

# Effect of shallow heterogeneities on wavefield gradients measurements

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Funded by the European Union's Horizon 2020 research and innovation programme  
under the Marie Skłodowska-Curie grant agreement No. 955515.



# How sensitive are the wavefield gradients to shallow localized velocity changes?

In this study we **investigate the sensitivity of the wavefield and the wavefield gradient measurements** to shallow localized velocity changes.

**Seismic simulation software**

Spectral element simulation software (SEM46<sup>1</sup>) modified: rotations and strains as a direct output

**Experiments performed**

Seismic array placed above a velocity anomaly

**Weak anomaly**  
(10% velocity drop)

**Strong anomaly**  
(70% velocity drop)

<sup>1</sup> Brossier et al. 2019, "Efficient time-domain 3D elastic and viscoelastic full-waveform inversion using a spectral-element method on flexible Cartesian-based mesh"



We performed each experiment two times, the first time considering a fully homogeneous medium, the second time including the seismic anomaly.

## How a seismic anomaly interacts with the wavefield and its observables?

### Phase shift

The seismic phases acquire a delay or advance with respect to the homogeneous case.



Very precise methods to measure it by frequency-time decomposition, adapted to slight velocity variation, the continuous wavelet transform<sup>2</sup>

### Amplitude variations

The normalized waveform changes its amplitude when the anomaly is placed in the medium



Comparison of the amplitude in the time domain

**Necessary for seismic monitoring, our goal!**

<sup>2</sup>Mao et al 2019, "On the measurement of seismic traveltimes changes in the time-frequency domain with wavelet cross-spectrum analysis"



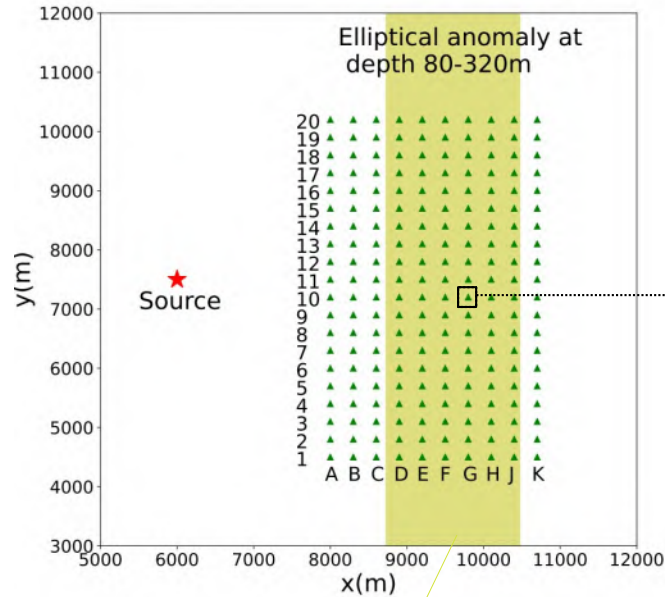
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# Weak anomaly experiment

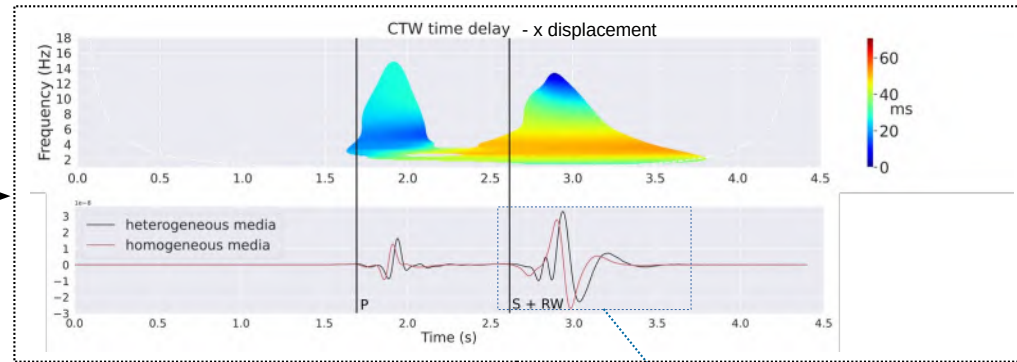
Continuous wavelet analysis

## Experiment configuration



10% velocity drop

Cross wavelet transform. UP: delay observed in the time frequency domain.  
BOTTOM: waveforms in the time domain



Measurement of the Rayleigh waves delay

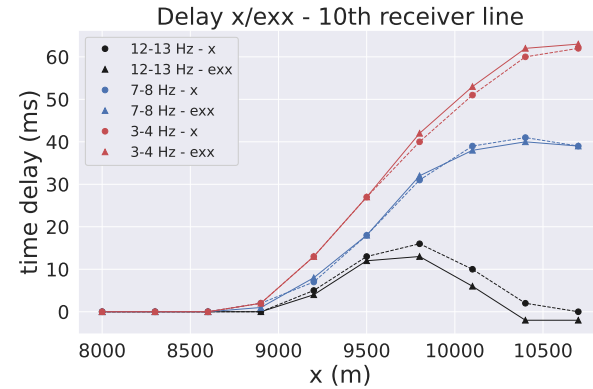
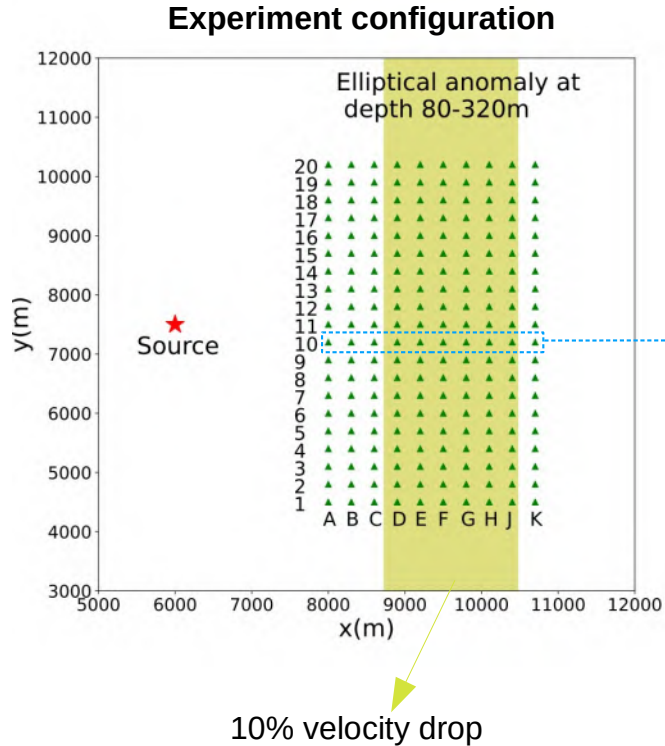


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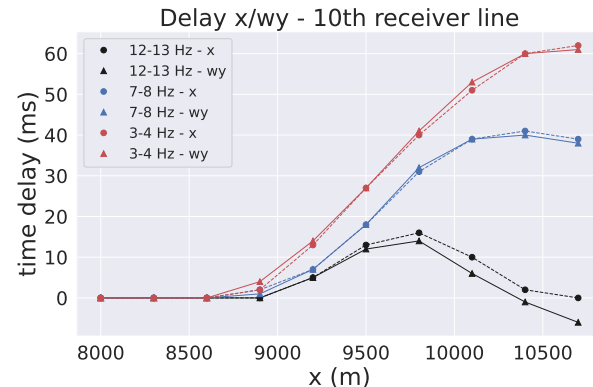


# Weak anomaly experiment

Delay measured at the 10<sup>th</sup> line of receivers in three frequency bands, 3-4Hz, 7-8Hz, 12-13Hz



Displacement versus strain



Displacement versus rotation

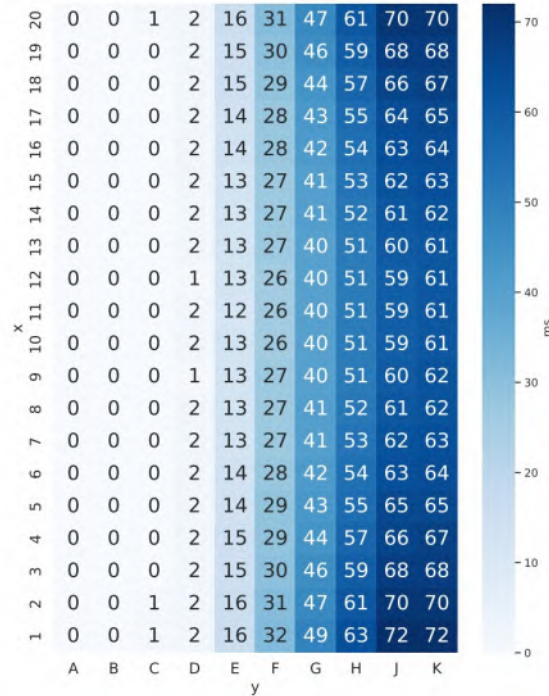


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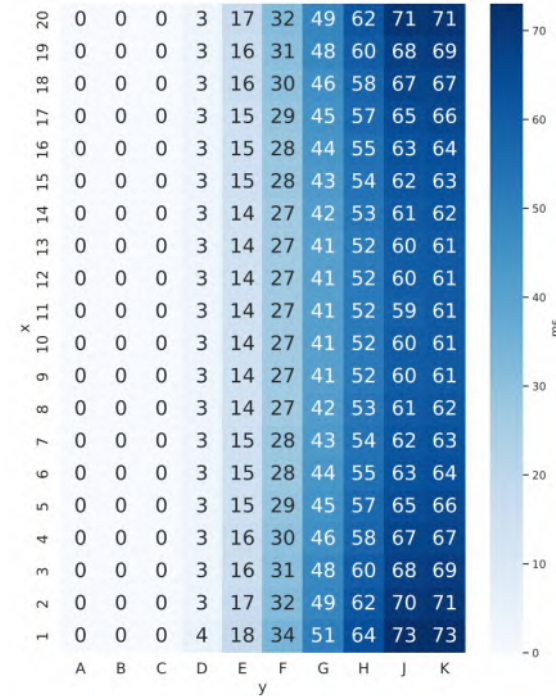


# Observables sensitivity comparison

Delay recorded at each receiver of the array (3-4Hz)



Displacement  $x$



Rotation  $\omega_y$

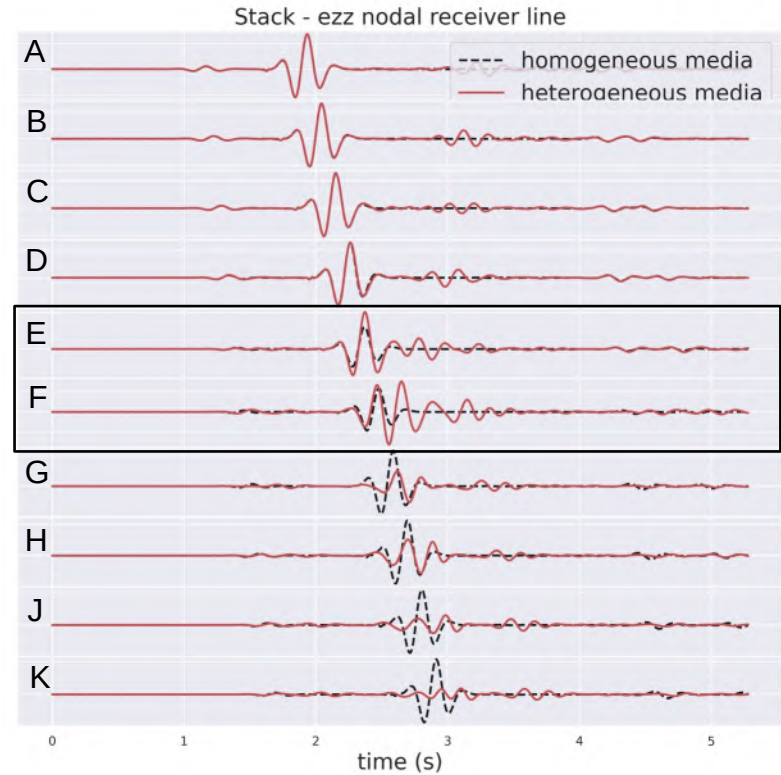
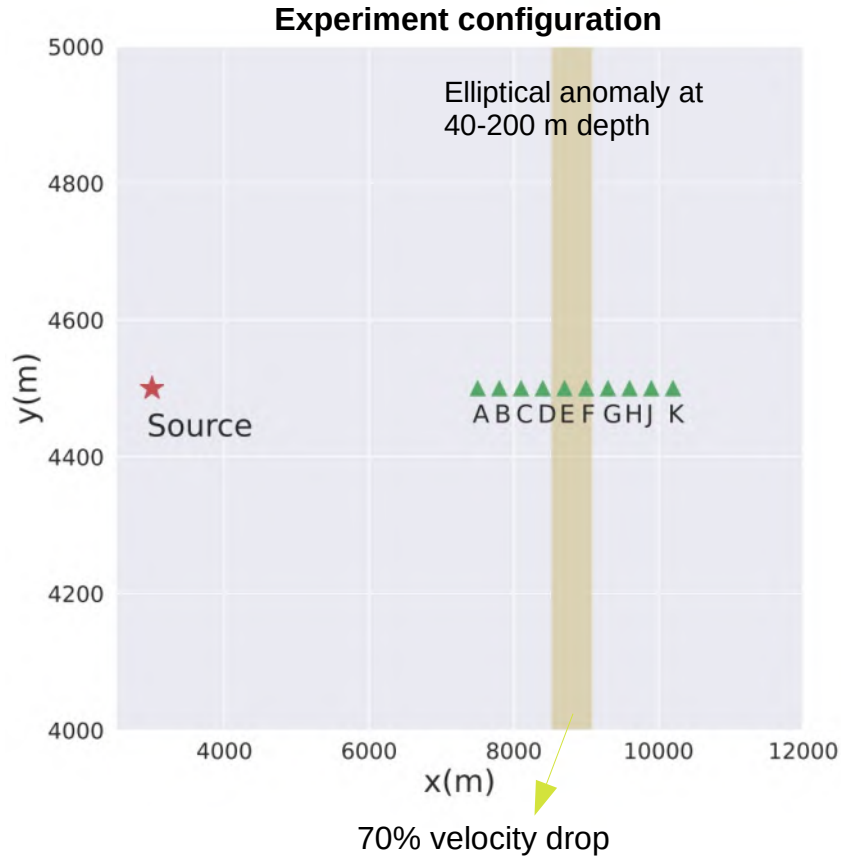


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# Strong anomaly experiment

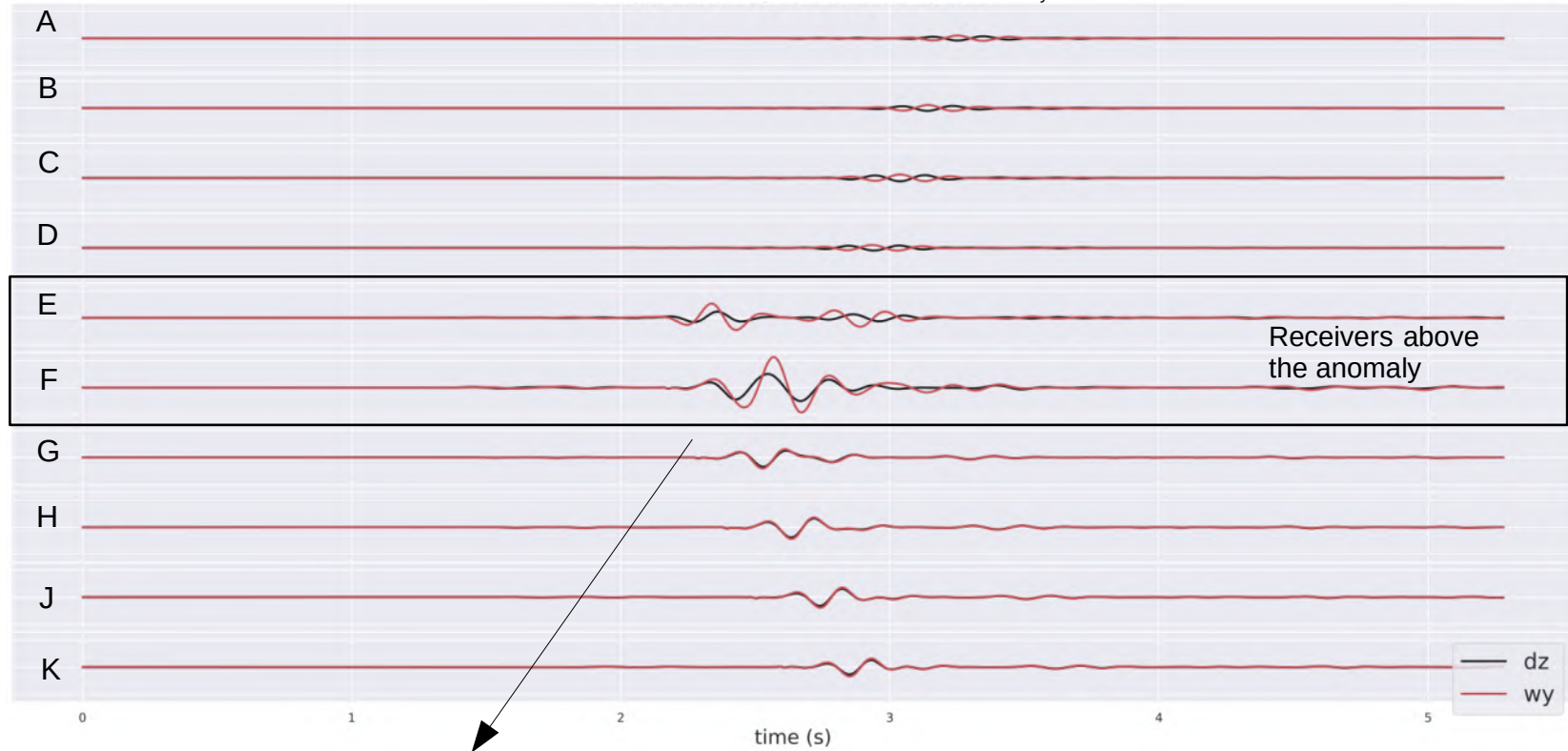


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# Strong anomaly experiment

Comparison of normalized scattered field  $dz/dt - \omega_y$  at the array



Local amplification of the rotation



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# Conclusions

## Weak anomaly experiment

For the purposes related to the seismic velocity monitoring, moderate changes up to 10% of velocity, the wavefield gradients do not show an increased sensitivity compared with the traditional observables

## Strong anomaly experiment

- Phase measurements are not effective
- Local amplification of the gradients

The larger sensitivity of the rotations could be used for FWI purposes

