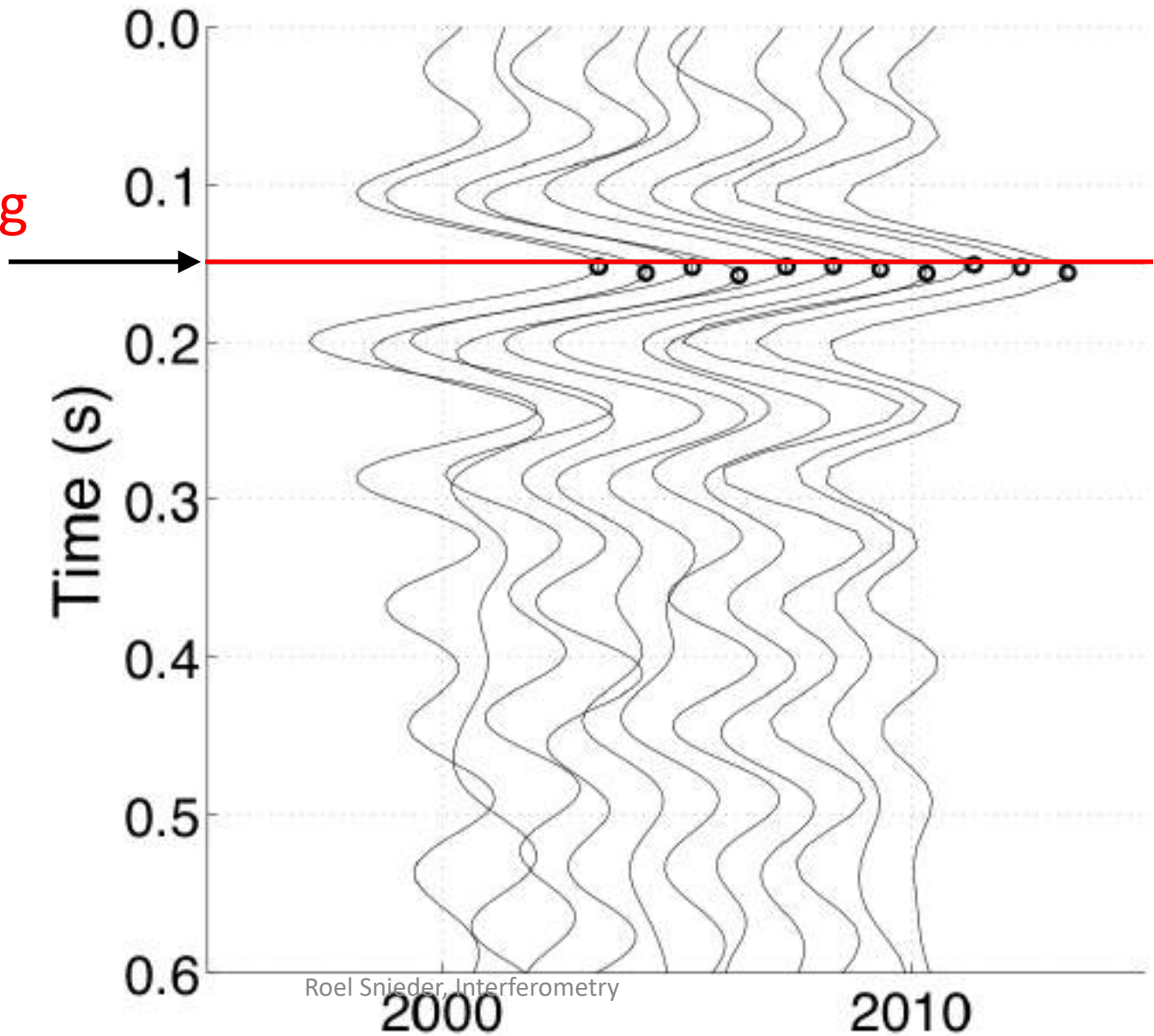
$$u_B = G_{BS_1} S_1 + G_{BS_2} S_2$$

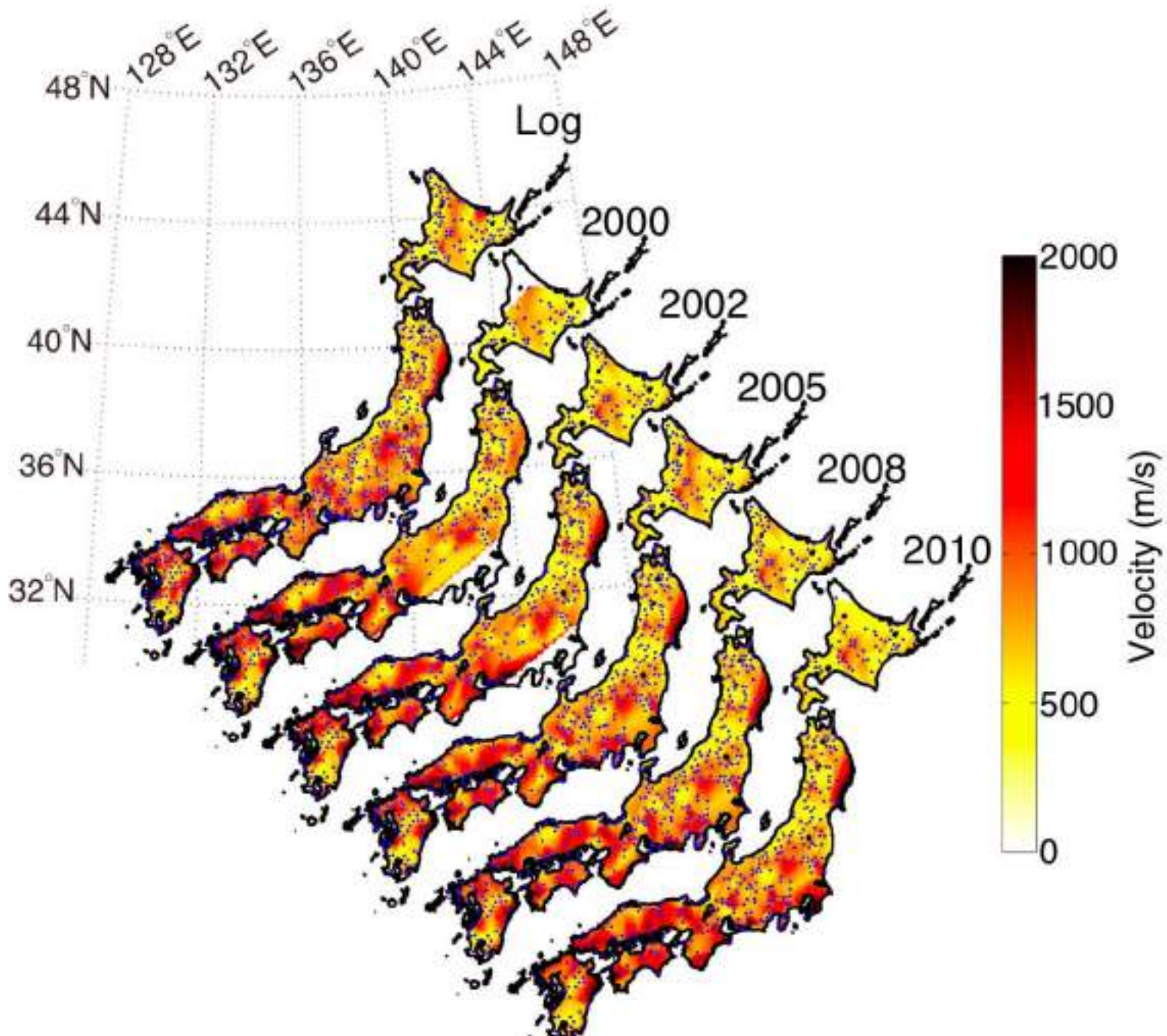
$$D = \frac{u_A}{u_B} = \frac{G_{AS_1} S_1 + G_{AS_2} S_2}{G_{BS_1} S_1 + G_{BS_2} S_2}$$

Source signal does not divide out

# Arrival time of shear wave

from logging  
data



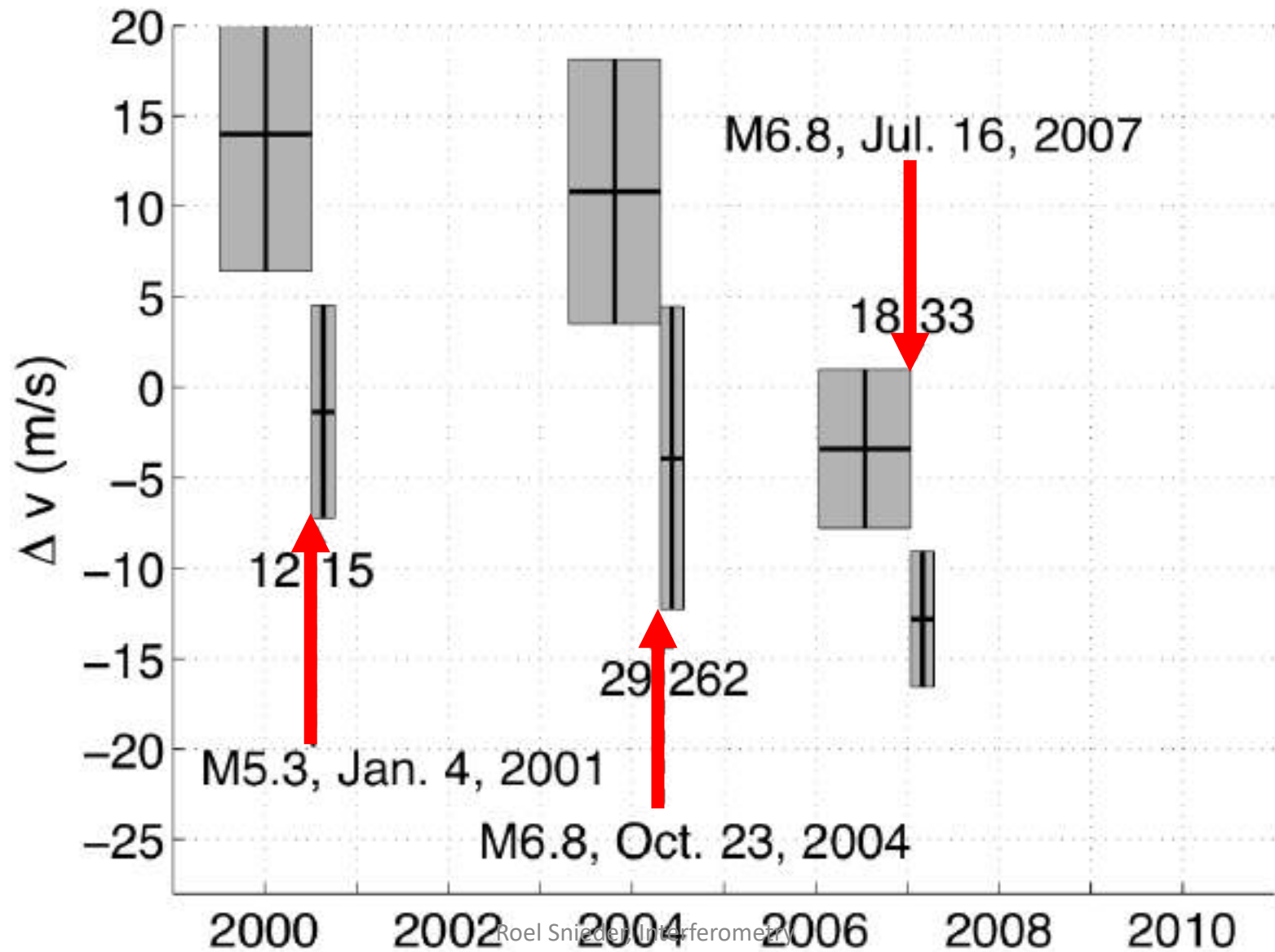


Roel Snieder, Interferometry

(Nakata and Snieder, JGR, 117, B01308, 2012)

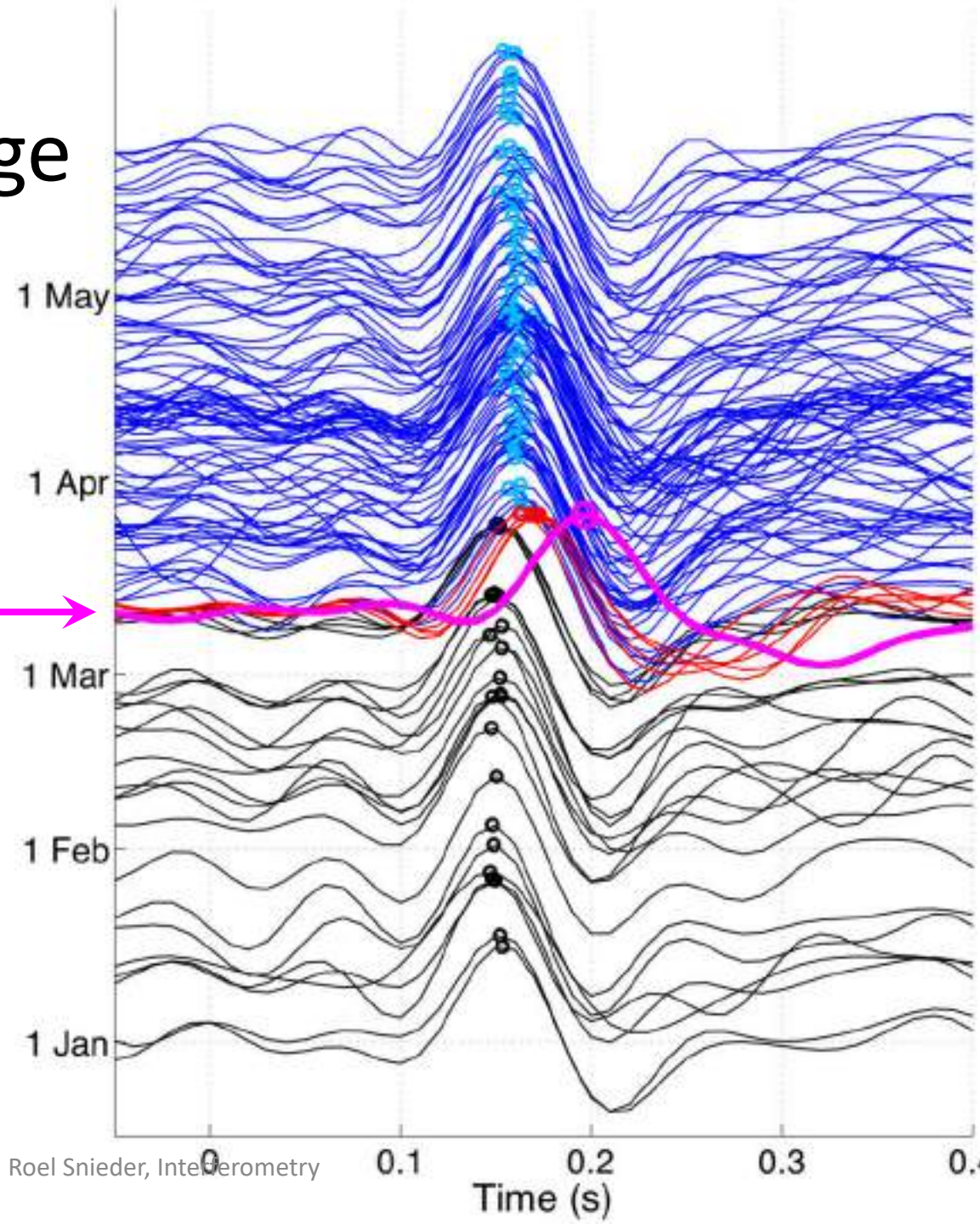


# S-waves in Niigata and earthquakes



# Time-lapse change

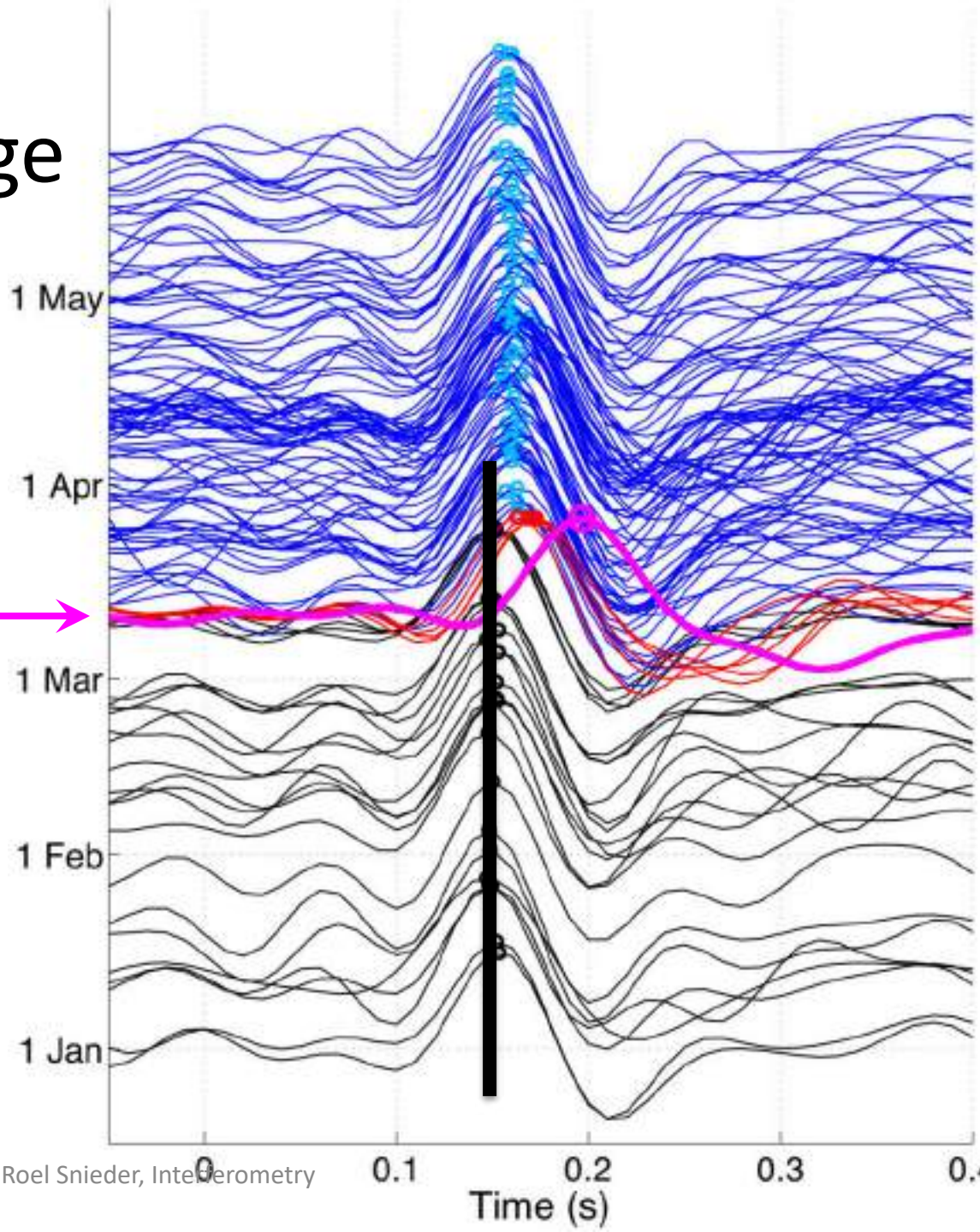
Tohoku-Oki  
earthquake



(Nakata and Snieder, Geophys.  
Res. Lett., 18, L17302, 2011)

# Time-lapse change

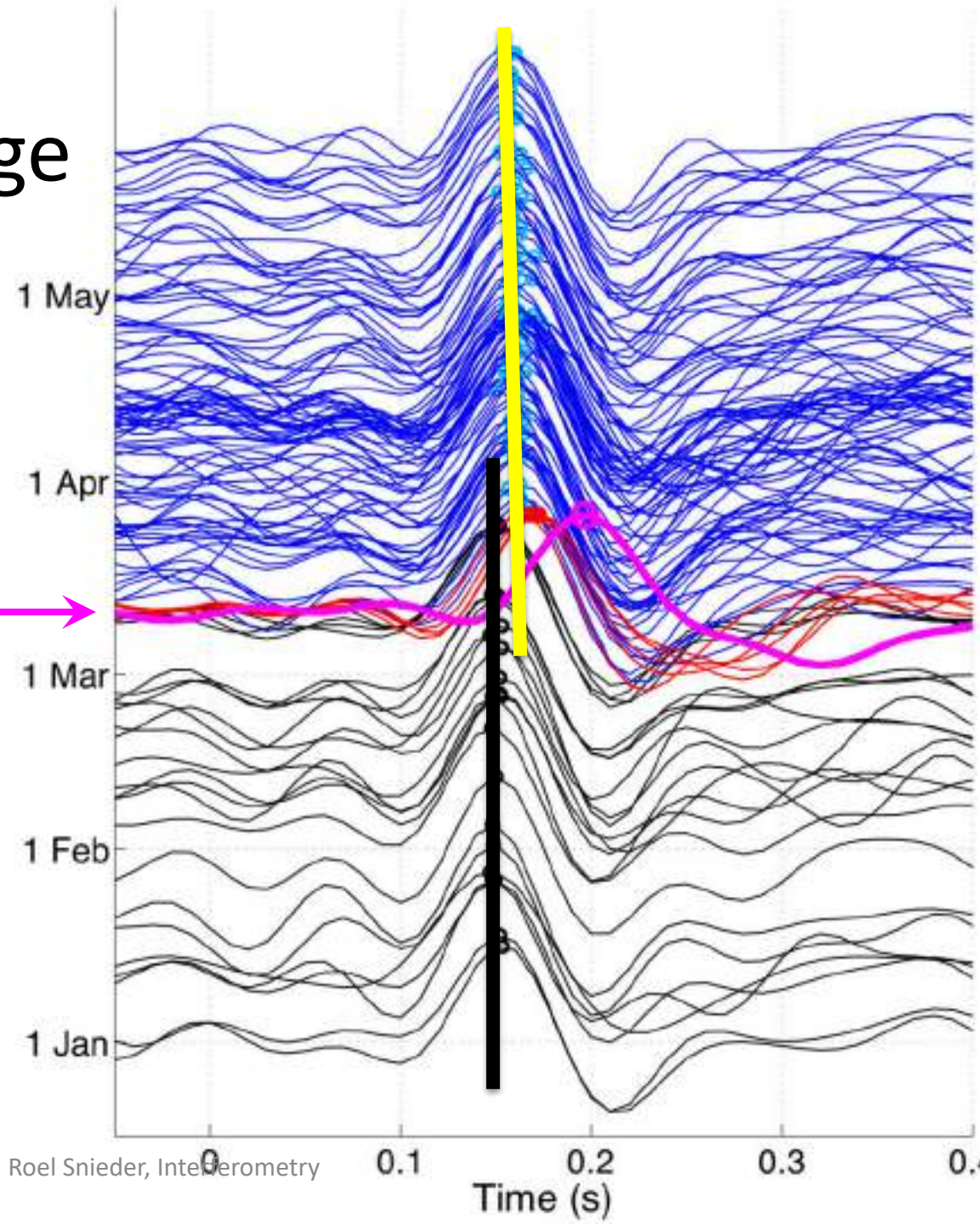
Tohoku-Oki  
earthquake



(Nakata and Snieder, Geophys.  
Res. Lett., 18, L17302, 2011)

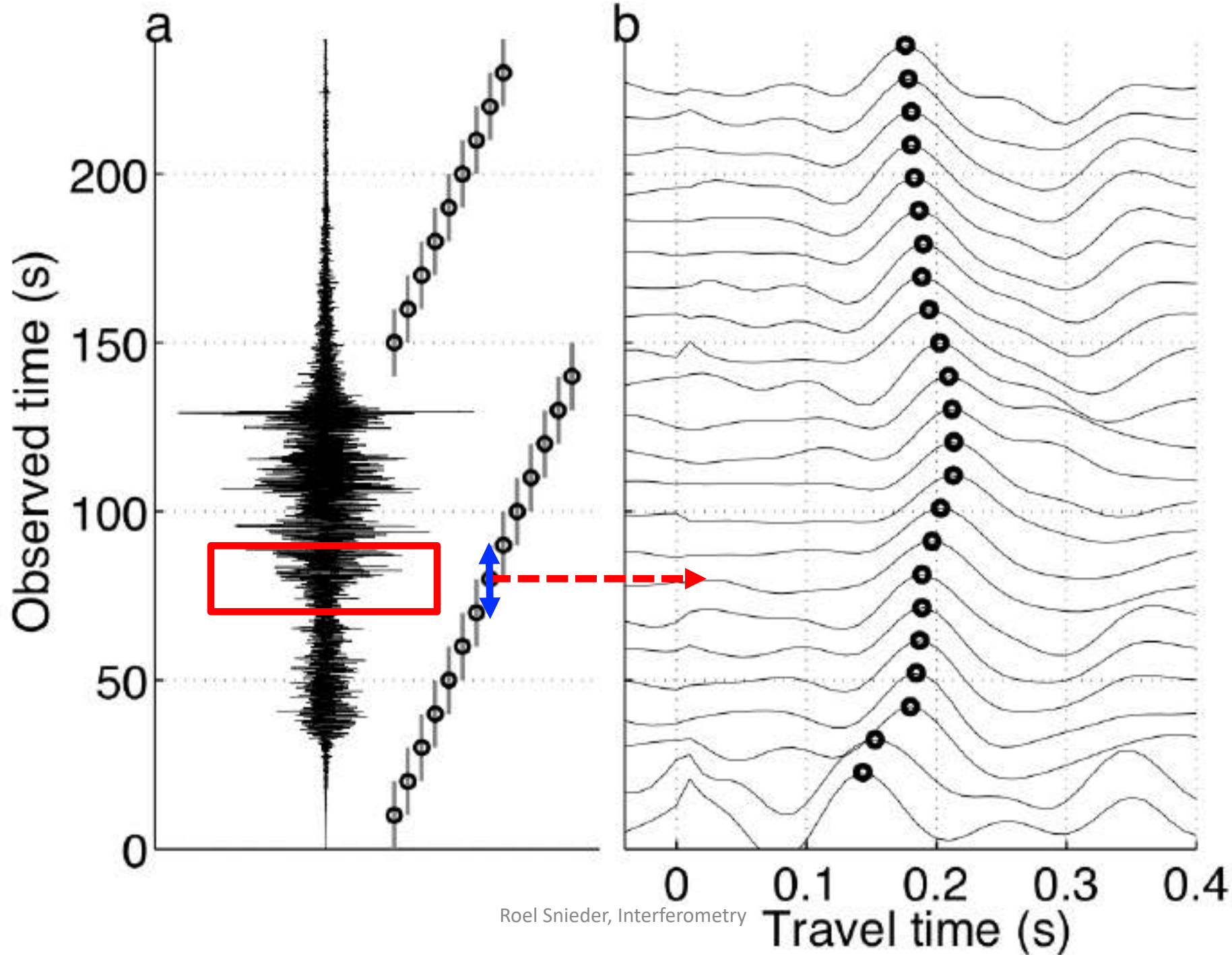
# Time-lapse change

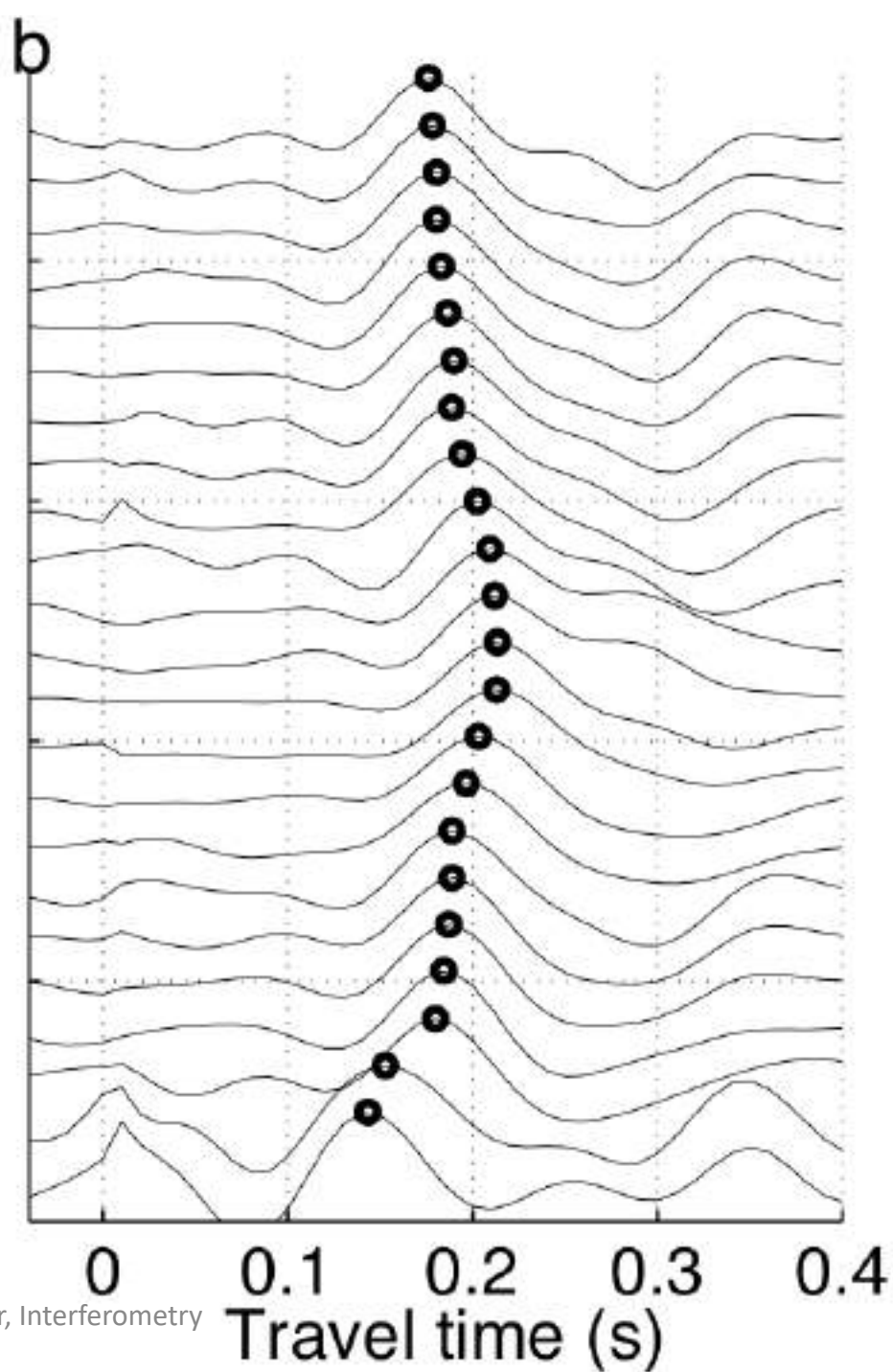
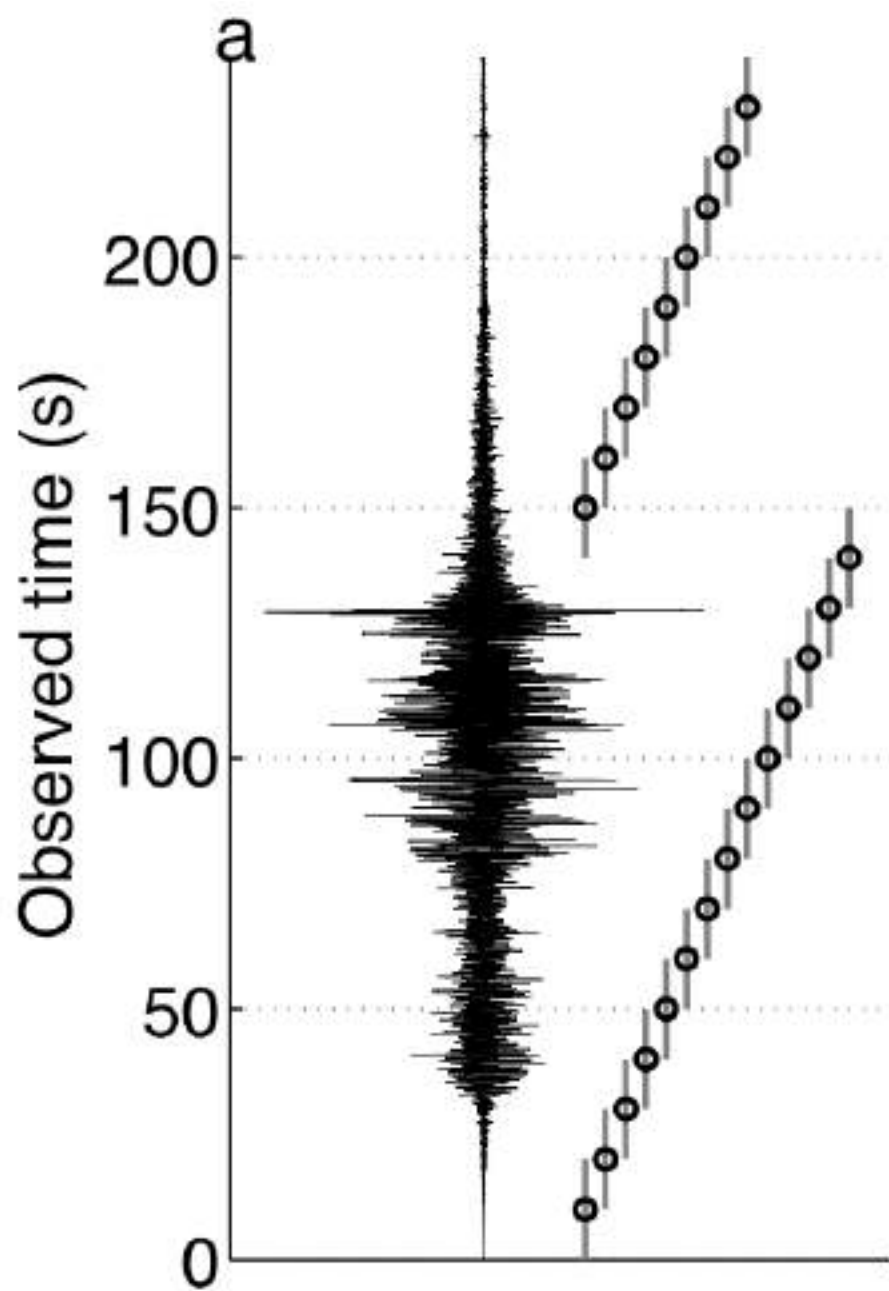
Tohoku-Oki  
earthquake

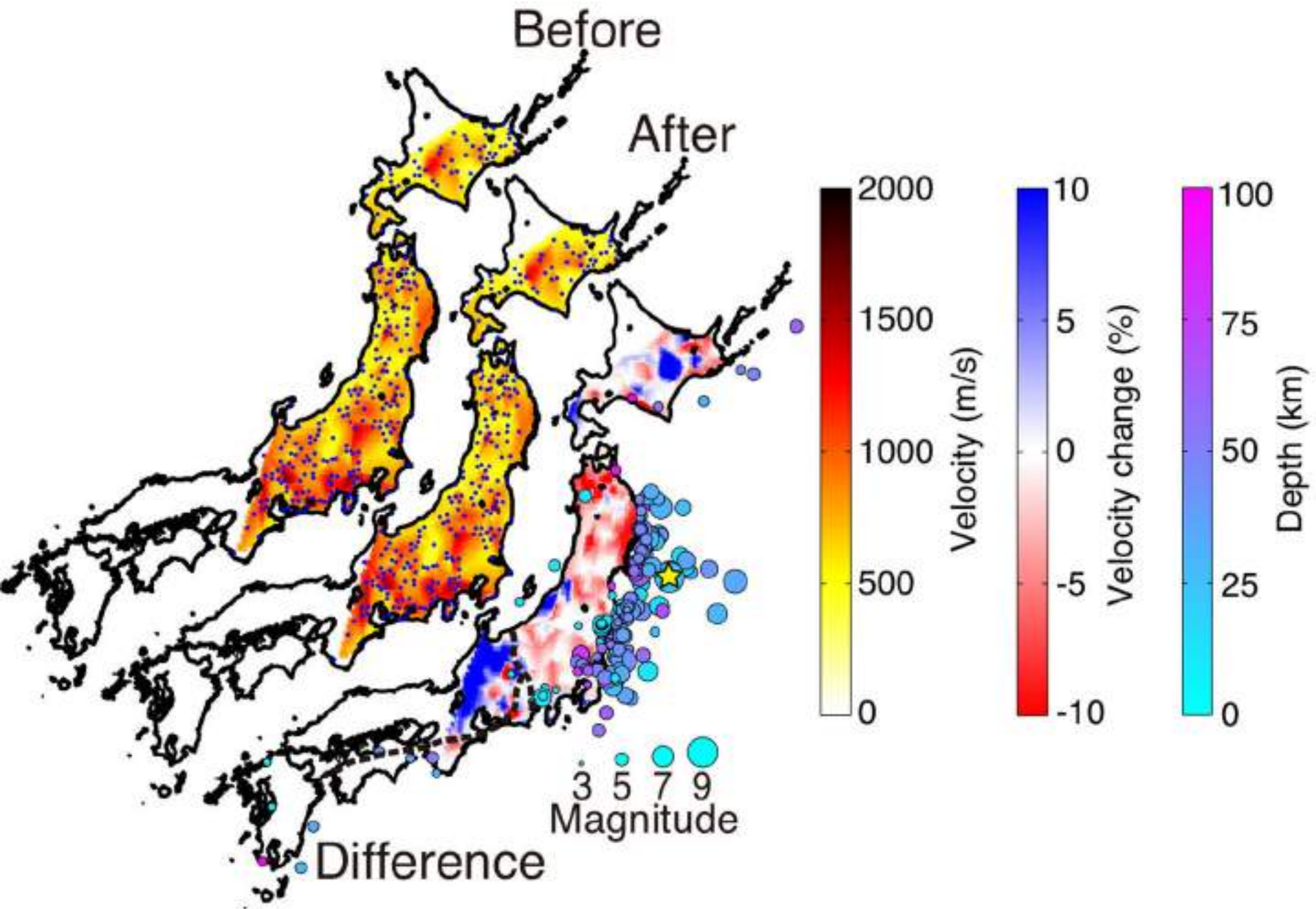


(Nakata and Snieder, Geophys.  
Res. Lett.,18, L17302, 2011)





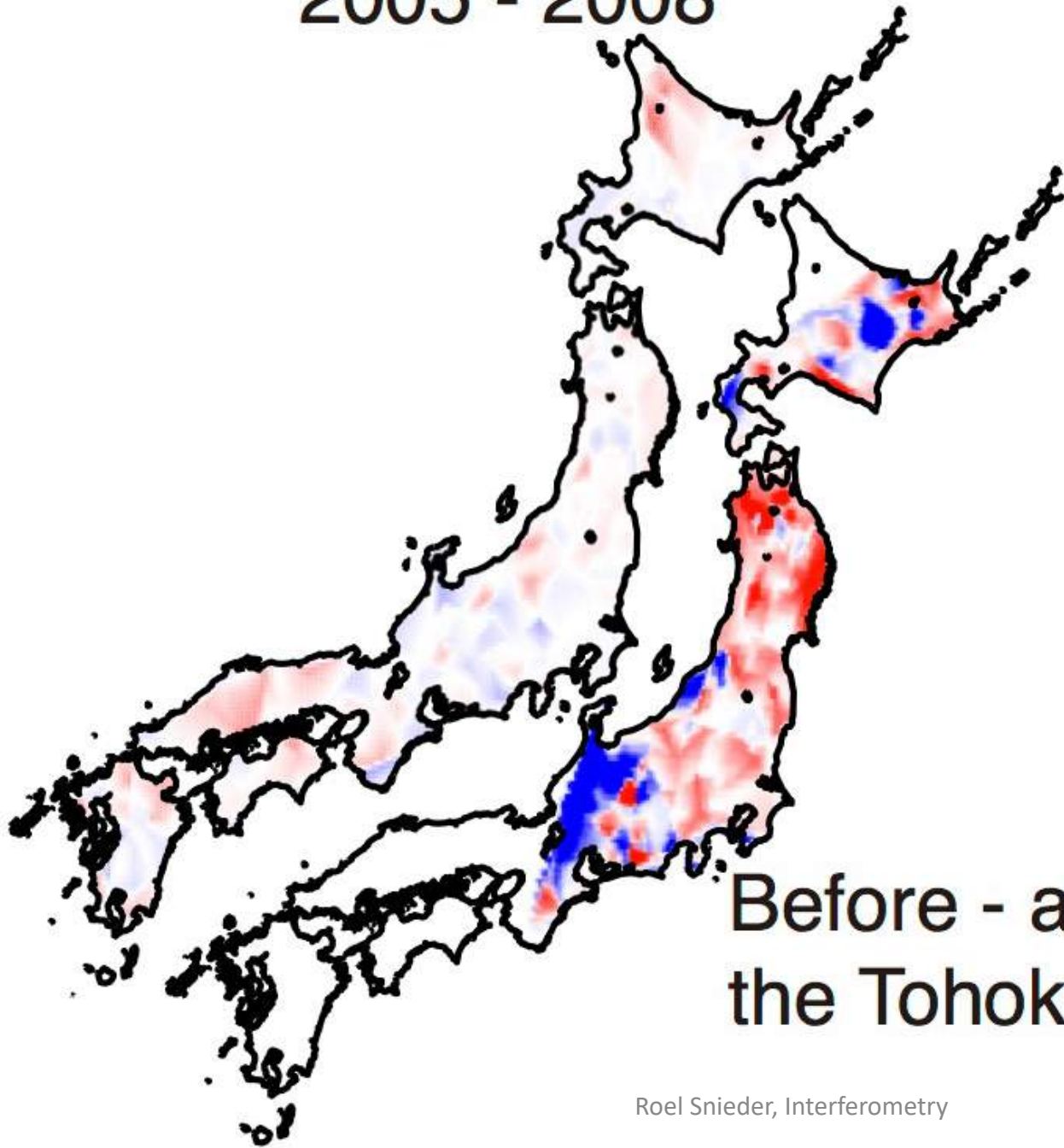




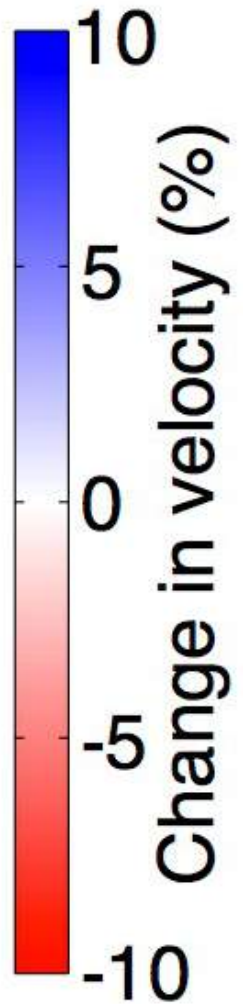
(Nakata and Snieder, *Geophys. Res. Lett.*, 18, L17302, 2011)



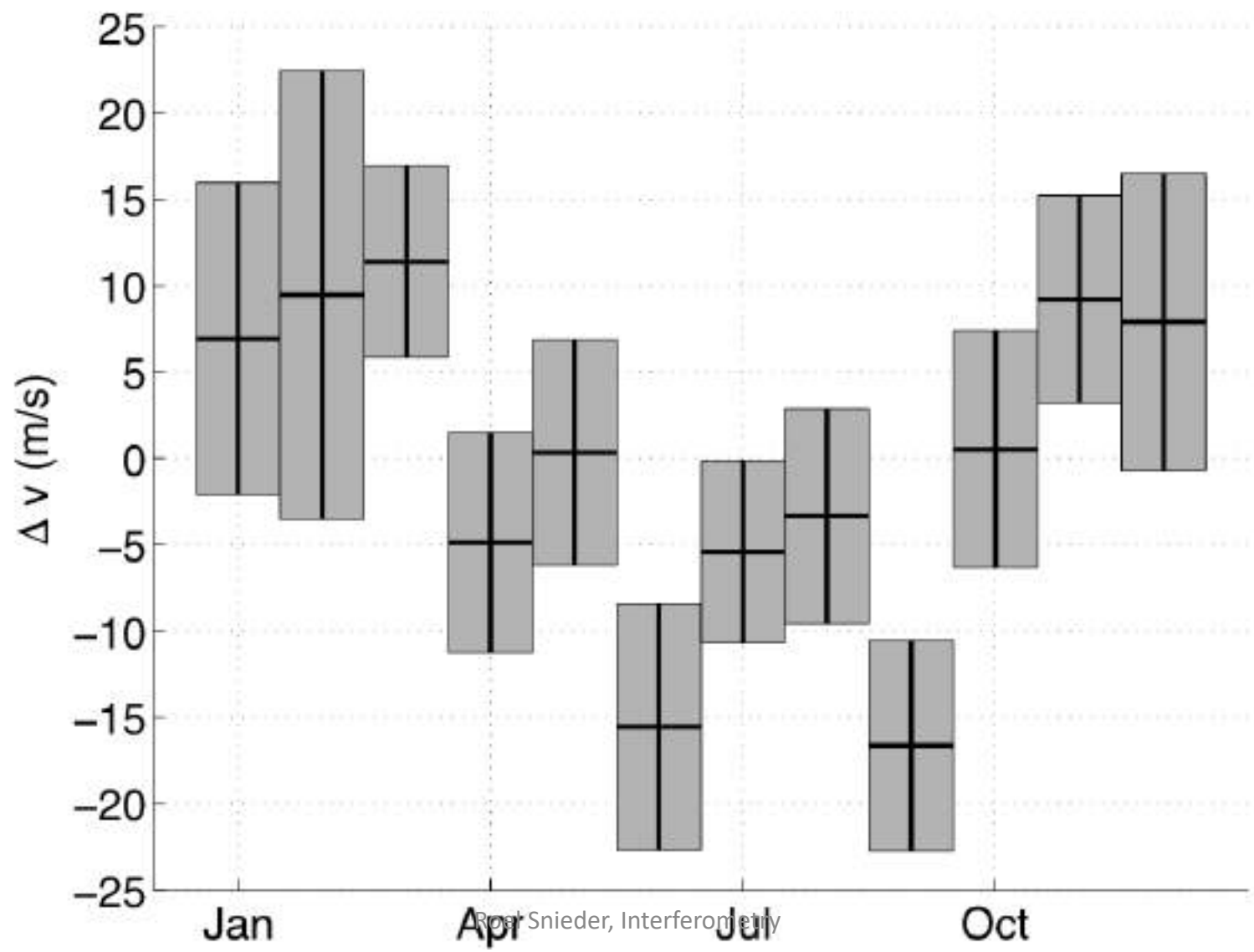
2005 - 2008



Before - after  
the Tohoku EQ

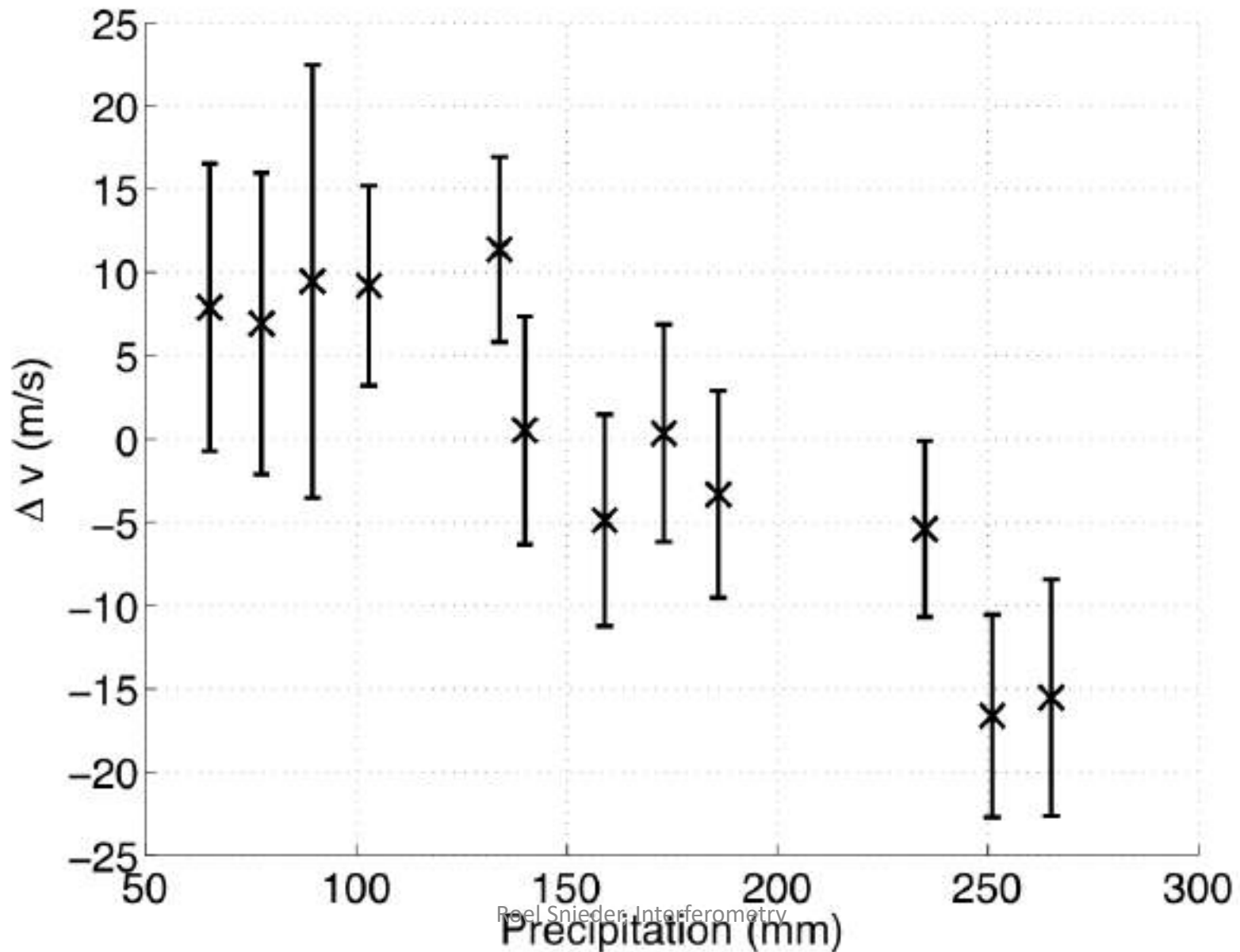


# S-velocity changes with seasons

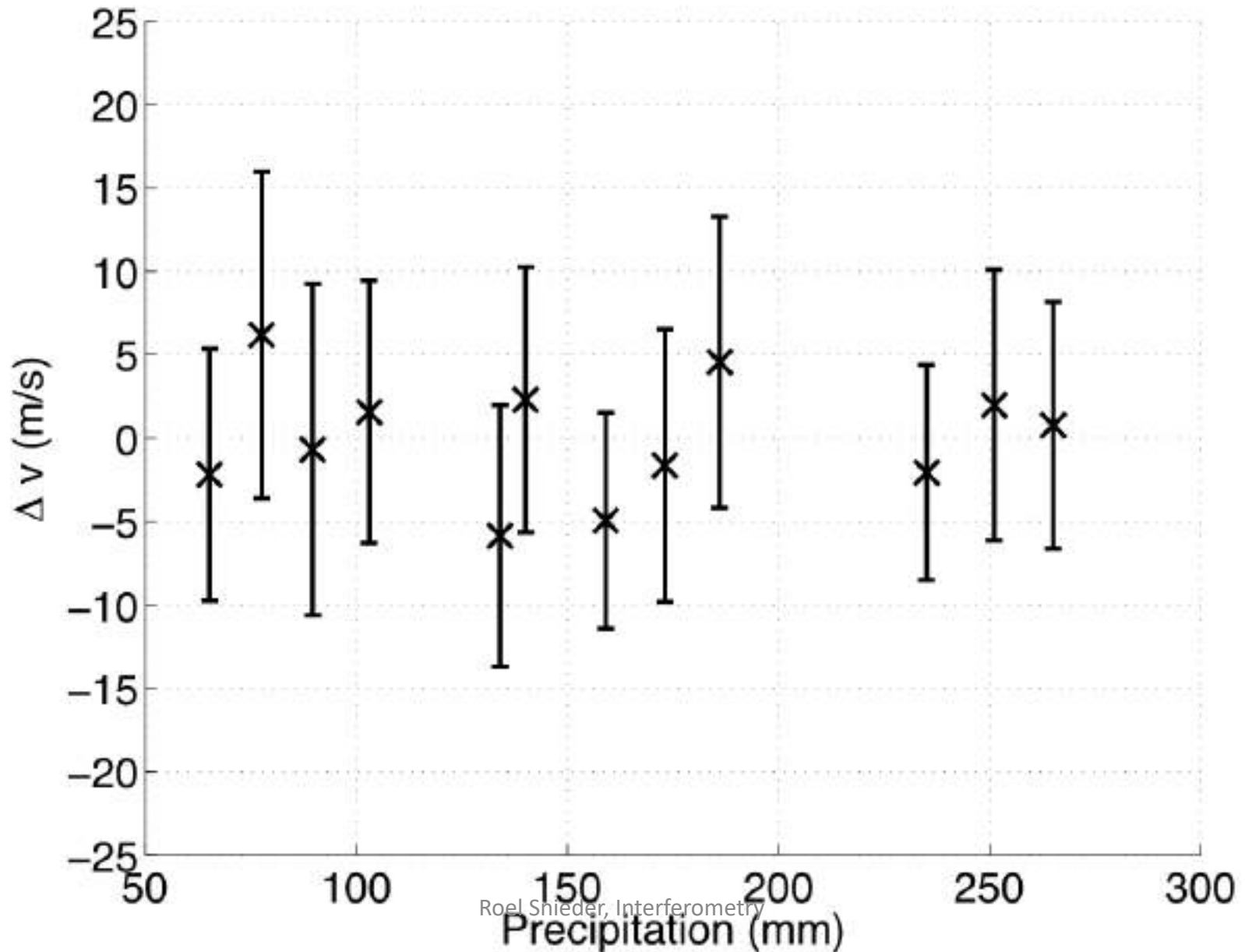


Roel Snieder, Interferometry

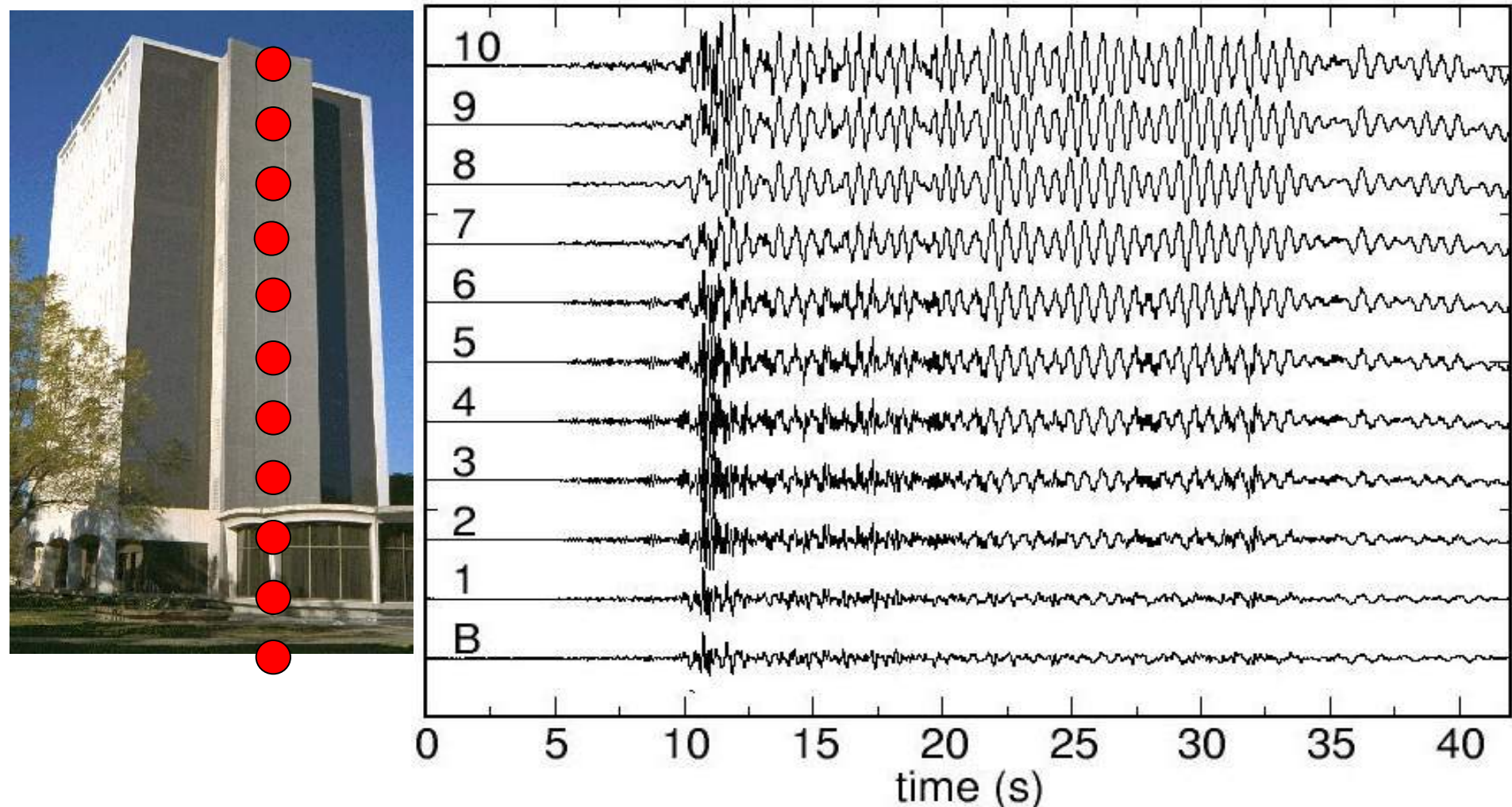
# Rainfall/ $v_s$ for soft-rock sites



# Rainfall/ $v_s$ for hard-rock sites

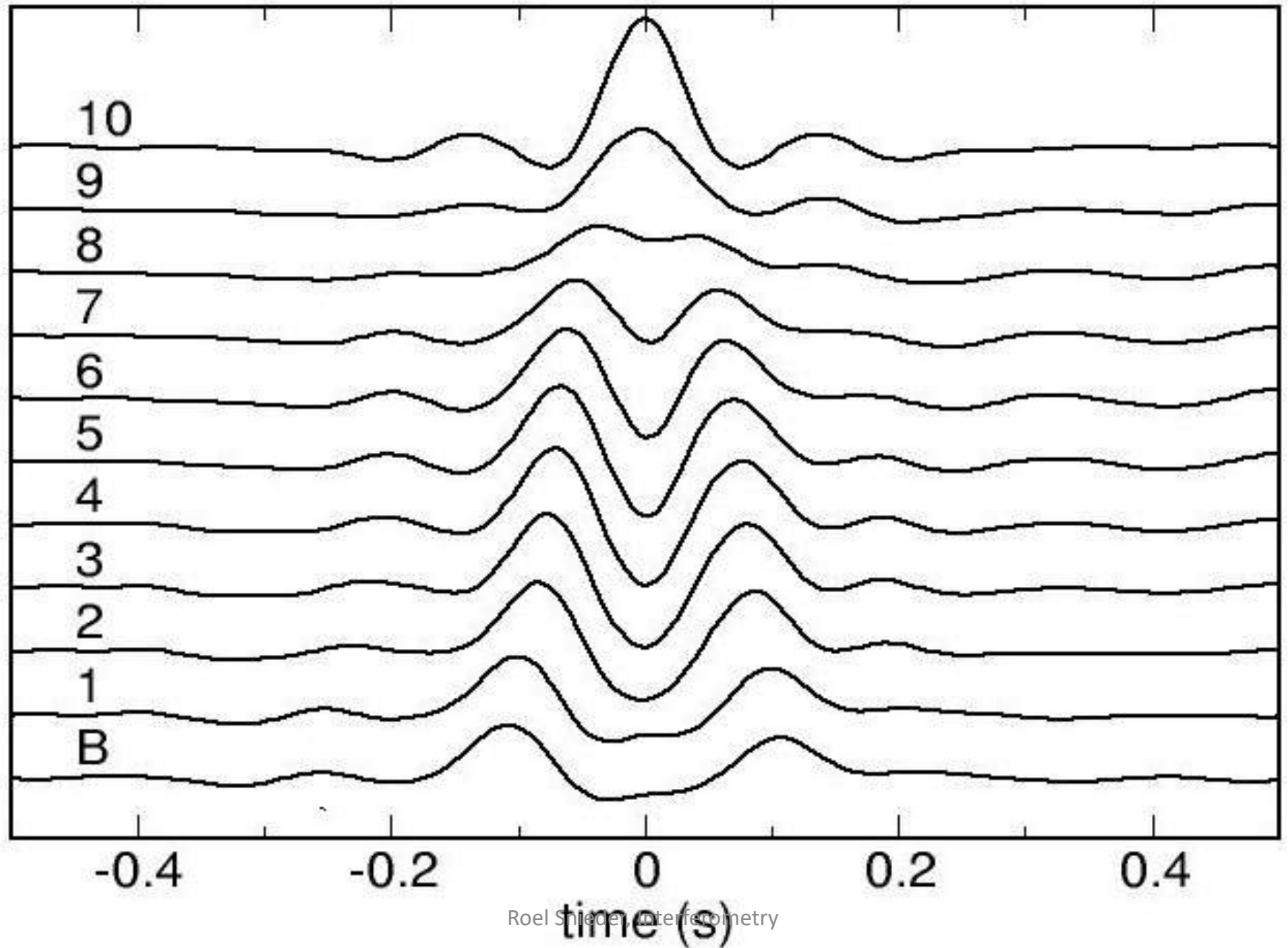


# Seismic interferometry in Millikan Library

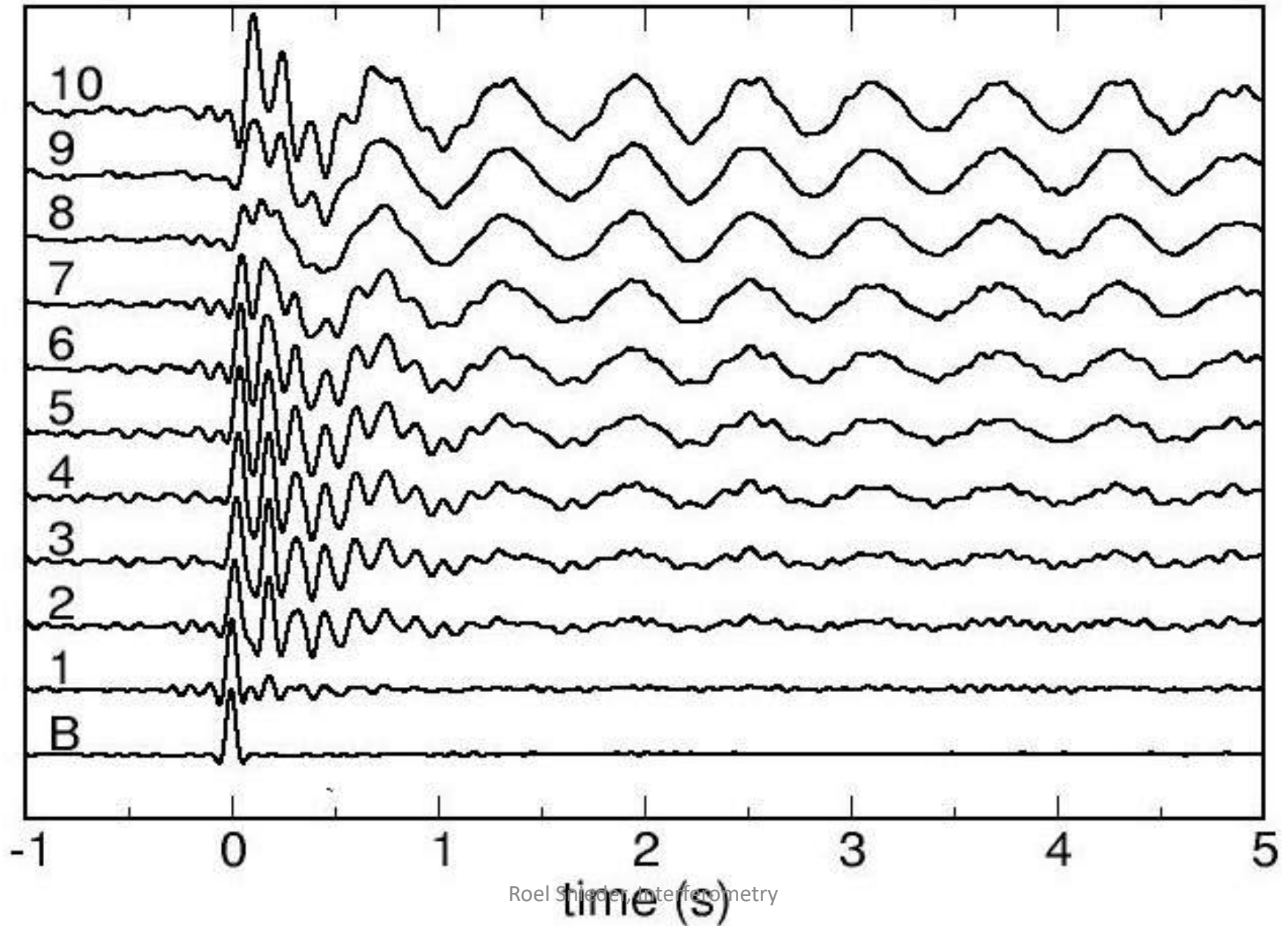


(Snieder and Safak, Bull. Seismol. Soc. Am., **96**, 586-598, 2006)

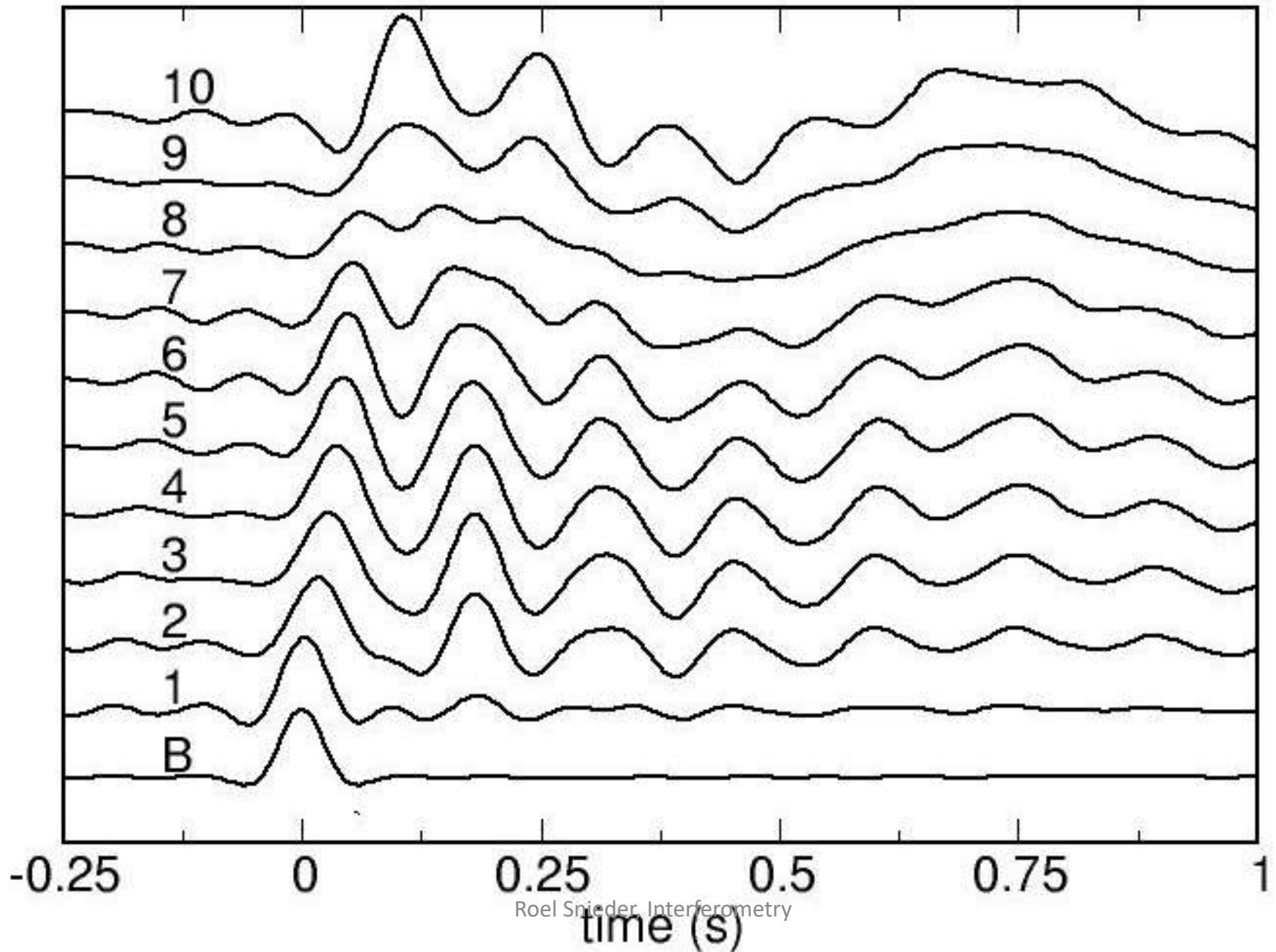
# Deconvolution with top floor



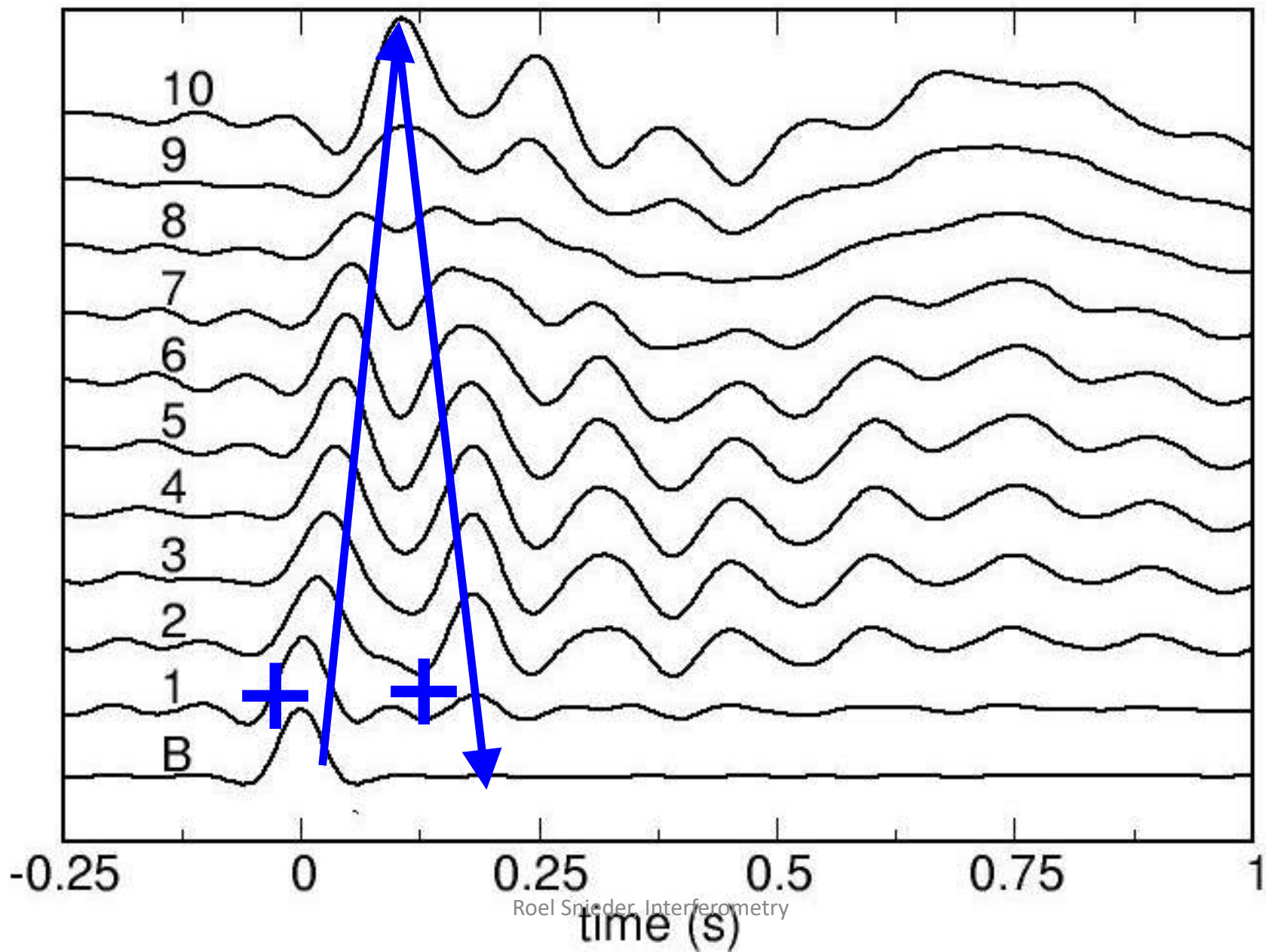
# Deconvolution with bottom floor

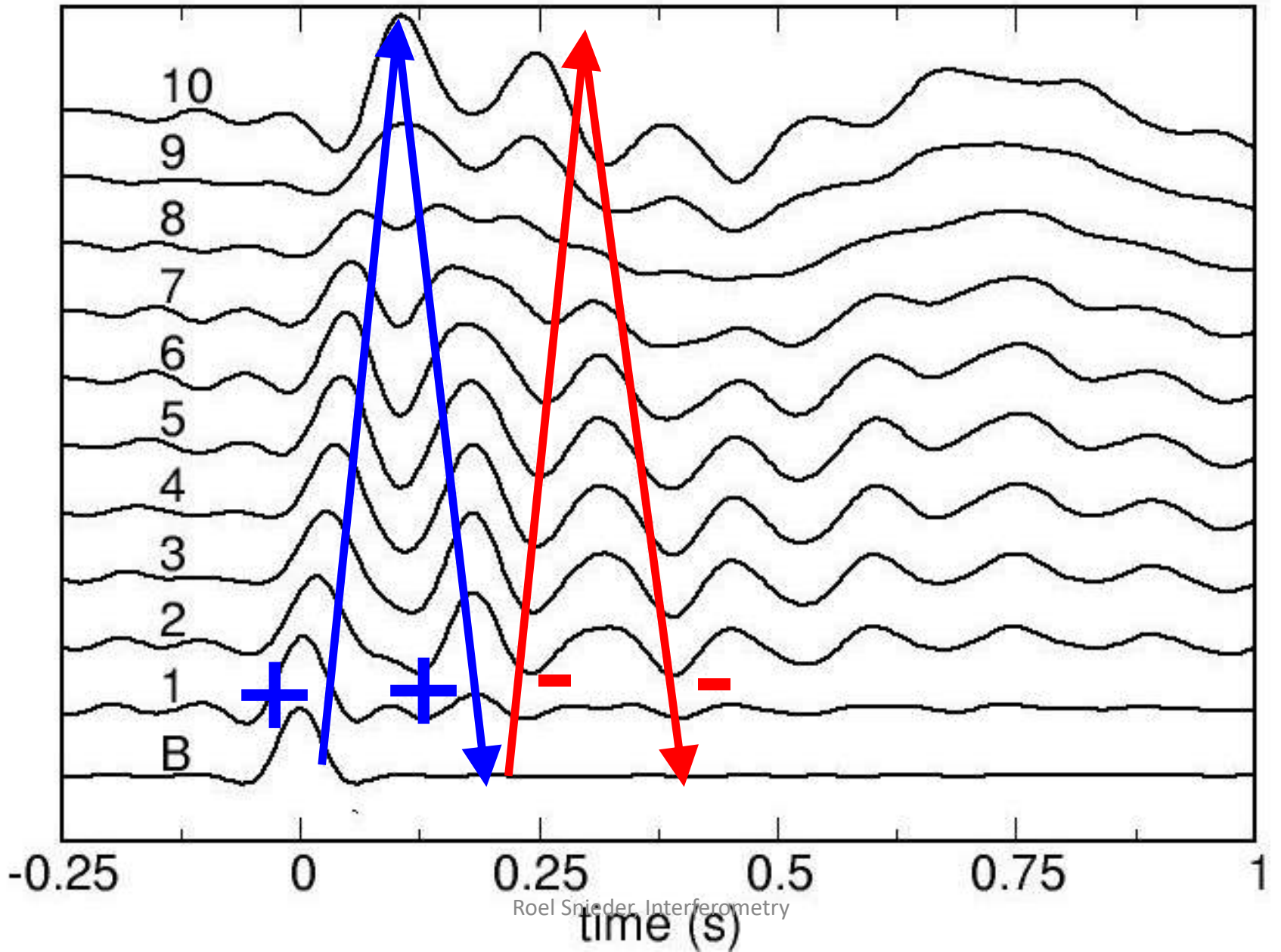


# Deconvolution with bottom floor

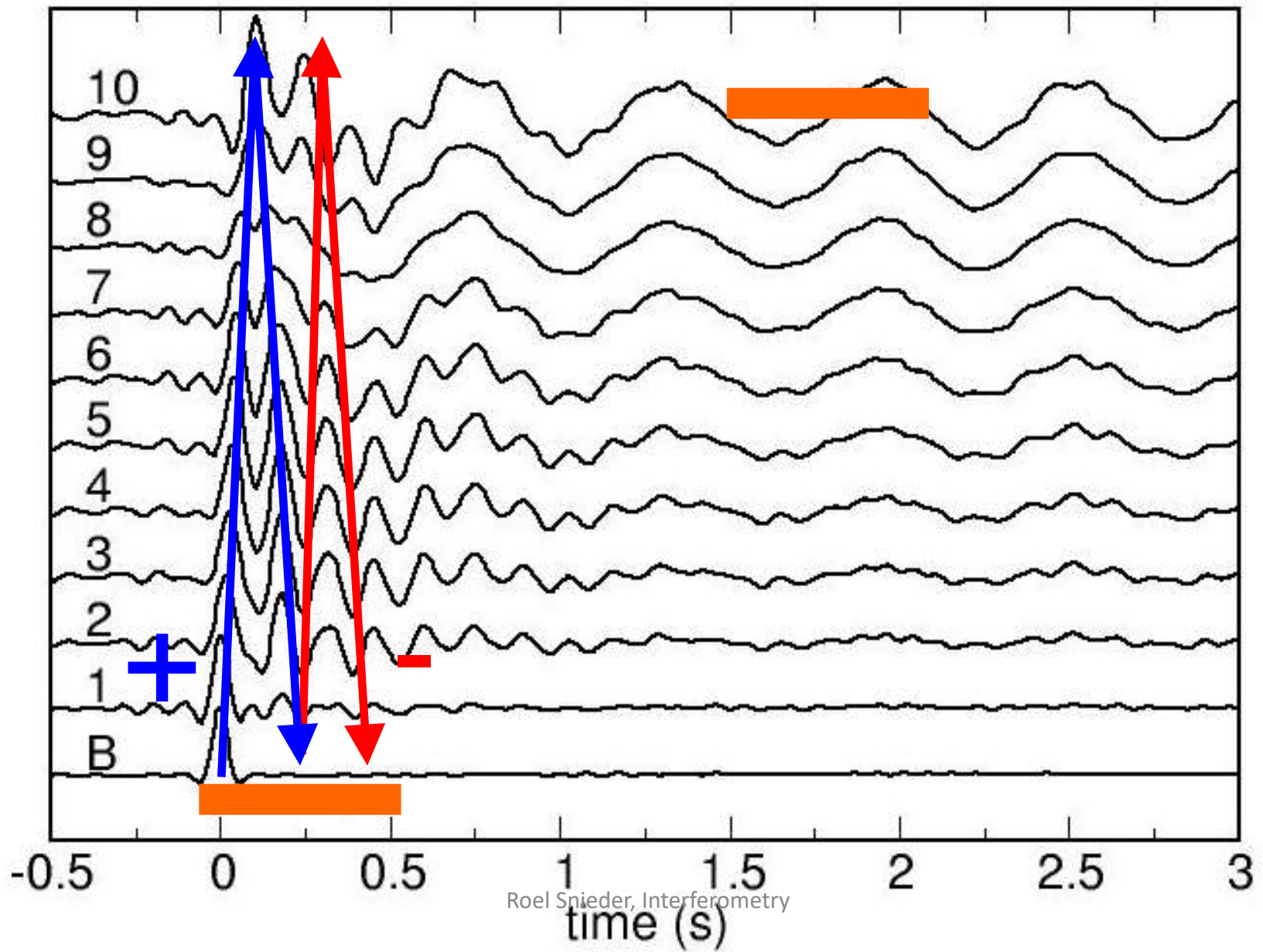






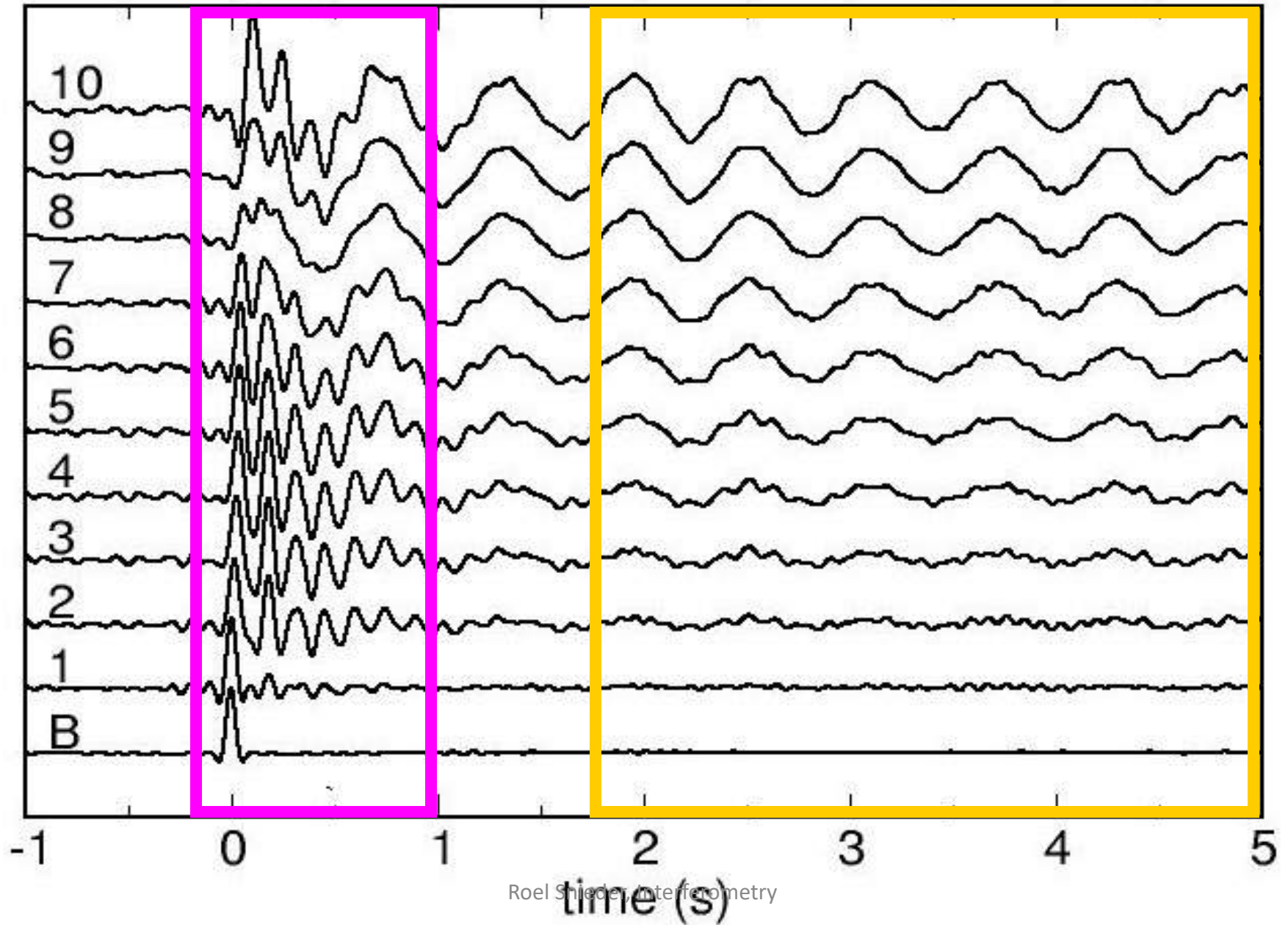


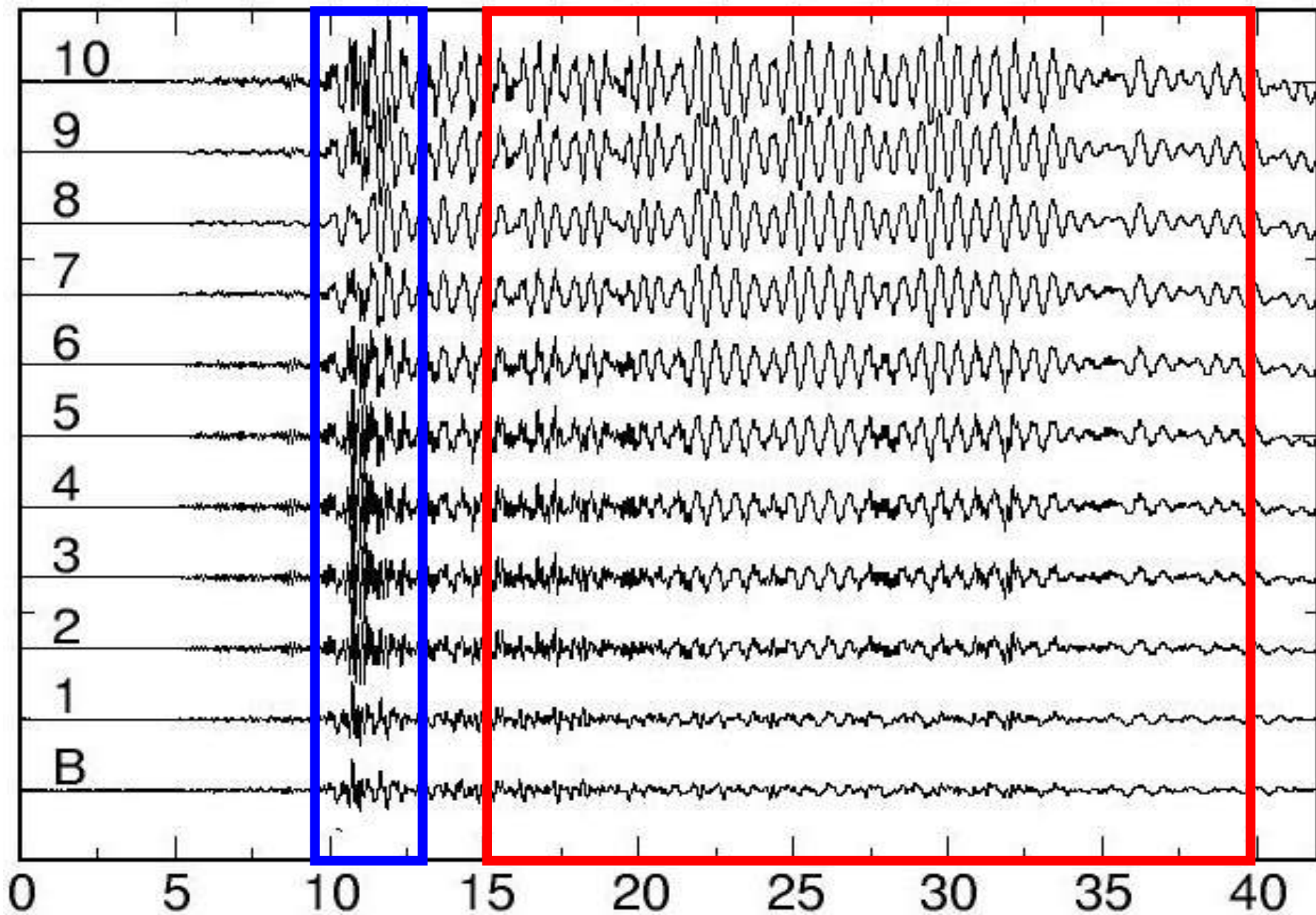
Fundamental mode:  $T = 4H/c$

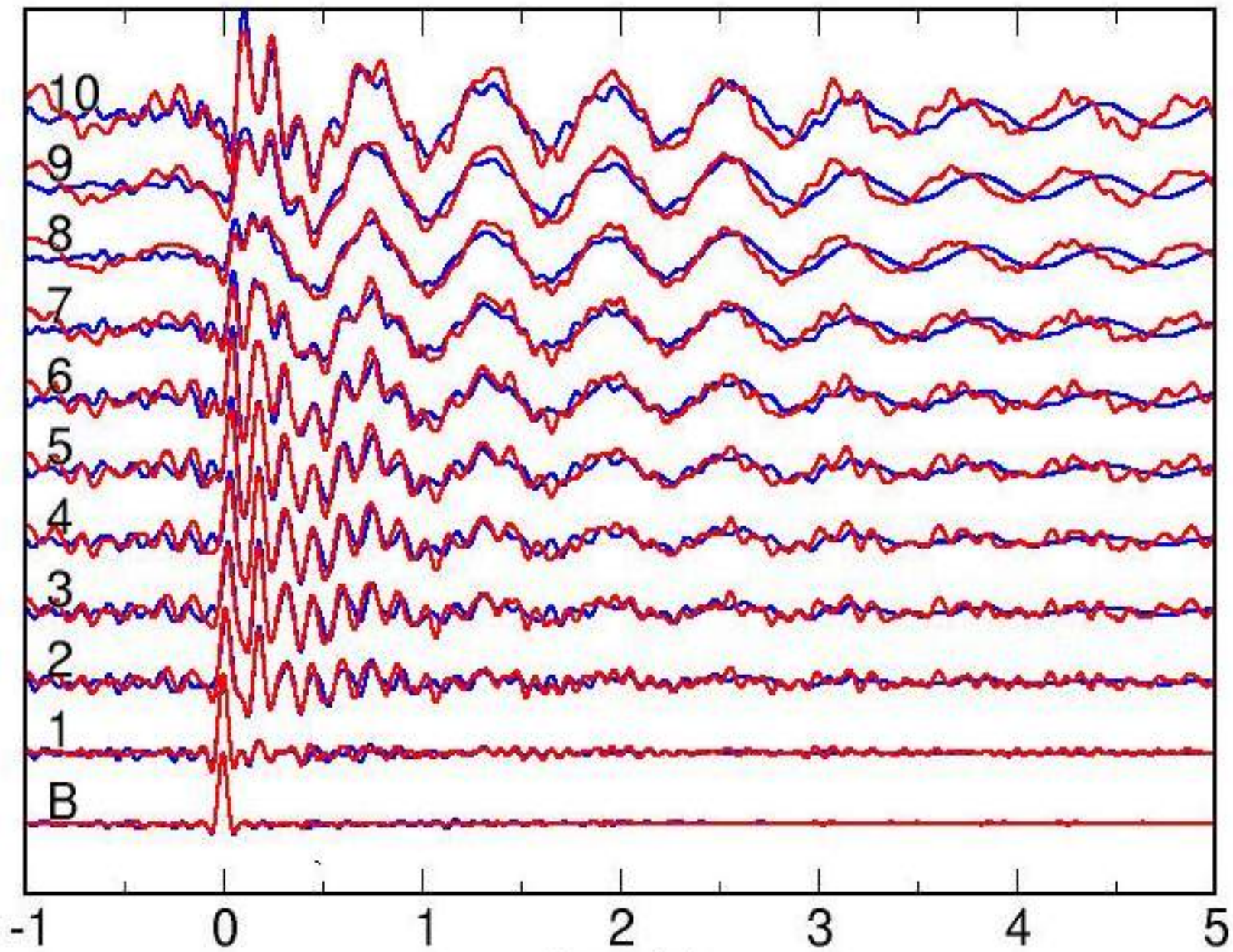


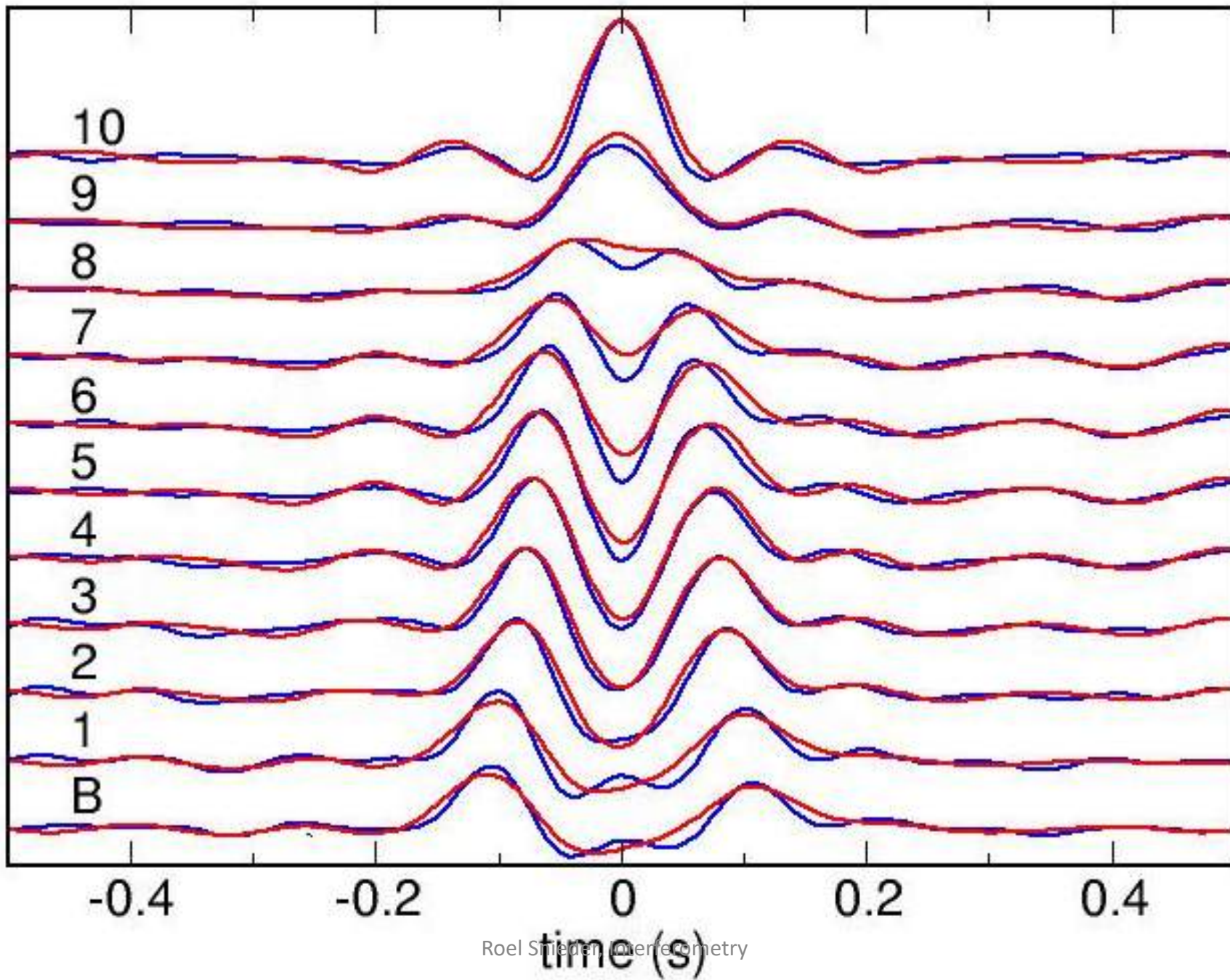
traveling waves

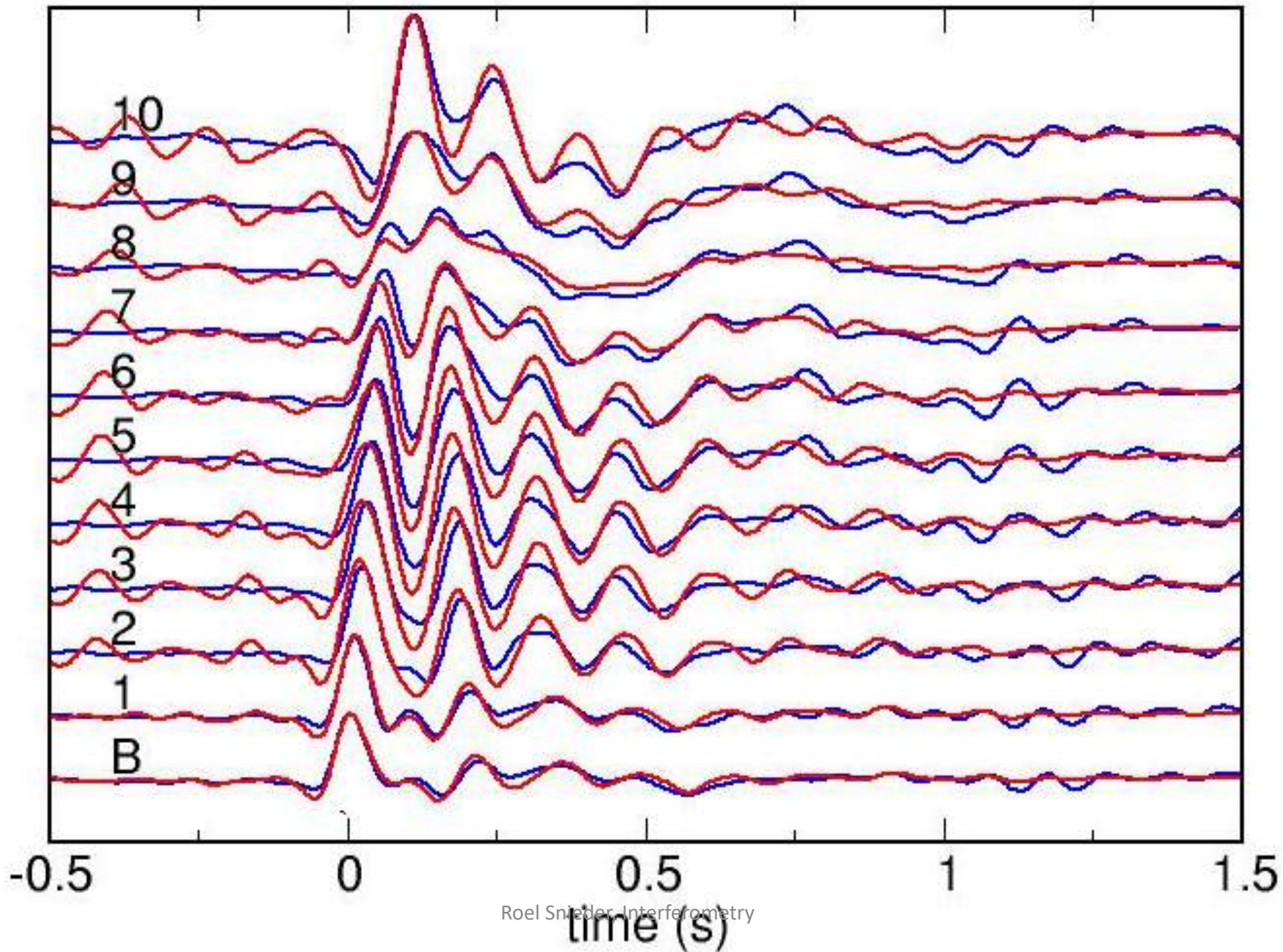
normal modes





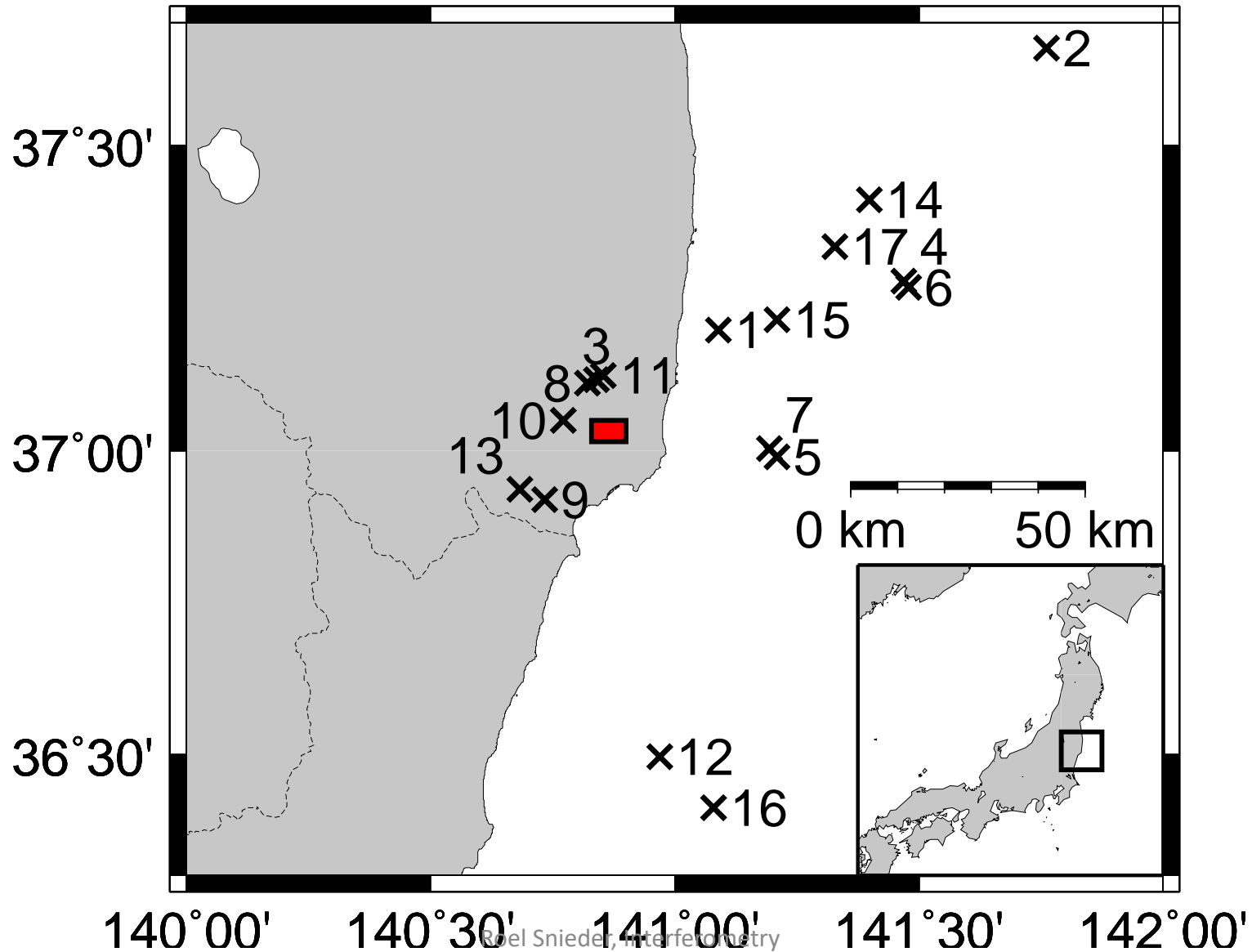








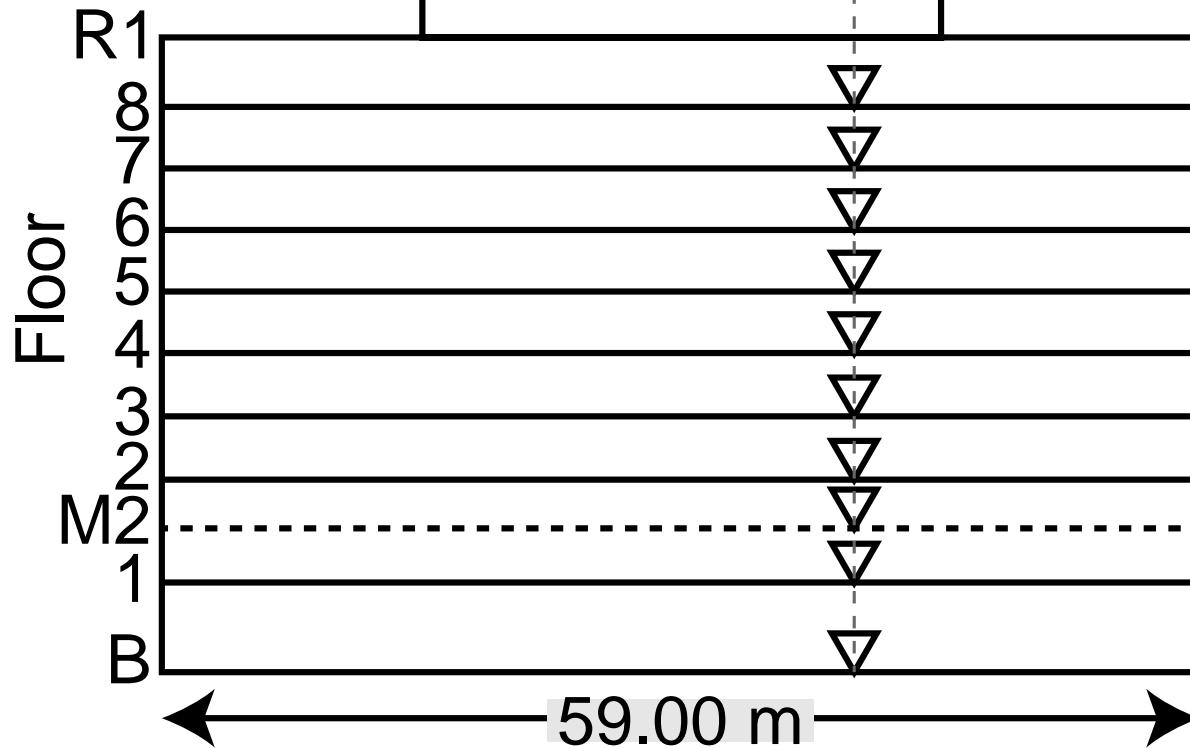
# Location of building and earthquakes



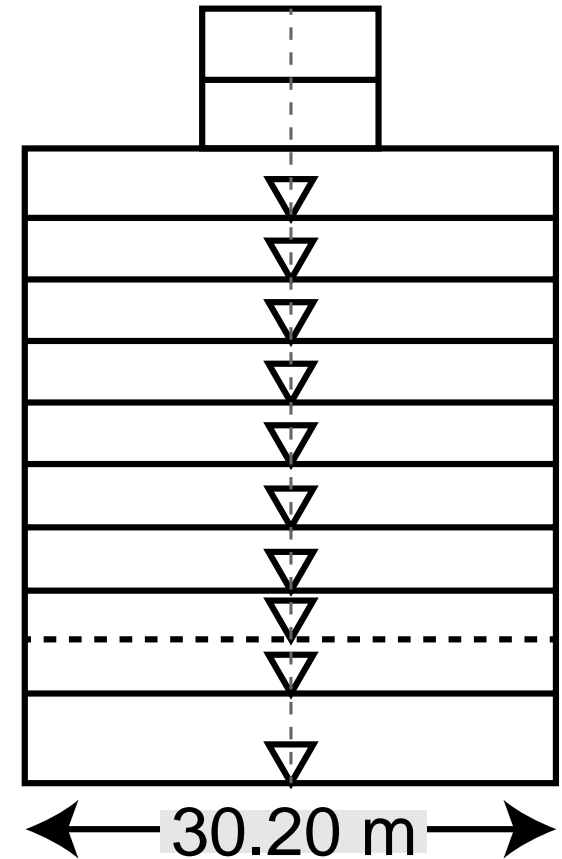
# Building and accelerometers

East-West

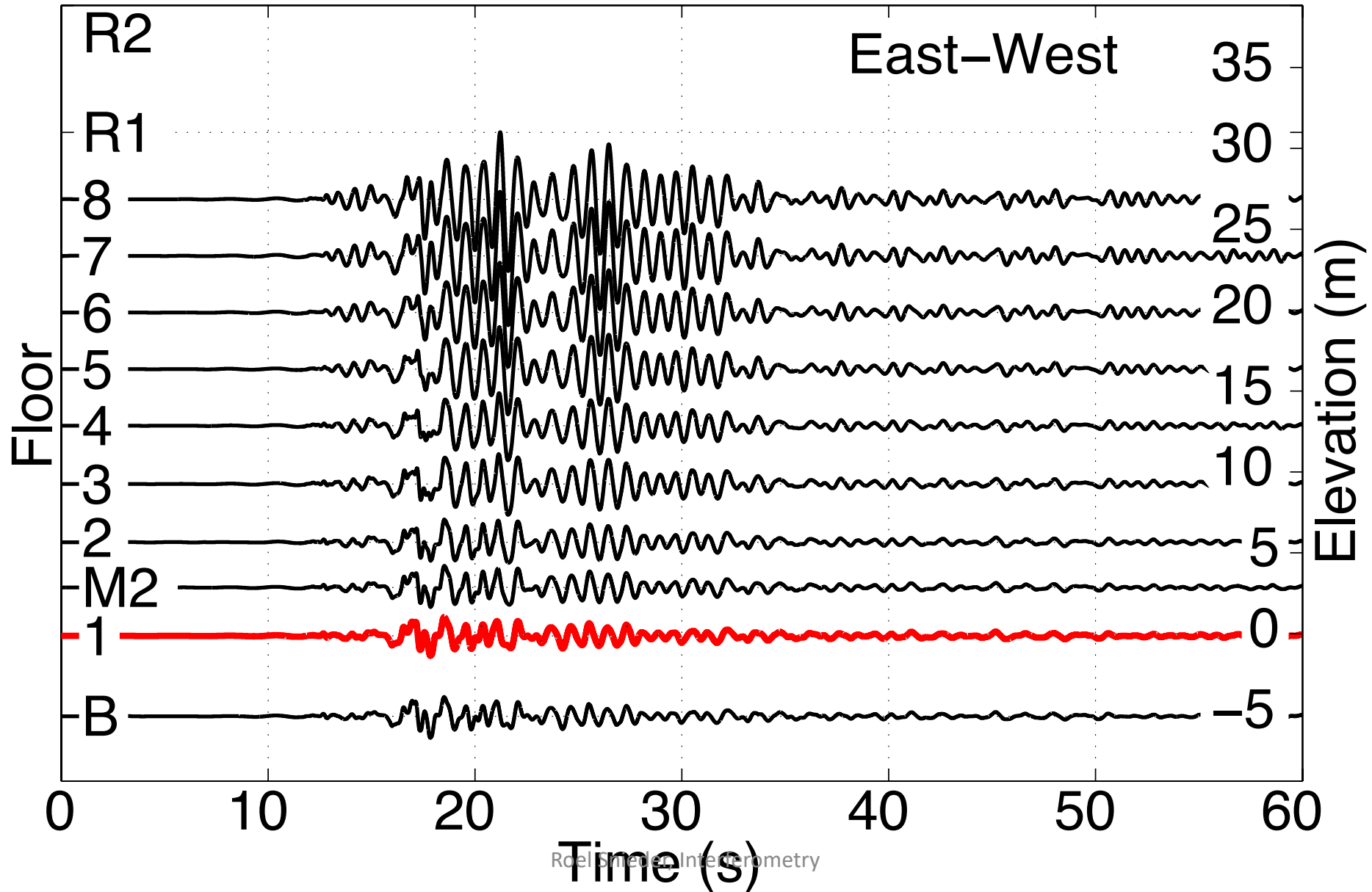
R2 (39 m)



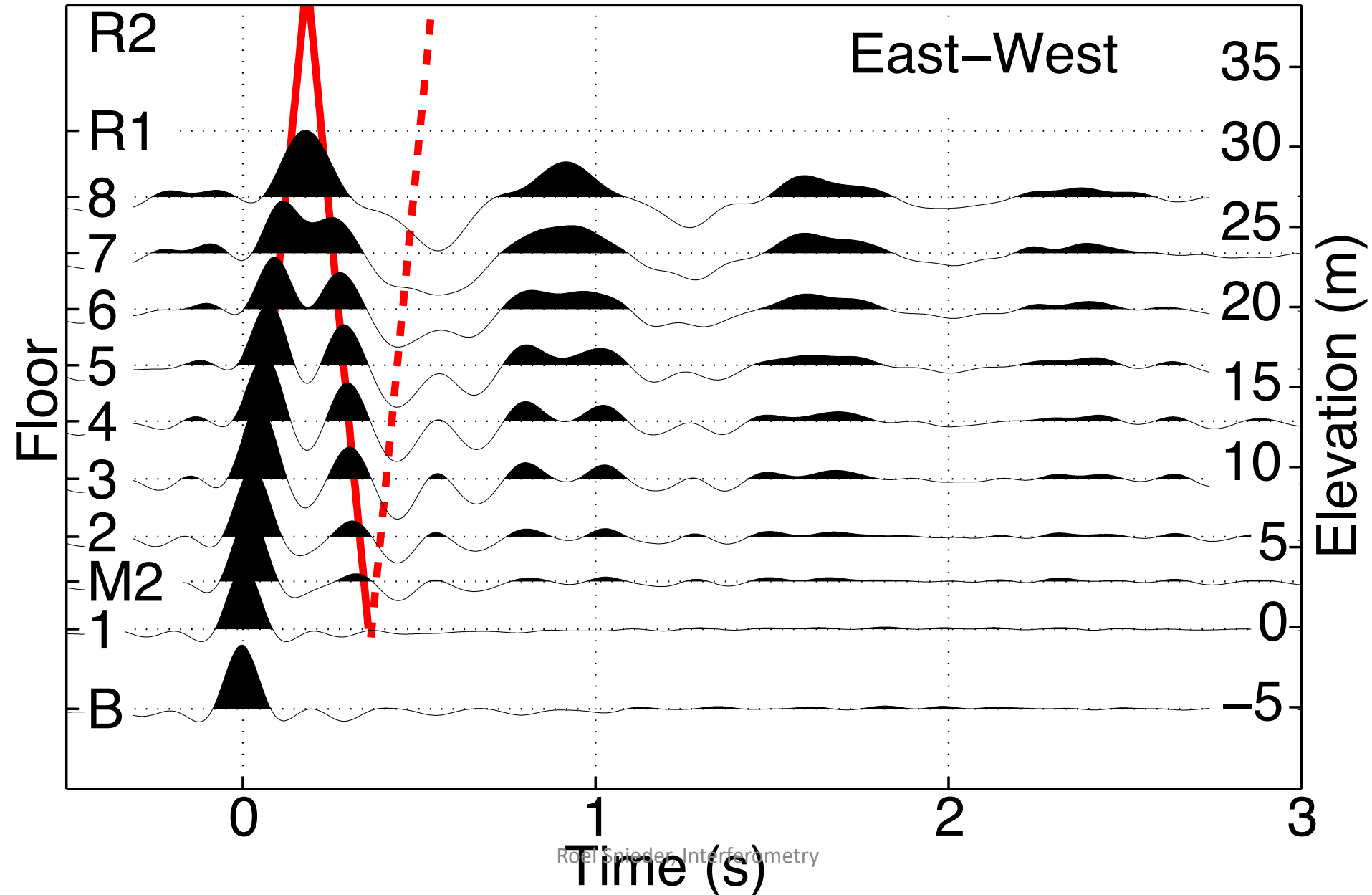
North-South



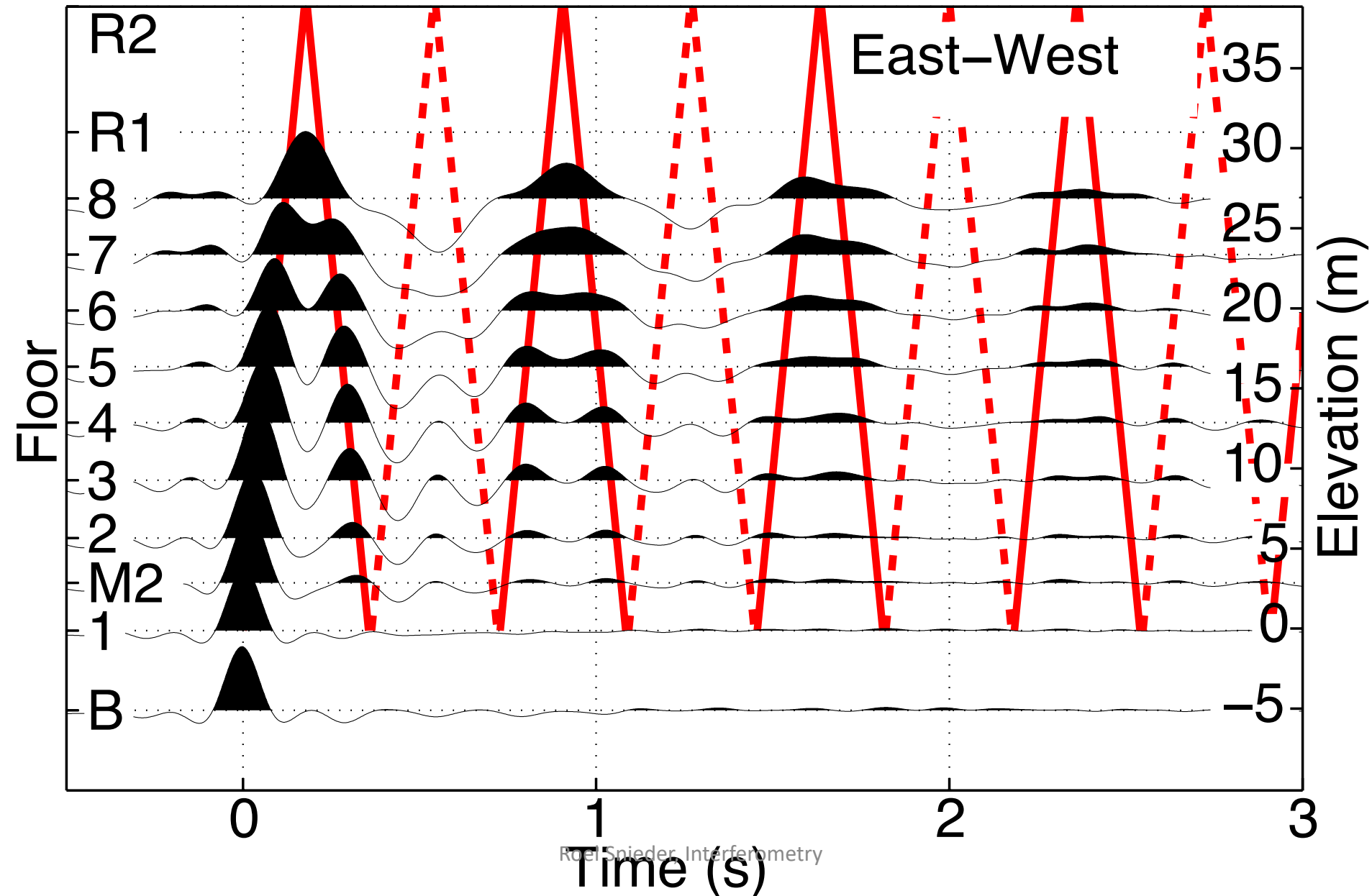
# Recorded motion



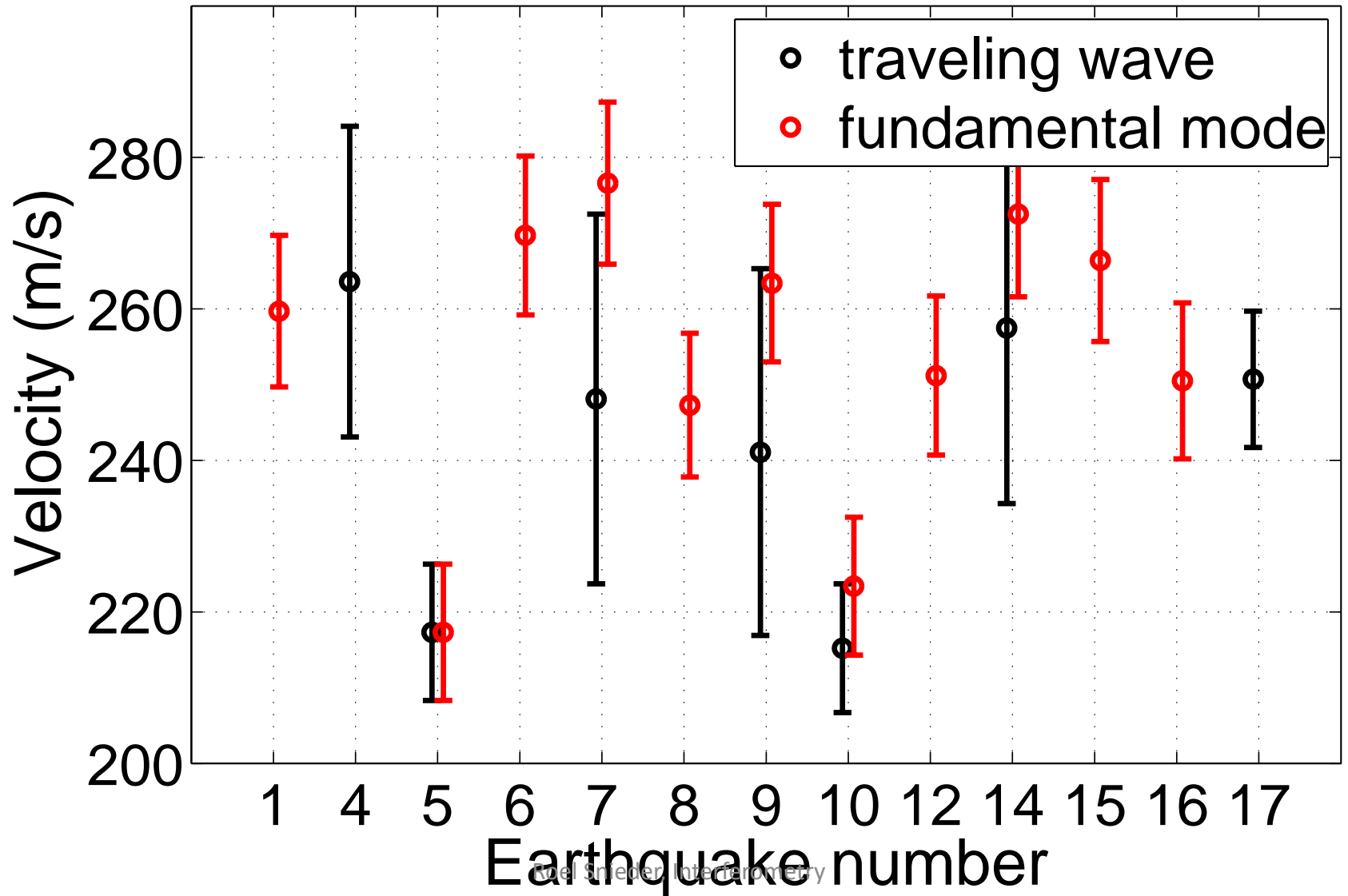
# Earthquake after deconvolution



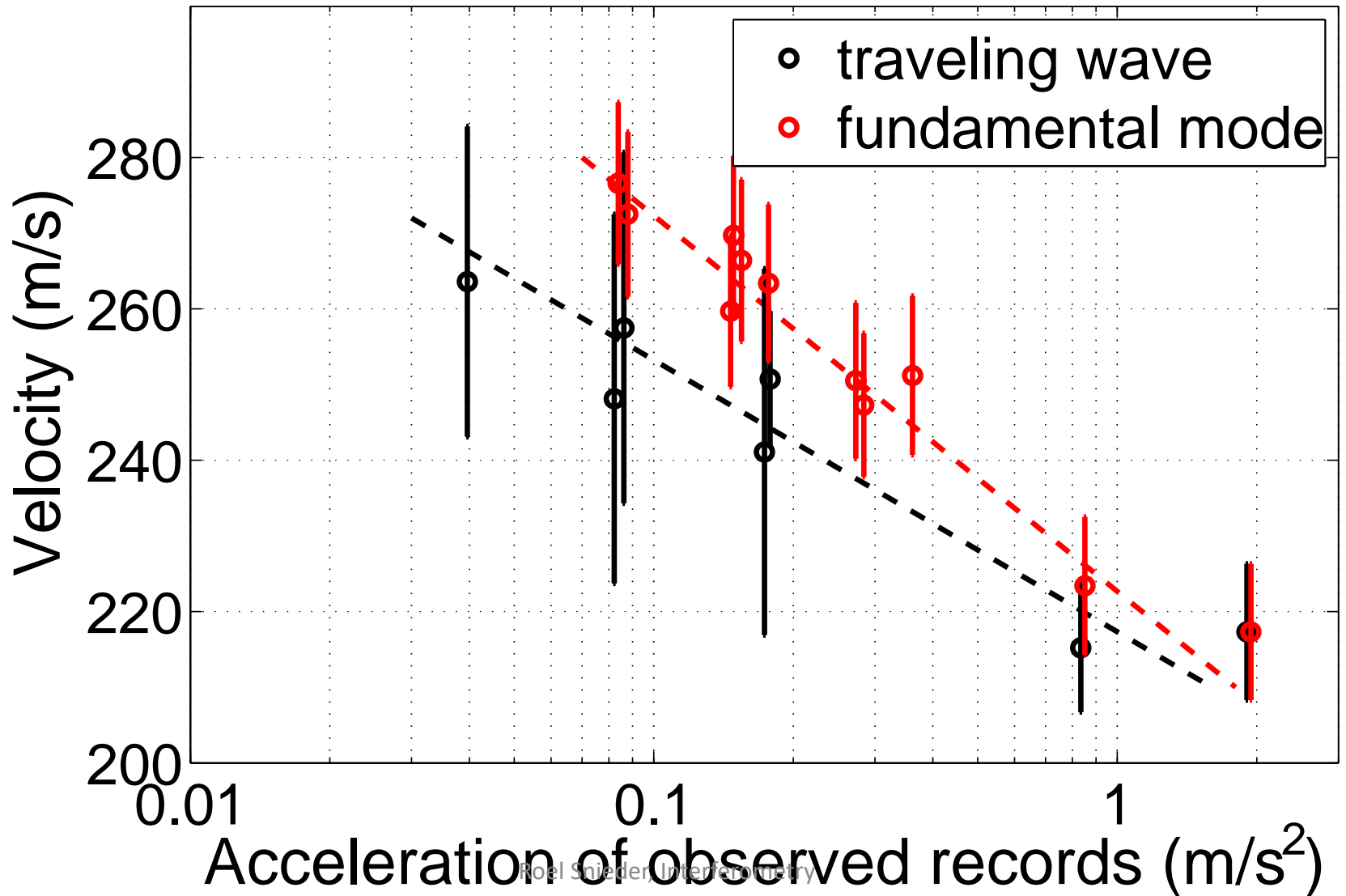
# Earthquake after deconvolution



# Shear velocity from different quakes



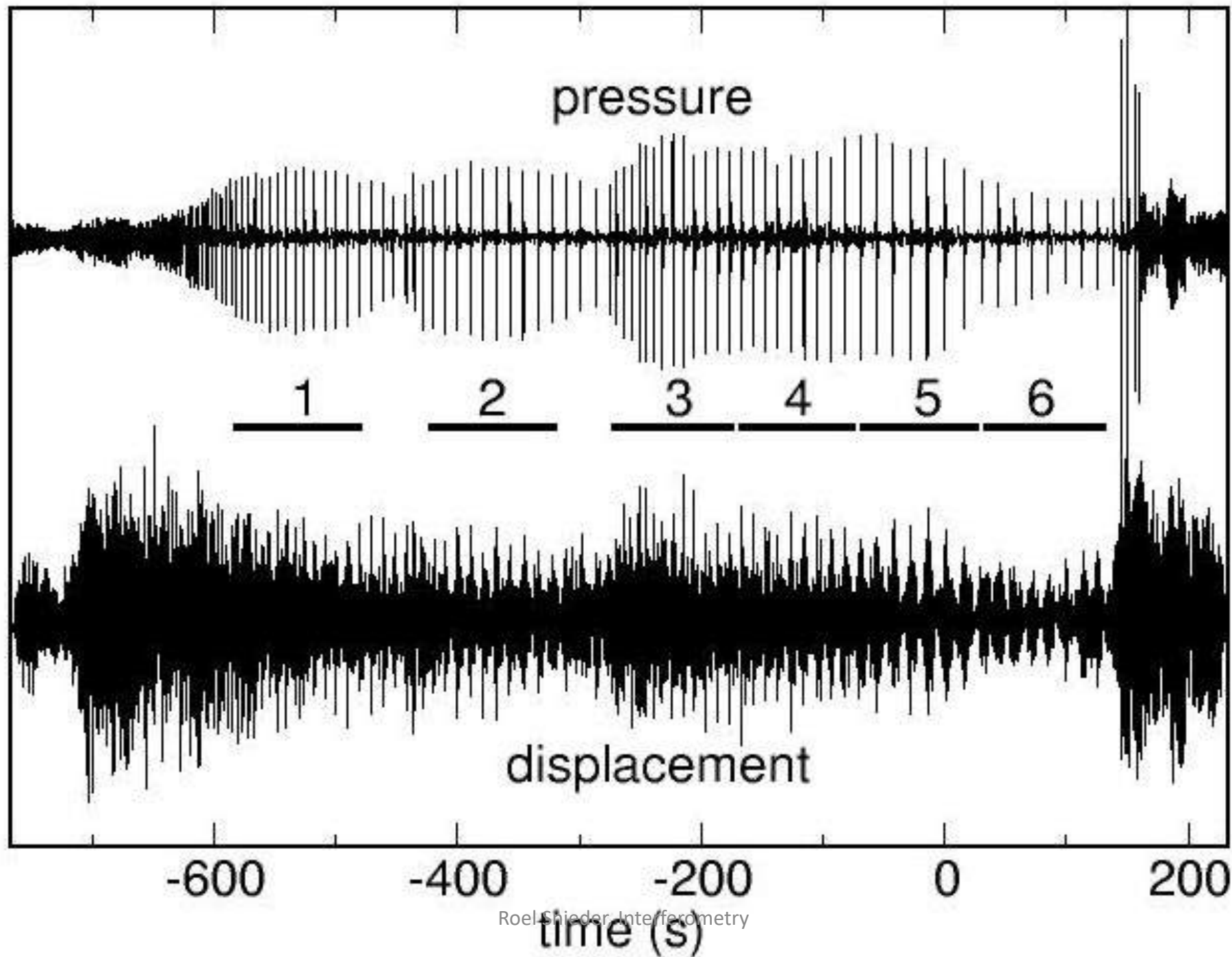
# Shear velocity vs. shaking



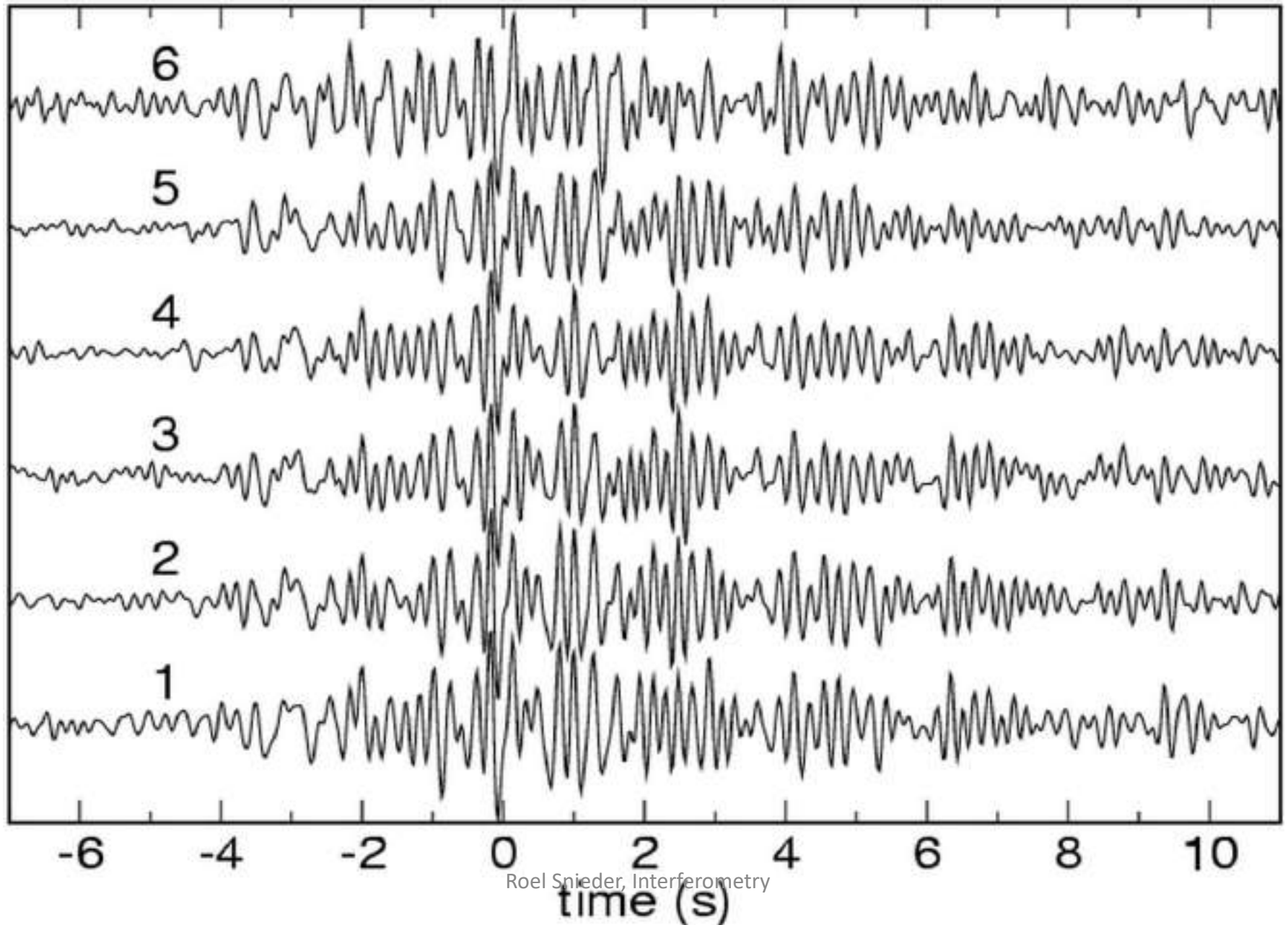
# When the source-timing can be used

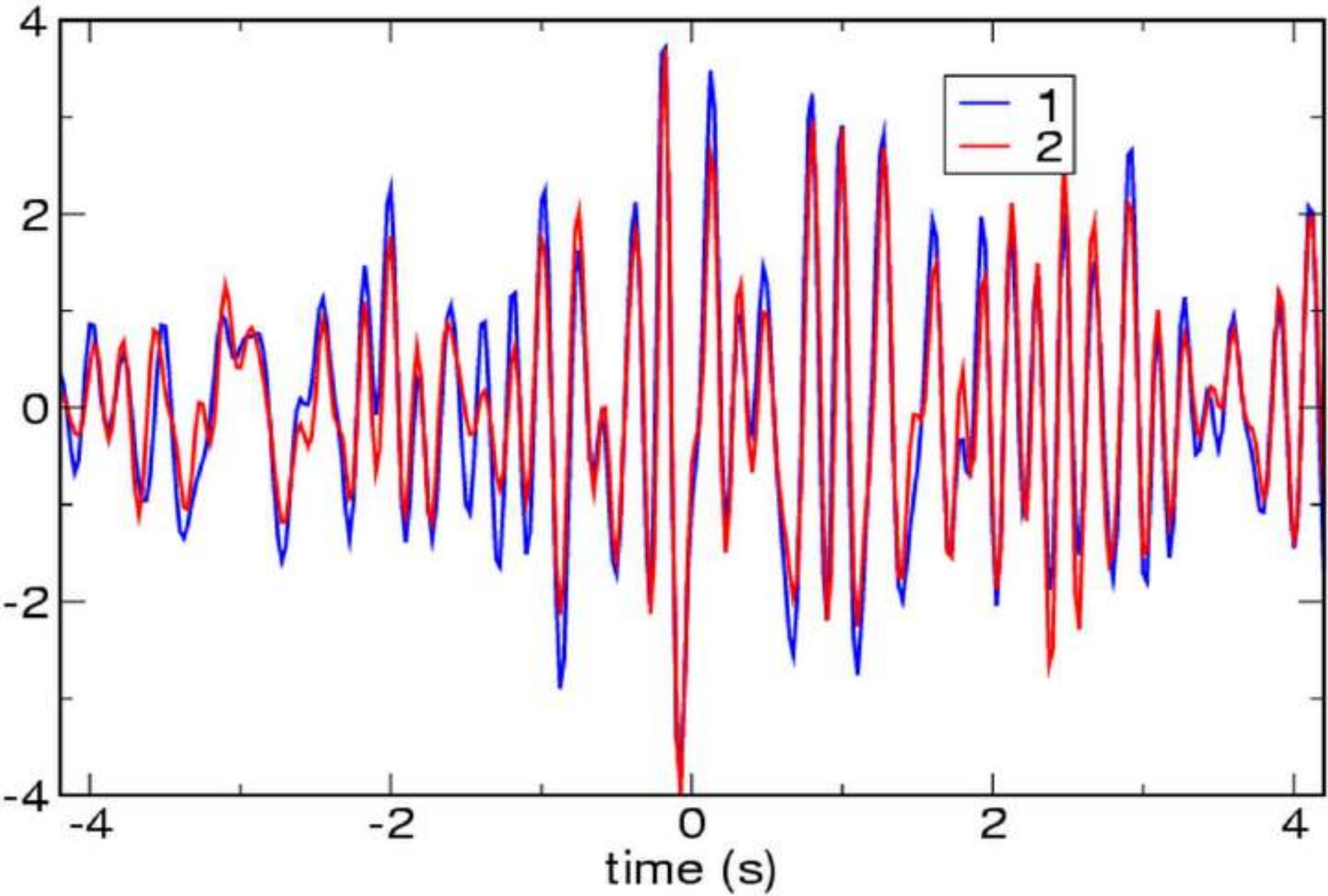






# Displacement deconvolved with pressure





(Snieder, R. and M. Hagerty, *Geophys. Res. Lett.*, 31, L09608, 2004)

# Talsperre Eibenstock



## Ultraschallsensorik aus der Bauphase

- ▶ Entwicklung des E-Moduls/ Druckfestigkeit
- ▶ 8 Geber, 4 Empfänger (UNG40/SW40, 40 kHz)
- ▶ VEB Projektierung Wasserwirtschaft



